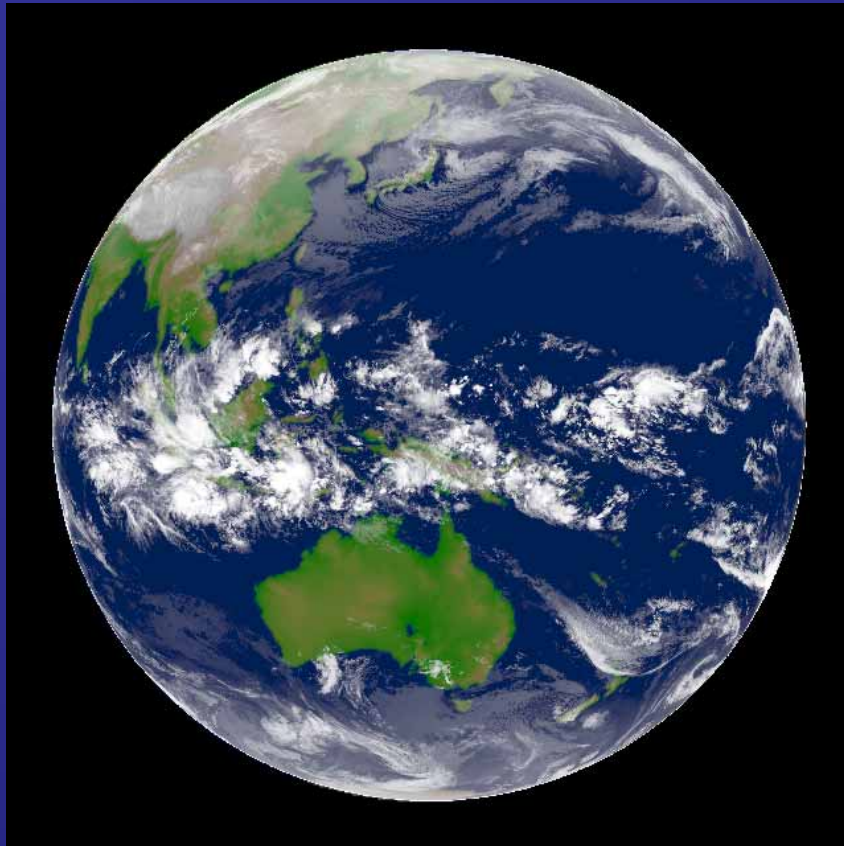
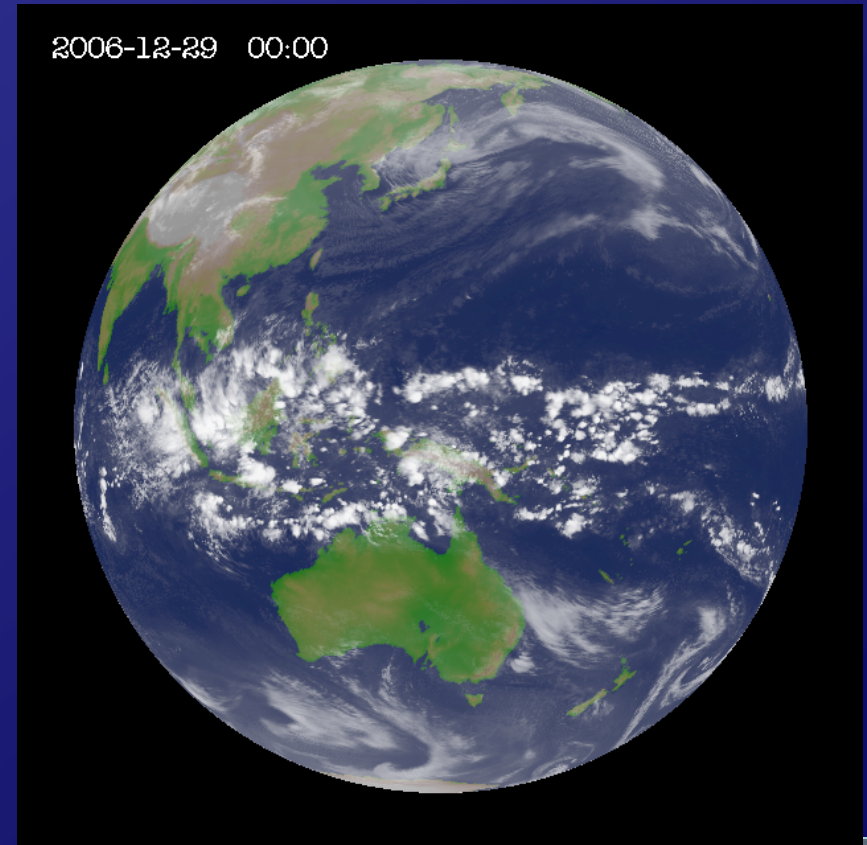


Influence of topography and zonal SST gradient on the MJO

Observation: TBB from MTSAT-1R



Simulation: OLR from a 3.5-km grid run



Hiroaki Miura (FRCGC, JAMSTEC)



• Dynamics (grid-scale)

Governing equations	Full compressible non-hydrostatic system
Spatial discretization	2 nd -order centered scheme (Tomita et al. 2002) 2 nd -order upwind biased scheme (Miura, MWR accepted)
Horizontal grid configuration	Icosahedral grid
Vertical grid configuration	Lorenz grid
Topography	Terrain-following coordinate
Conservation	mass, tracers, total energy (Sato 2003)
Temporal scheme	Slow mode - explicit scheme (RK2) Fast mode - Horizontal Explicit Vertical Implicit scheme

• Physics (subgrid-scale)

Turbulence / surface flux	Modified Mellor & Yamada → Yamada and Mellor (1979) /Louis(1979), Uno et al.(1995)
Radiation	MSTRNX (Sekiguchi 2004)
Cloud physics	Grabowsky(1998)
Cloud parameterization	no
Shallow clouds	no
Land process	Mixed layer/bucket



Initial conditions:

Interpolated from NCEP tropospheric analyses (6 hourly, 1.0x1.0 degree grids)

Initial time:

14-km run: 2006-12-15 00:00:00

Boundary conditions:

Reynolds SST, Sea ICE (weekly data): These are linearly interpolated in time.

ETOPO-5 topography, Matthews vegetation

UGAMP ozone climatology (for AMPI2)

Horizontal grid spacing:

14 km

Vertical domain:

0 m ~ 38,000 m

40-levels (stretching grid)

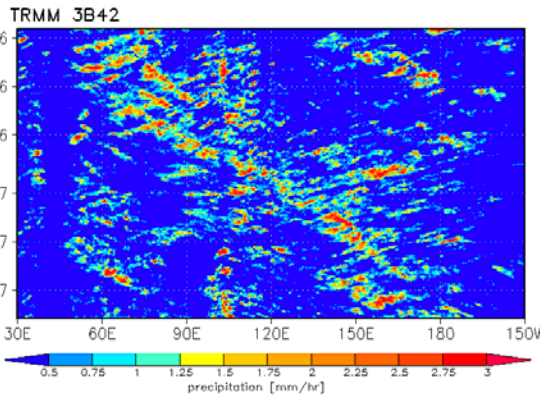
Duration (current situation):

