Global Chemical Transport in a Cloud-Resolving GCM Rosa, D., Collins, W.D., & Lamarque, J.F.

✓ MOTIVATION:

✓ In models, the treatment of cloud convective processes affects the atmospheric vertical transport and mixing therefore the chemical state and radiative forcing of the atmosphere. We investigate the implications for the concentrations of short lived atmospheric passive tracers, radon (Rn) and methyl iodide (CH3I) from explicitly resolving cloud physical and dynamical processes. Conventional climate models use implicit treatments for cloud convective processes called "parameterizations"; we use one of those parameterizations here as control case. To test the fidelity of the model we compare simulated v. measured vertical profiles available from field campaigns

✓ CONTROL CASE:

- ✓ Community Atmosphere Model (CAM) in Chemical Transport Mode (CTM) with NCEP meteorology: 28 levels; 1.9x2.5; FV; dt=30min
- ✓ Zhang & McFarlane convective scheme for clouds
- ✓ Model for Ozone and Related Chemical Tracers (MOZART)
- ✓ Rn from land masses
- ✓ CH3I monthly climatologies (Bell et al JGR). Marine areas and rice paddies.

✓ TEST CASE:

✓ As in CONTROL but the cloud convective processes are from the Colorado State University (CSU) Multi-scale Modeling Framework (MMF): SUPER-PARAMETERIZED CAM (SPCAM)

✓ Embedded 2D Cloud Resolving Model (CRM): 64x1; dx = 2000m; t=20s

✓ RUNS:

✓ 11 Years: 1996 → 2006

✓ PROFILE BINNING:

✓ Observed and simulated values are binned on discrete vertical levels with midpoints at 0.5, 2, 5, 10, and 18 Km.

✓ LOCAL CONVECTIVE INDEX (LCI):

- ✓ LCI = $\Delta ln(q)/\Delta h$ where q is volume mixing ration and h is height
- ✓ The anomalies in the simulated LCI with respect to observations are calculated for the atmospheric layers referred to here as the Planetary Boundary Layer (**PBL**; 0.5–2 Km), Low Troposphere layer (**LTL**; 2–5 Km), Mid Troposphere layer (MTL; 5–10 Km), and Upper Troposphere layer (UTL; 10–18 Km).

✓ CONCLUSIONS

✓ Simulated vertical profiles of tracers concentrations from both models are compared with observed ones and suggest that the cloud-resolving GCM is better than the conventional GCM. Contrasting climatological maps of tracers concentrations from simulations we found consistent and appreciable relative differences between the cloud-resolving GCM and the conventional case that might have important implications for climate and atmospheric chemistry simulations but require further investigations.