14-km run: 30 days

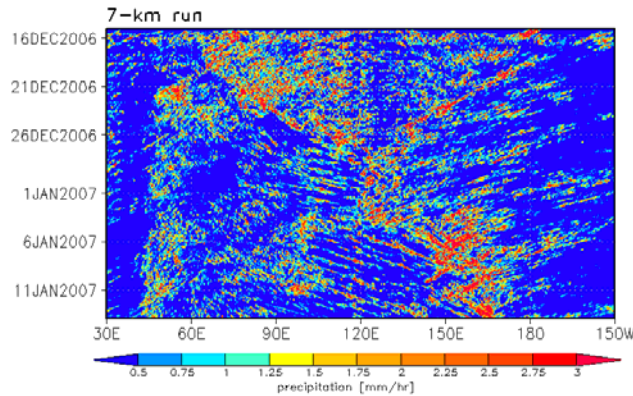


Precipitation: lat=5S-5N

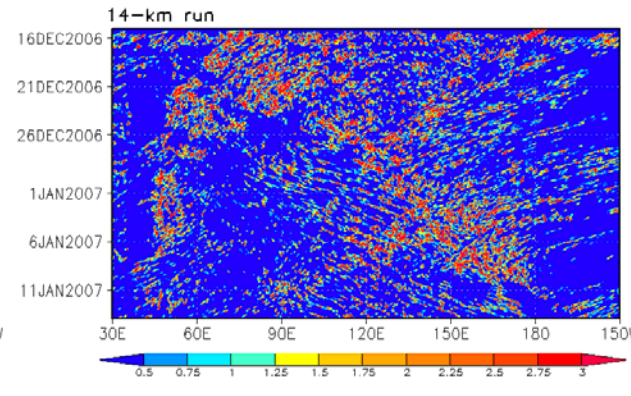
TRMM 3B42



7-km run

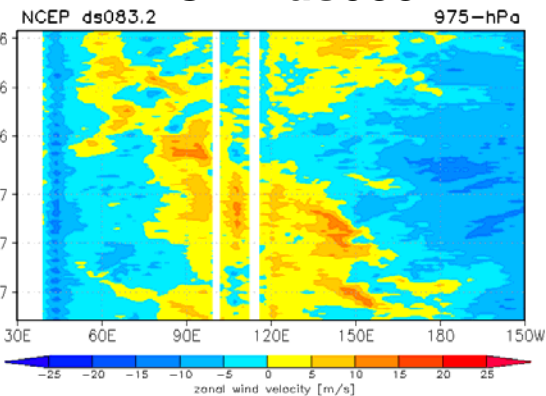


14-km run

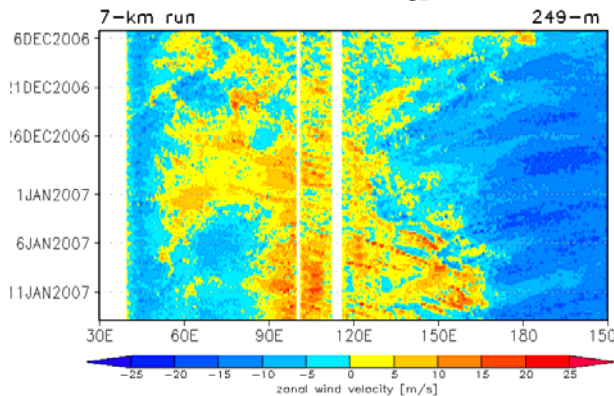


Zonal wind: lat=0, p=975 hPa/z=249 m

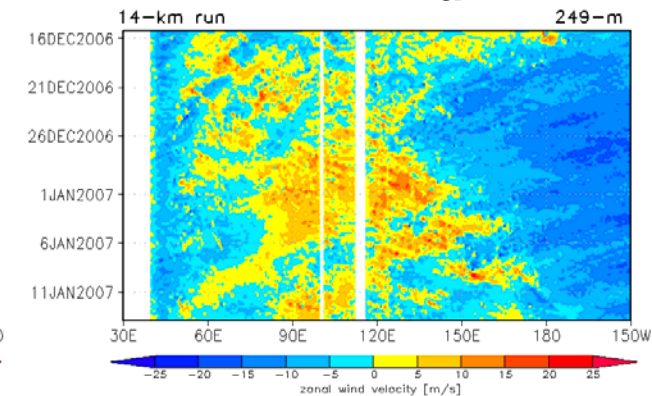
NCEP ds083.2



7-km run



14-km run



Time change of zonal wind

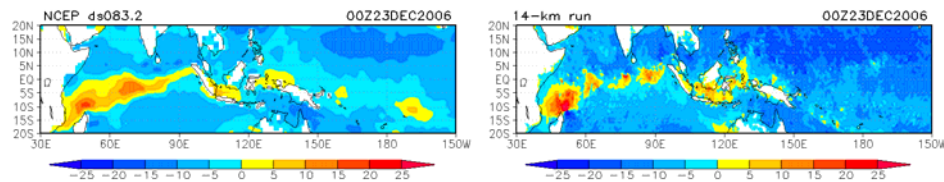
NCEP ds083.2: p=975 hPa

14-km run: z=249 m

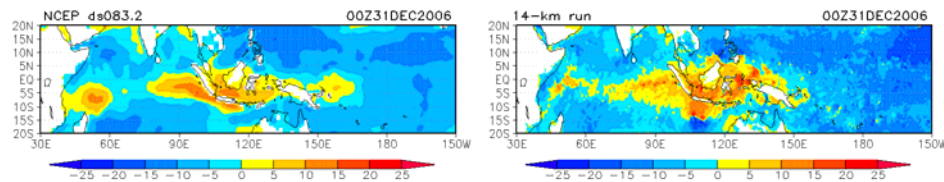
NCEP ds083.2:lat=0

14-km run:lat=0

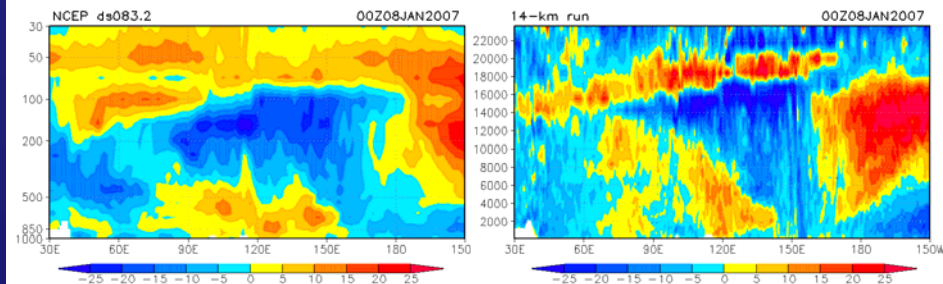
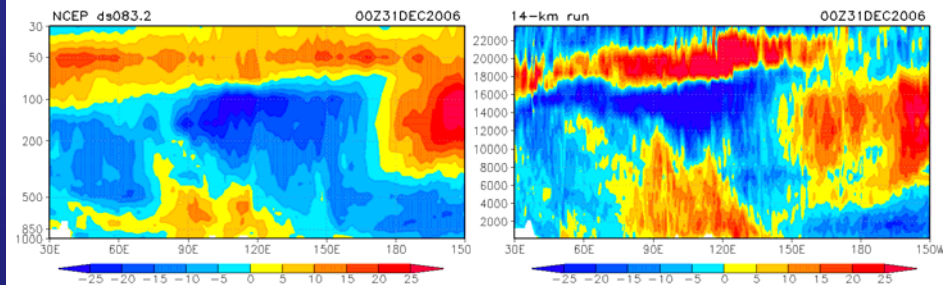
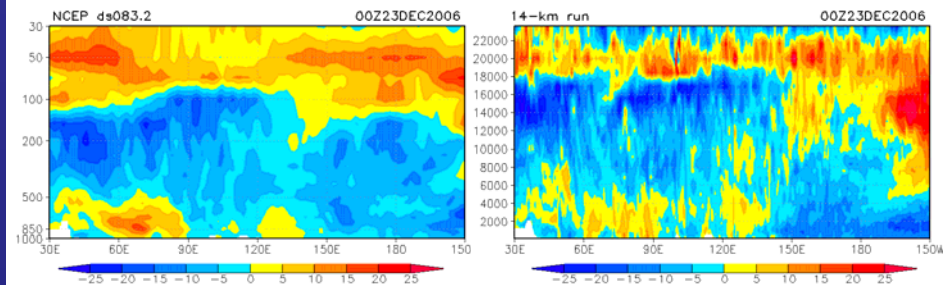
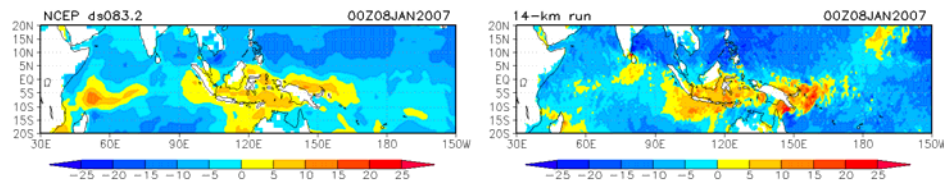
2006-12-23



2006-12-31



2007-01-08



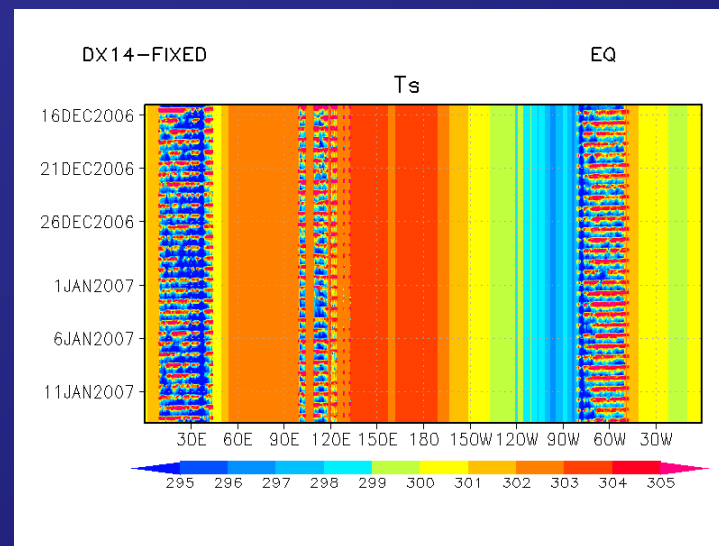
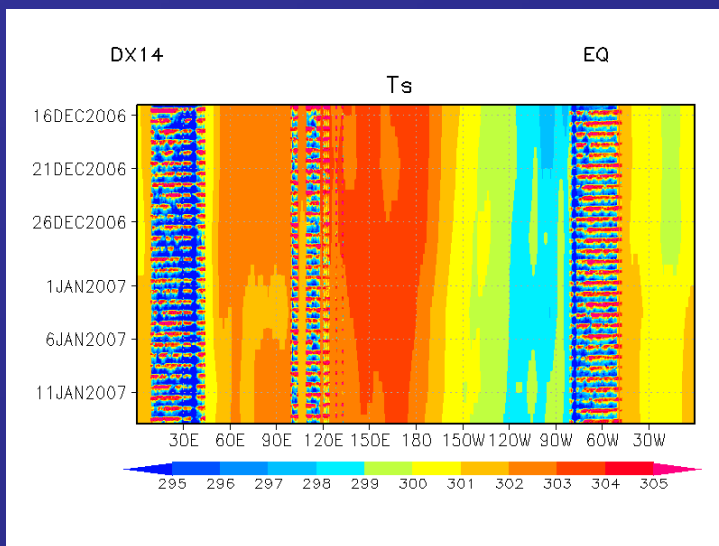
Test-1: No time change of SST

FIXED: SST was fixed to data at 2006-12-13.

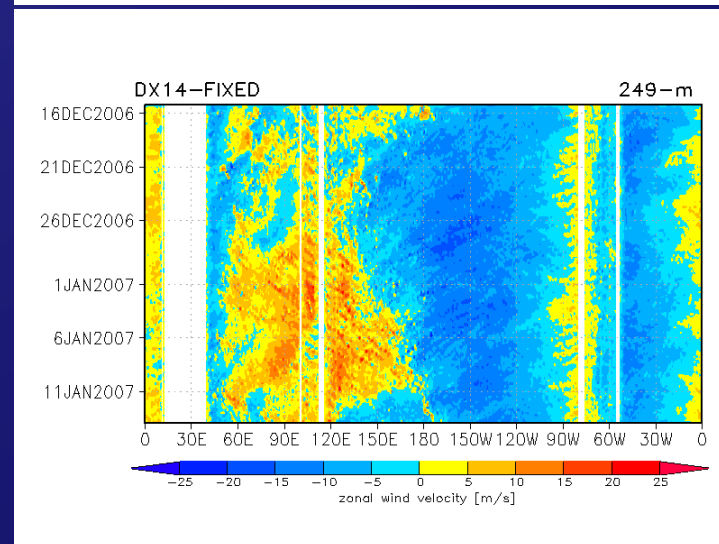
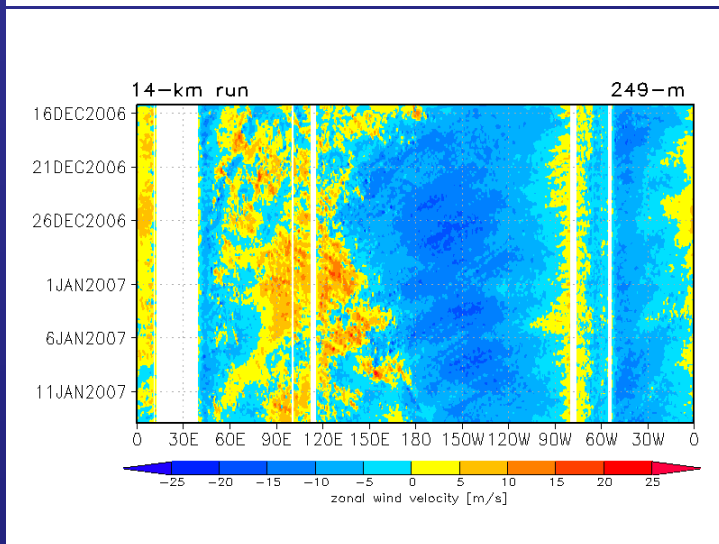
CTL (Control)

FIXED

Surface
temperature:
lat=0



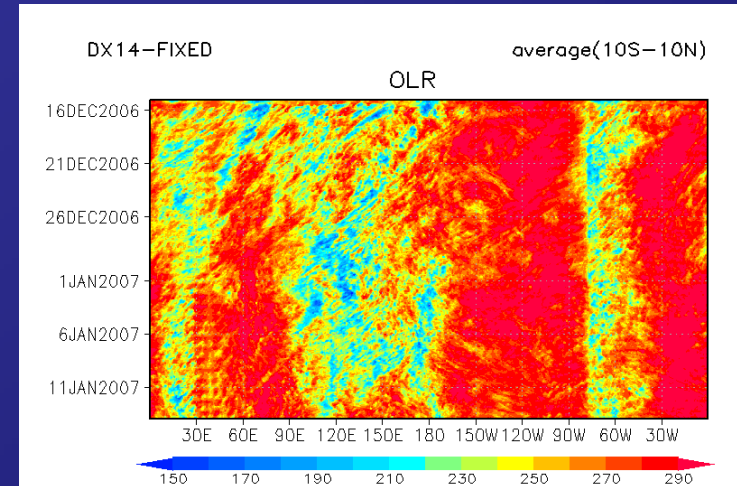
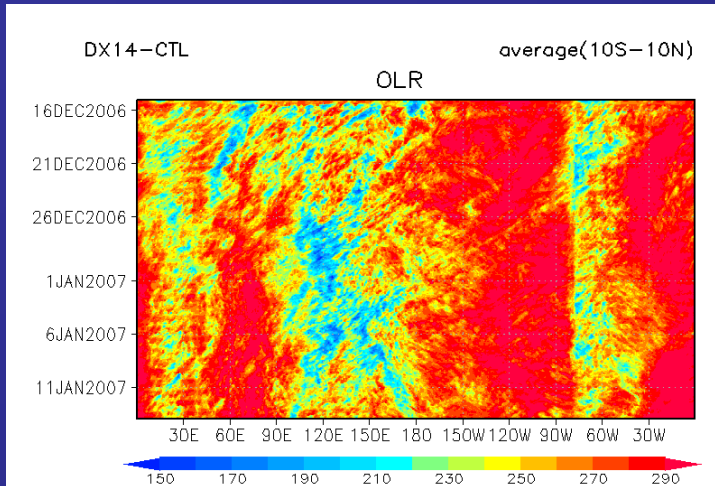
Zonal wind:
lat=0



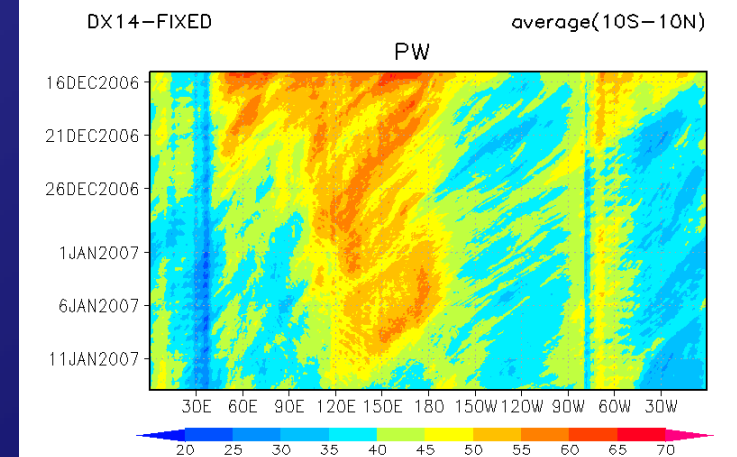
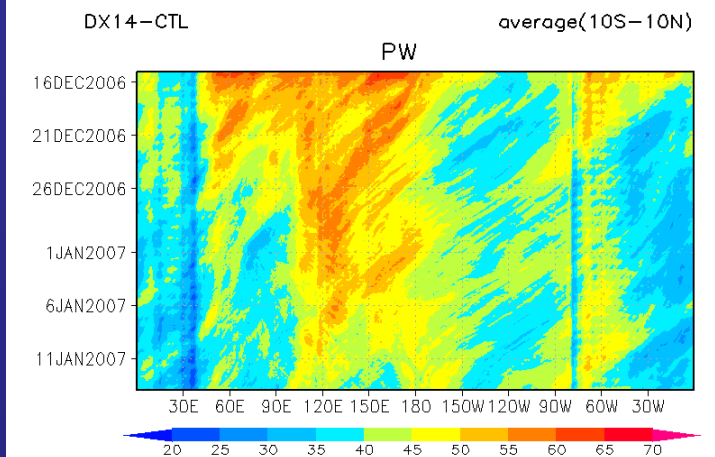
CTL (Control)

FIXED

Outgoing
long-wave
radiation:
lat=10S-10N



Precipitable
water:
lat=10S-10N



- At least in the simulations, the SST feedback was not essential for the slow eastward movement of the MJO.



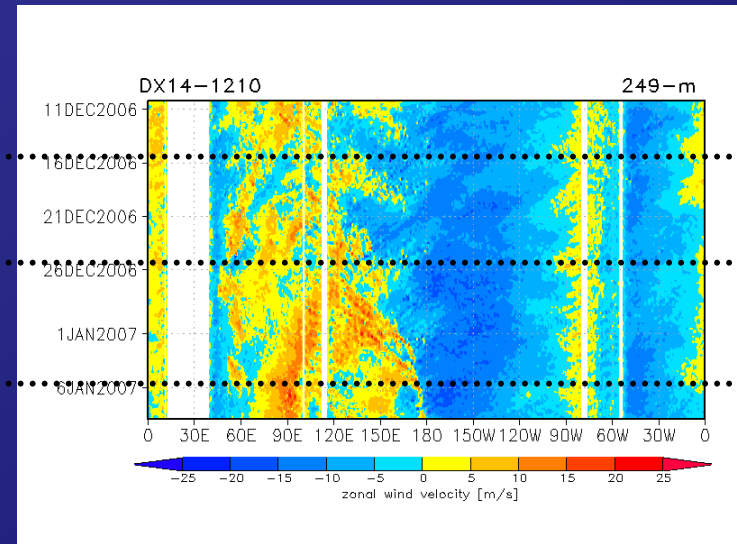
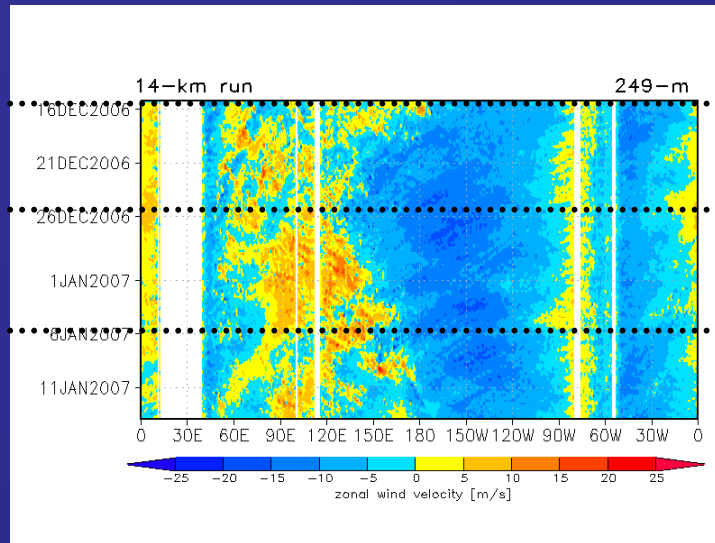
Test-2: Another initial date

1210: Initial conditions were given at 2006-12-10.

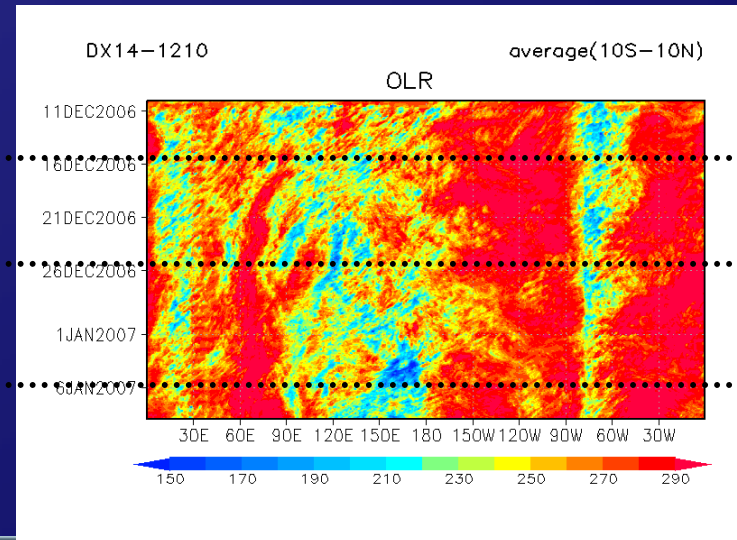
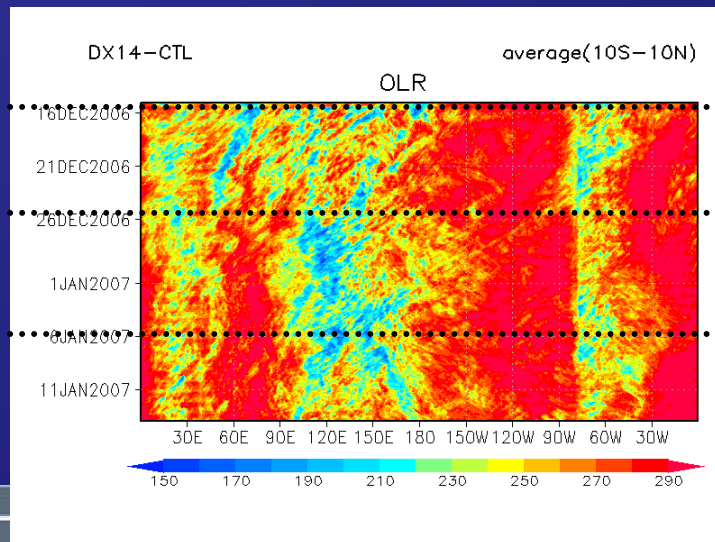
CTL (Control)

1210

Zonal wind:
lat=0

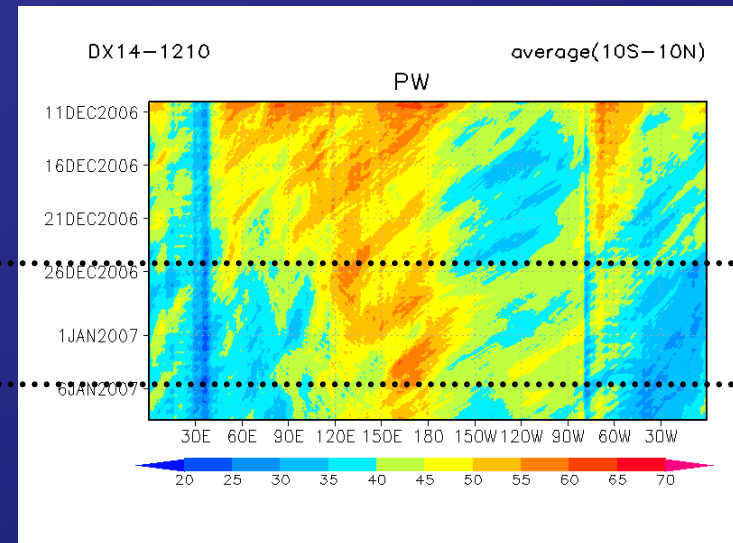
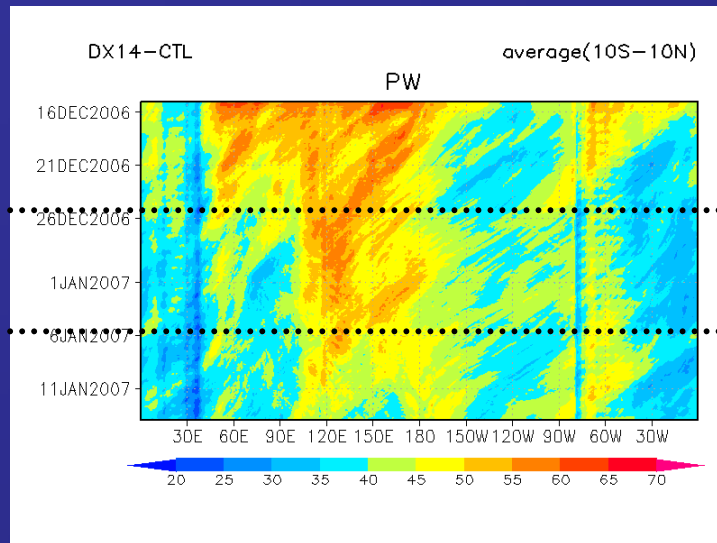


Outgoing
long-wave
radiation:
lat=10S-10N



CTL (Control)

1210



Precipitable
water:

lat=10S-10N

- Convective activity over the maritime continents in 1210 was enhanced at the same timing as in CTL.
- However, eastward movement was faster in 1210 than in CTL.
- There were smaller amount of moisture over the maritime continents.



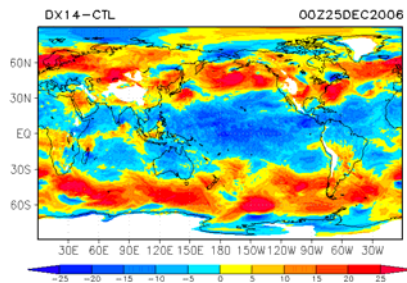
Test-3: Influence of topography

TOP00: Topographic height was zero over the globe.

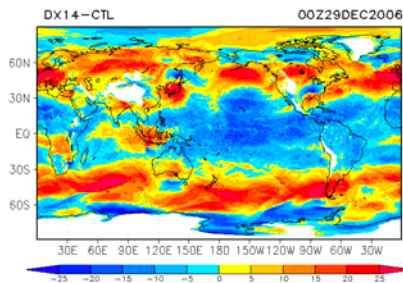
Time change of zonal wind: z=1570 m

CTL (control)

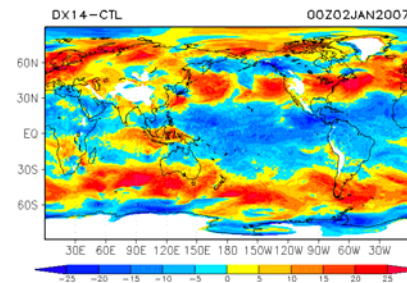
2006-12-25



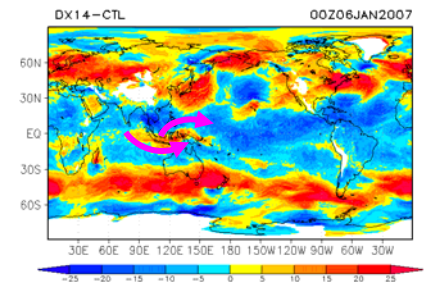
2006-12-29



2007-01-02

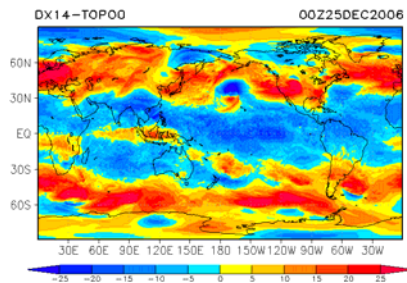


2007-01-06

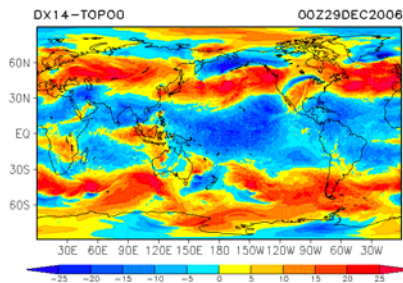


TOP00

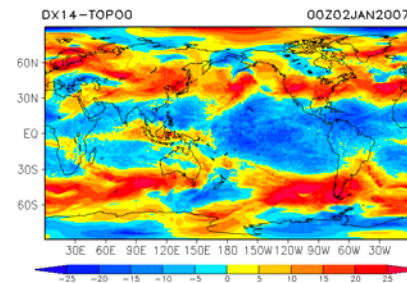
2006-12-25



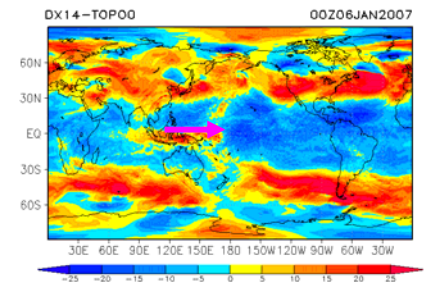
2006-12-29



2007-01-02



2007-01-06



Time change of zonal wind: lat=-5

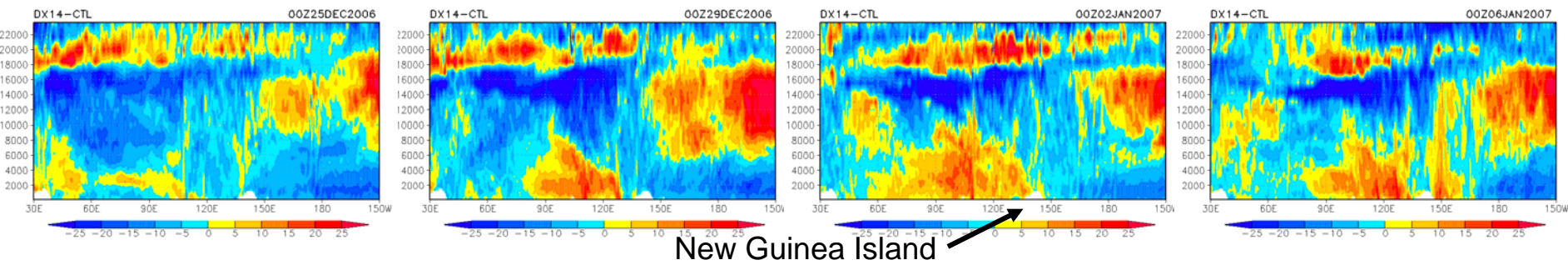
CTL (control)

2006-12-25

2006-12-29

2007-01-02

2007-01-06



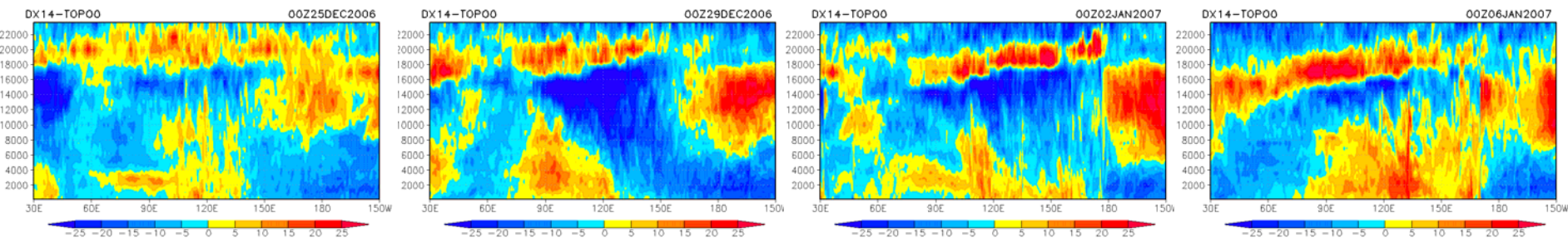
TOP00

2006-12-25

2006-12-29

2007-01-02

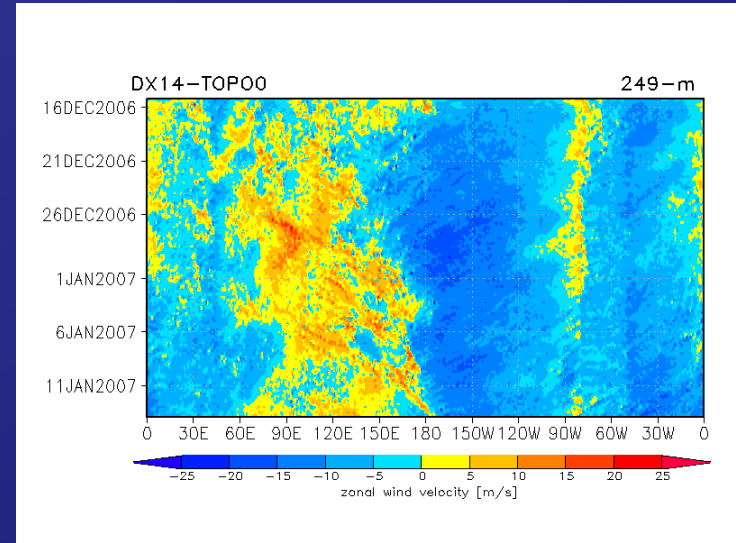
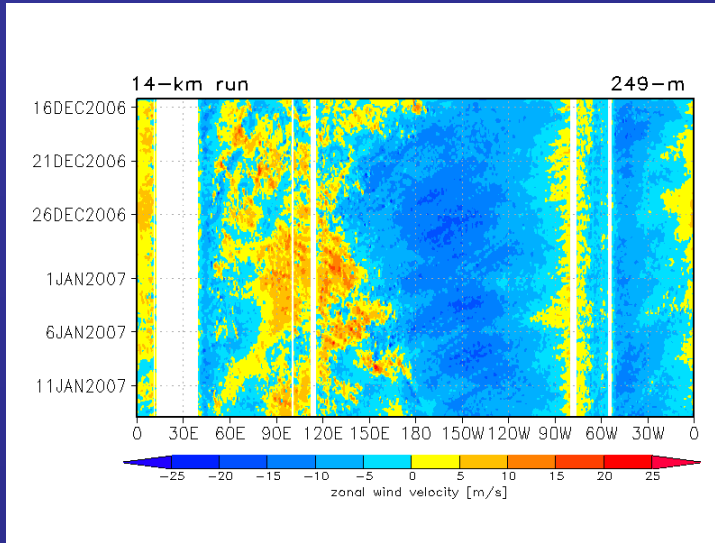
2007-01-06



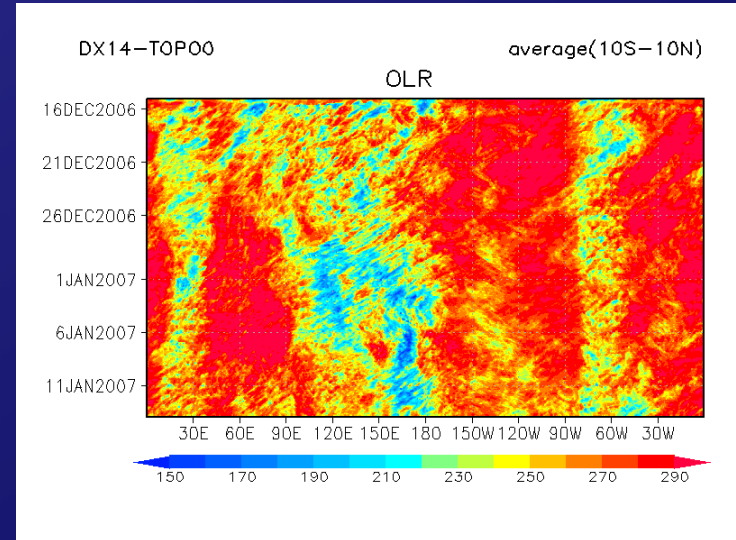
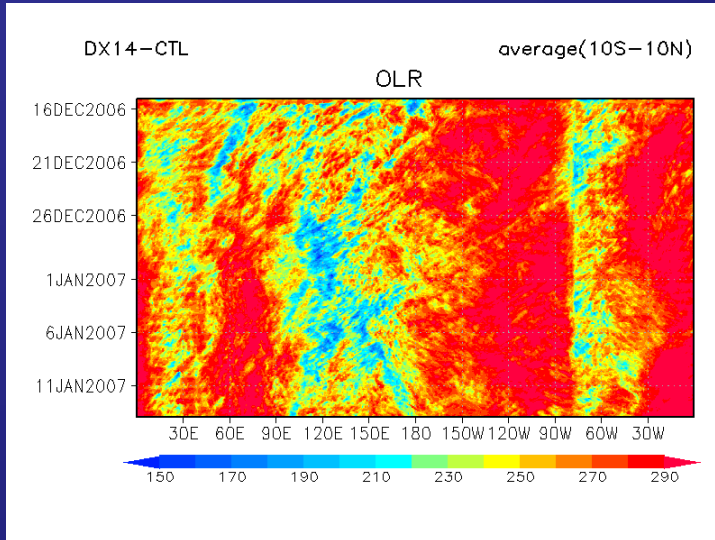
Control

TOPO0

Zonal wind:
lat=0

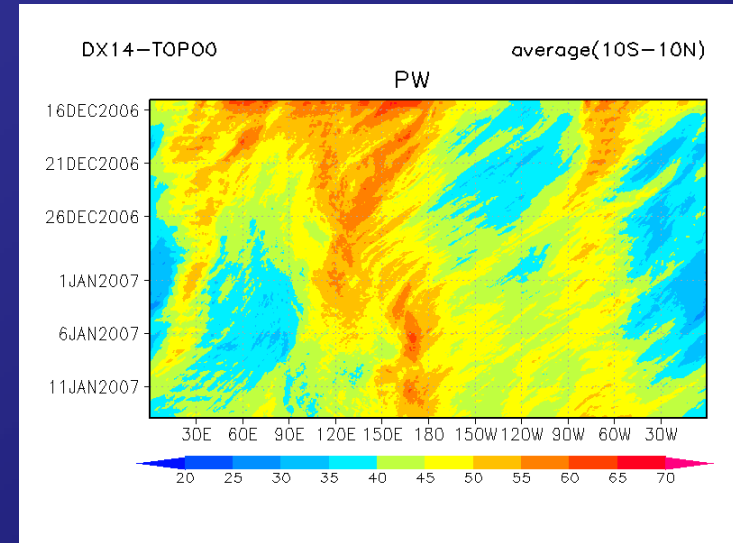
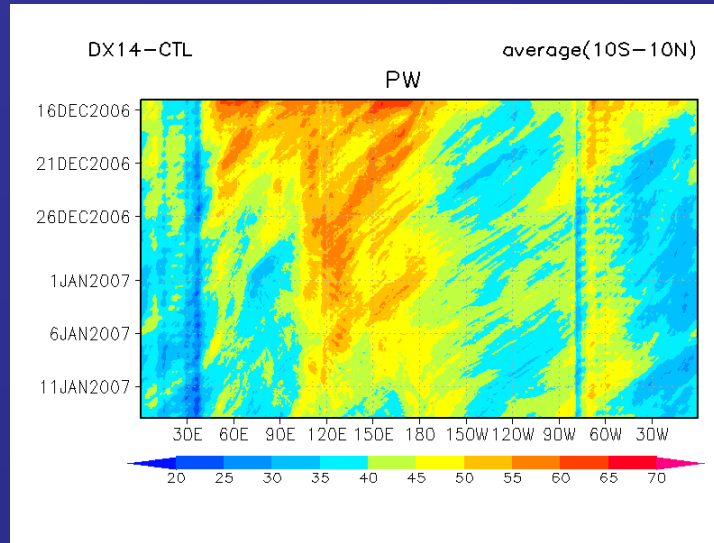


Outgoing
long-wave
radiation:
lat=10S-10N



Control

TOPO0



Precipitable
water:
lat=10S-10N

- The eastward moving signal (OLR) was separated into two signals in TOPO0.
- A fast propagating signal (~15 m/s) and a localized signal (at 120E).
- The fast propagating signal in CTL was apparently blocked by topography.
- Influence of mid-latitude signals seemed small, but it is not confirmed now.



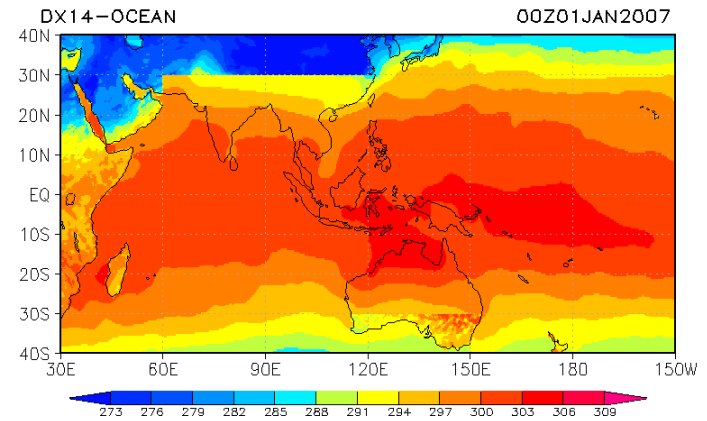
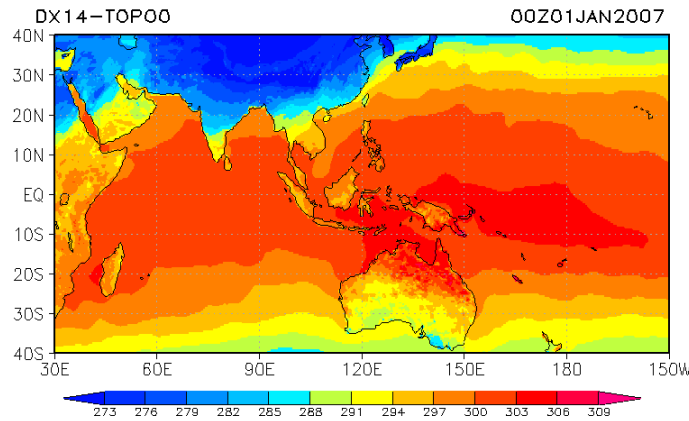
Test-4: Influence of land surface

OCEAN: TOPO0+surface condition was ocean over 30S-30N and 60E-180.

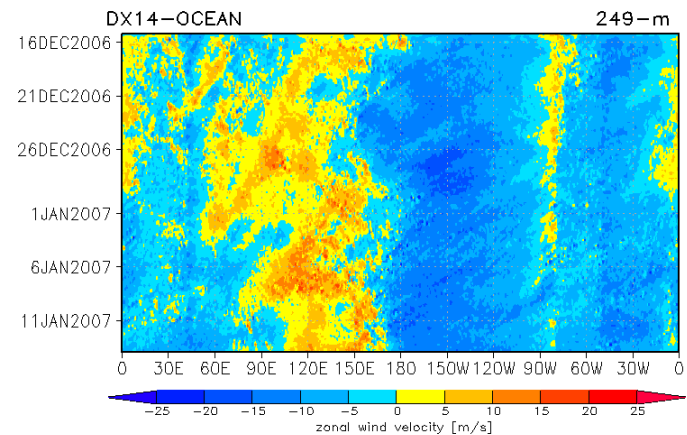
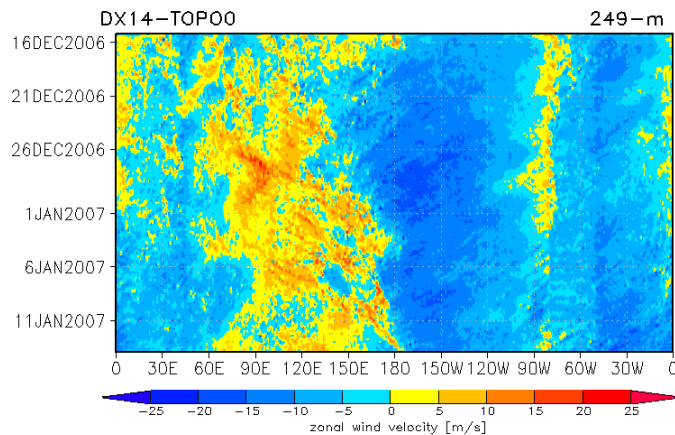
CTL (Control)

OCEAN

Surface
temperature:
2007-01-01



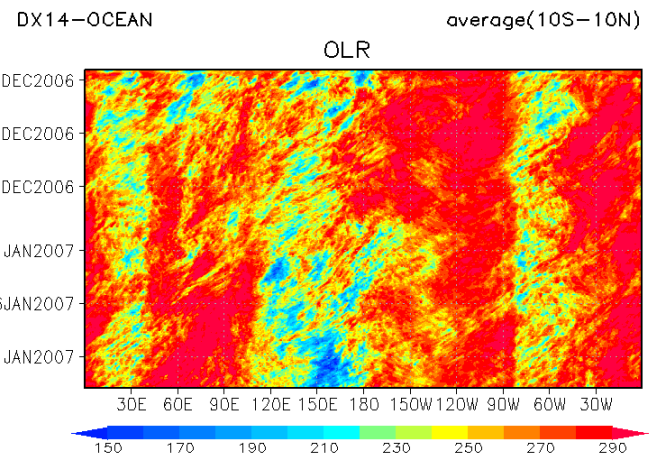
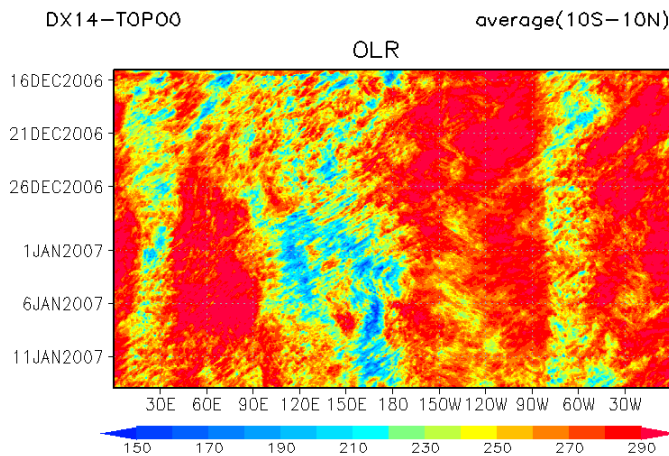
Zonal wind:
lat=0



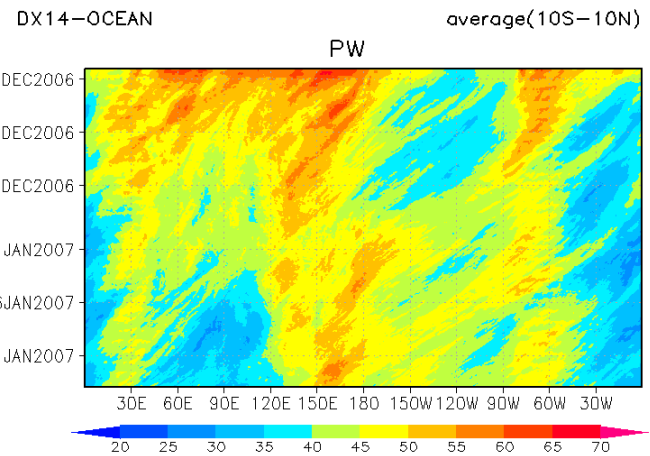
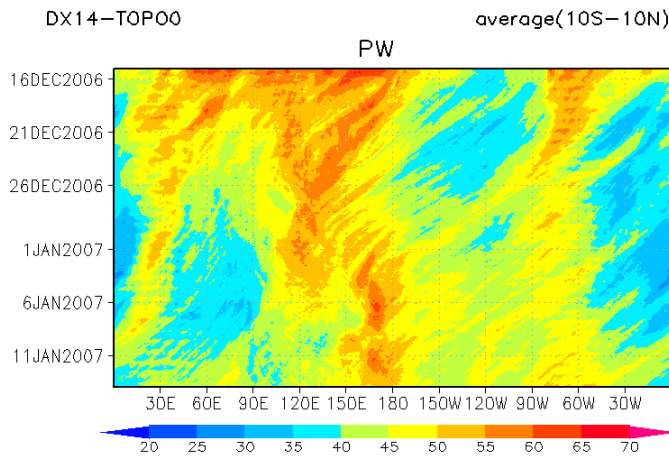
CTL (Control)

OCEAN

Outgoing
long-wave
radiation:
lat=10S-10N



Precipitable
water:
lat=10S-10N



- Land surface (surface heating) seemed to help fast organization of convection.



Test-5: Influence of zonal SST gradient

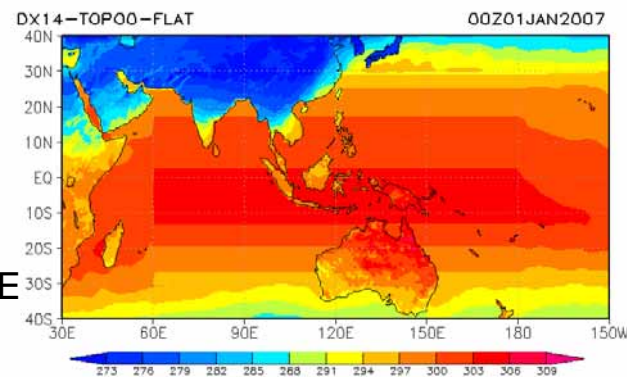
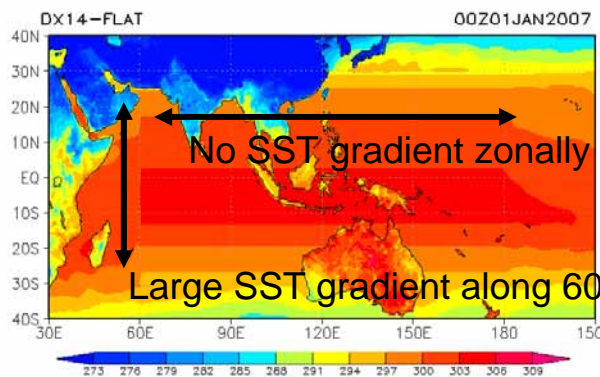
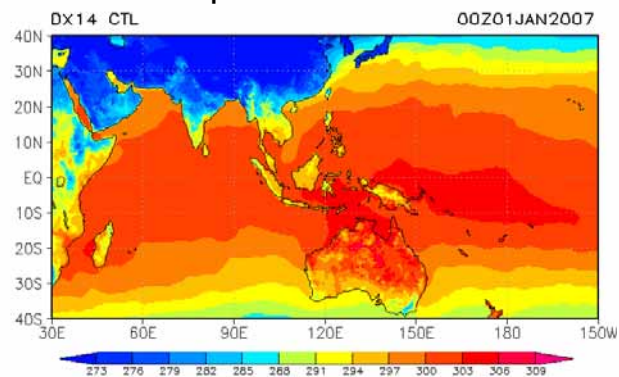
FLAT: SST over 30S-30N and 60E-180 was the same as SST of lon=180.

CTL (Control)

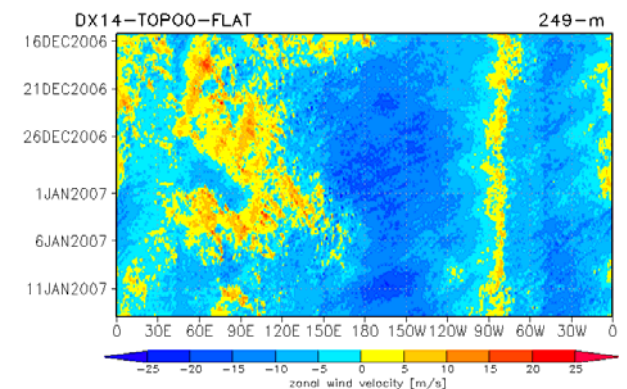
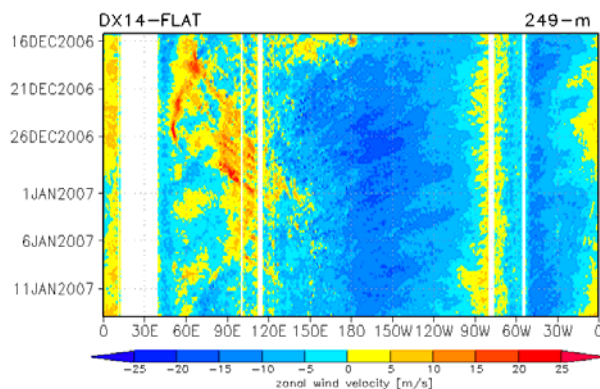
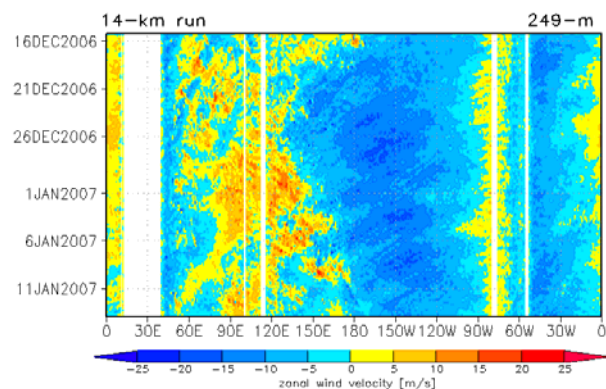
FLAT

TOPO0-FLAT

Surface temperature: 2007-01-01



Zonal wind: lat=0

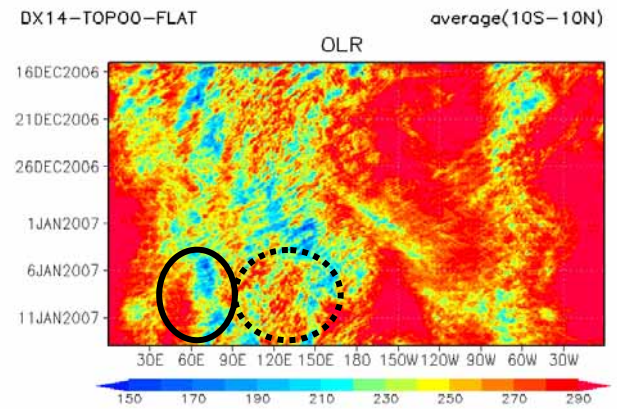
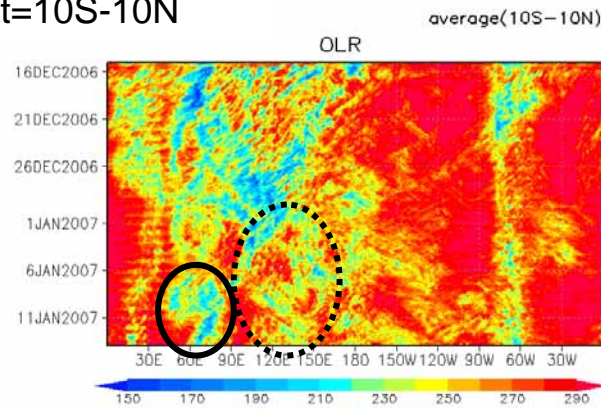
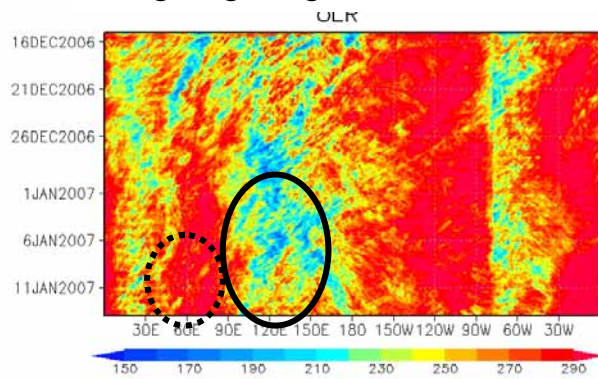


CTL (Control)

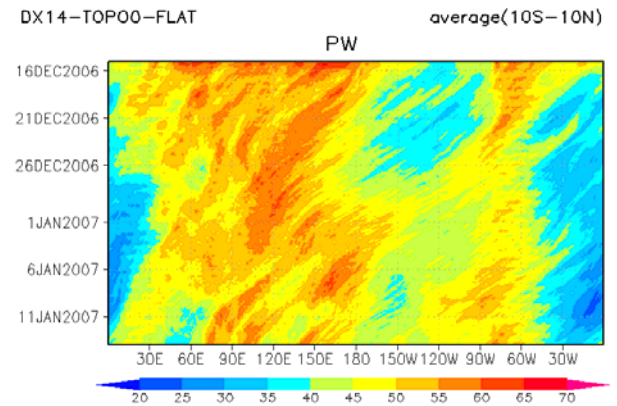
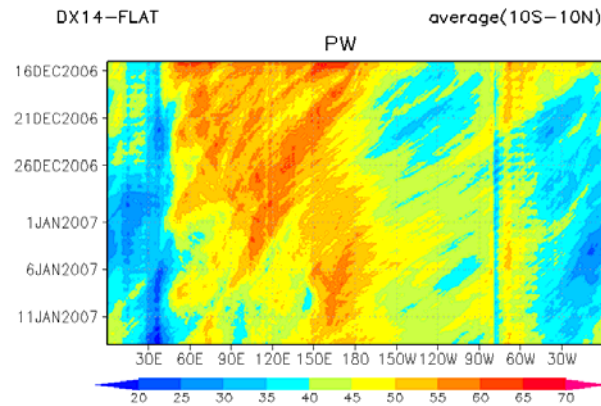
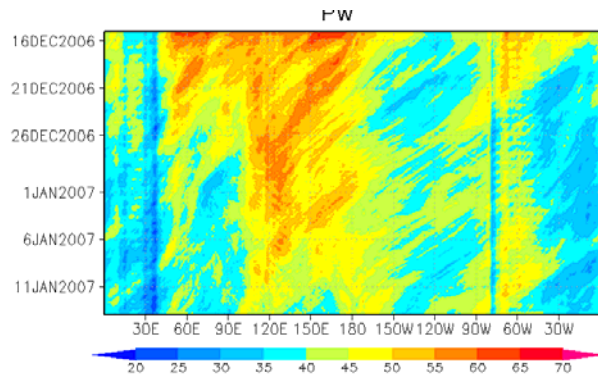
FLAT

TOPO0-FLAT

D Outgoing long-wave radiation: lat=10S-10N



D Precipitable water: lat=10S-10N



- Convection over the maritime continents was not maintained in FLAT cases.
 - Moisture passed through over the maritime continents.
 - Convection was enhanced around 60E in FLAT cases.
- A eastward propagating signal in TOPO0-FLAT was likely a single wave.



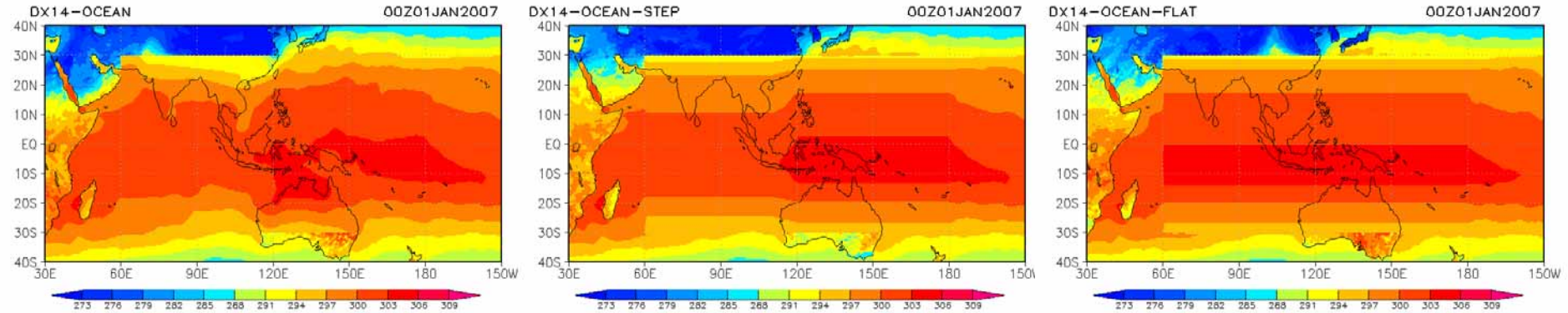
**STEP: SST was 1 K cooler than SST of FLAT over 30S-30N and 60E-110E.
SST(110E<lon<120E) was interpolated zonally and linearly.**

OCEAN

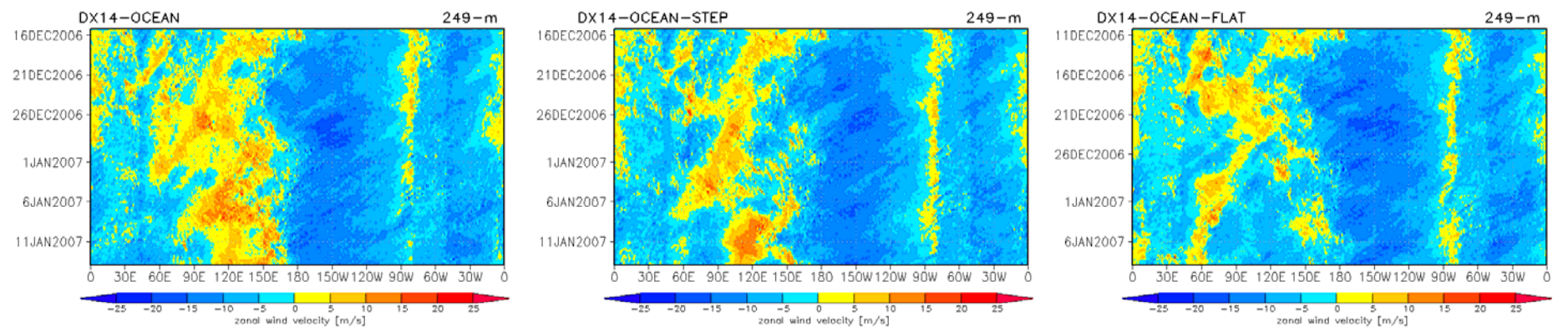
OCEAN-STEP

OCEAN-FLAT

Surface temperature: 2007-01-01



Zonal wind: lat=0



OCEAN

OCEAN-STEP

OCEAN-FLAT

DX: Outgoing long-wave radiation: lat=10S-10N

average(10S-10N)

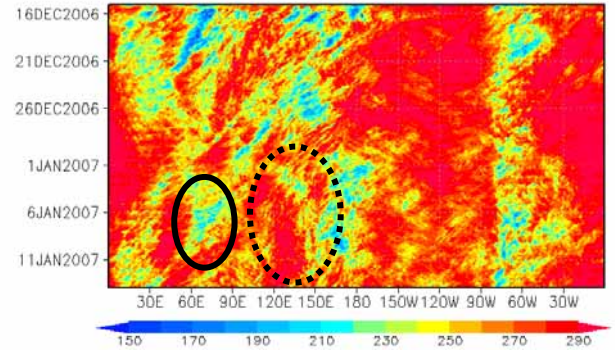
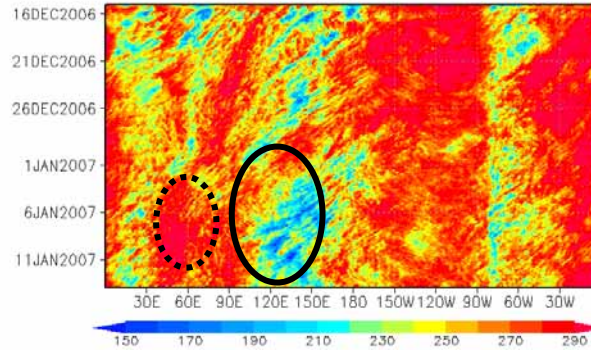
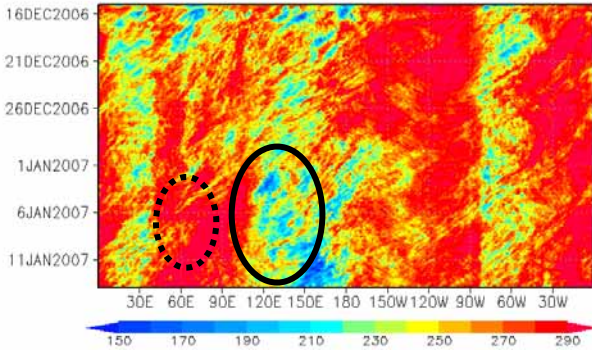
DX14-OCEAN-FLAT

average(10S-10N)

OLR

OLR

OLR



DX: Precipitable water: lat=10S-10N

14-OCEAN-STEP

average(10S-10N)

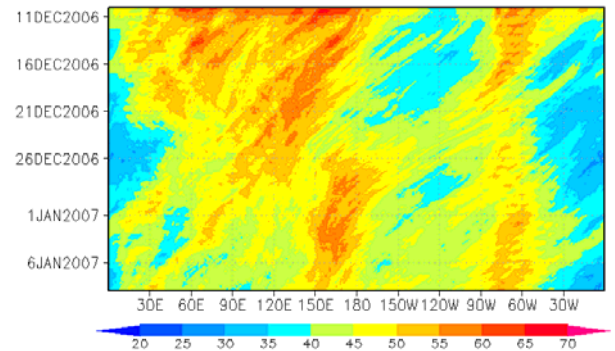
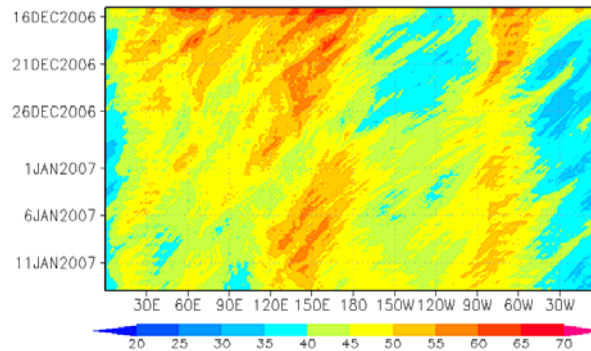
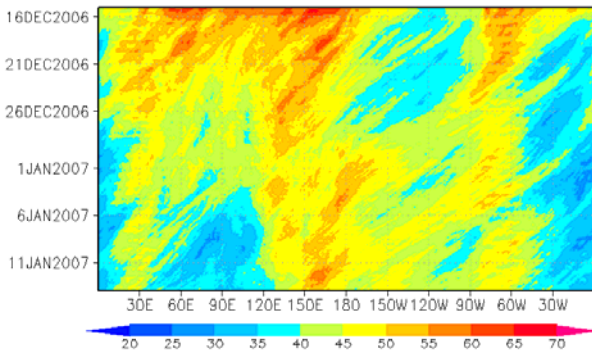
DX14-OCEAN-FLAT

average(10S-10N)

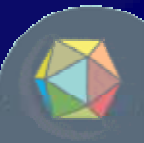
PW

PW

PW



- Convection over the maritime continents was maintained in OCEAN-STEP, but it was not simulated in OCEAN-FLAT.
- Which is important?
 - local equilibrium, meso-scale circulation, and lifting of warmer air over cooler air.



- Some sensitivity tests were done using a 14-km grid.
- Influence of SST variation
 - The SST feedback was not essential for the slow eastward movement of the MJO at least in this event.
- Influence of topography
 - There was a westerly signal which proceeded over the maritime continents at the end of 2006.
 - The westerly signal seemed to be blocked by topography.
 - And, the other westerly signals took over a role in the eastward movement of the MJO.
- Influence of zonal SST gradient
 - Convective activity over the maritime continents could not be maintained without zonal SST gradient.
 - Convection was forced around the region where zonal SST gradient was large.