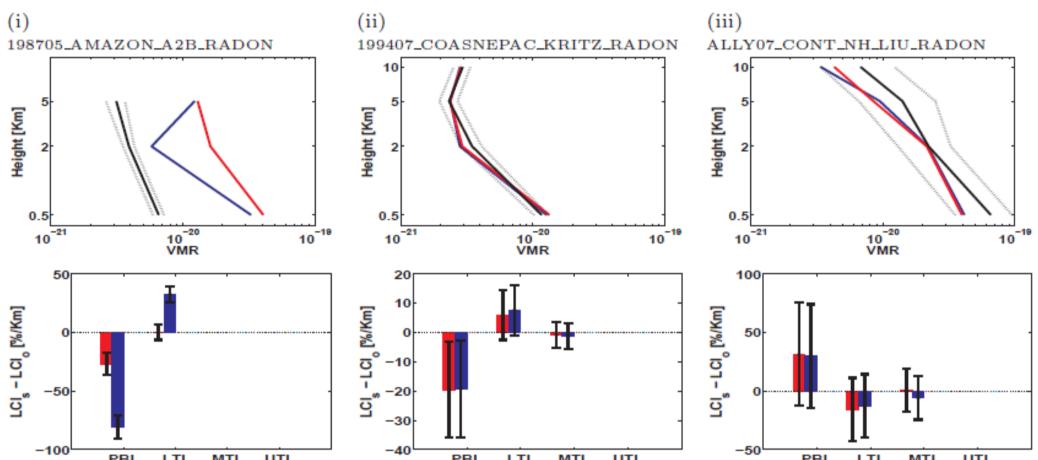


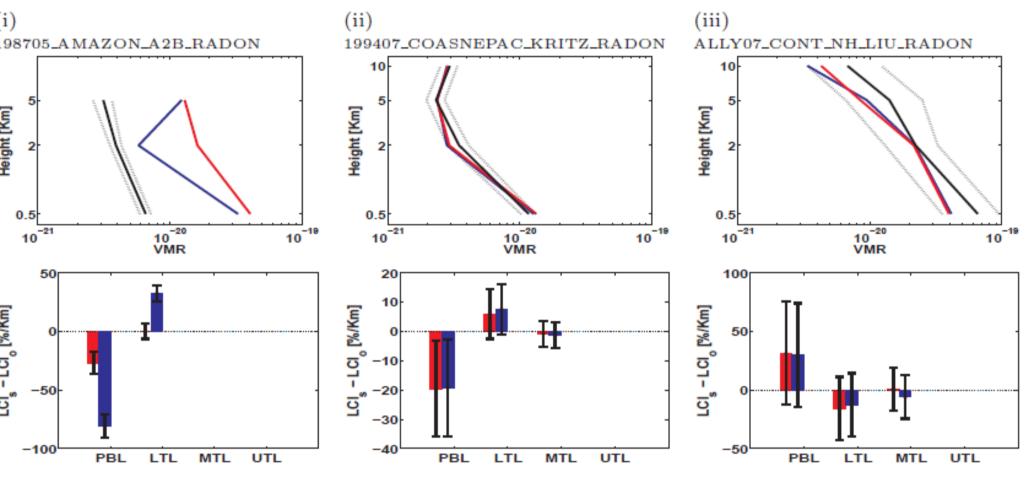


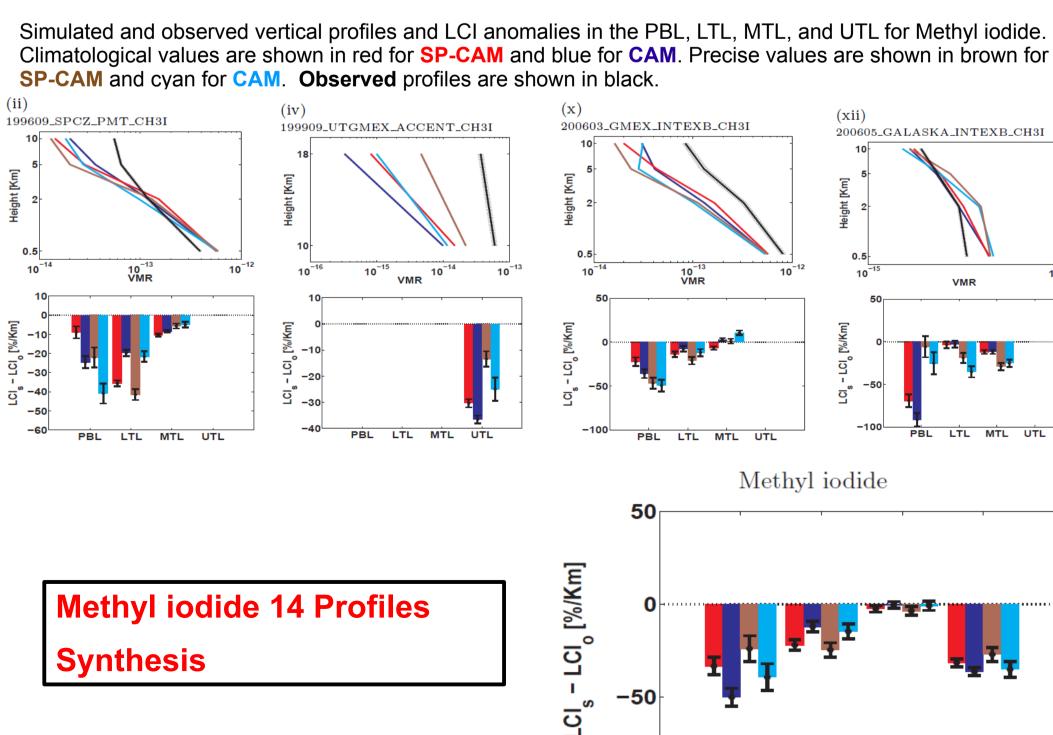


199609_SPCZ. 199609_ITCZ 199904_TROI 199909_UTG1 200103_SUBCH 200103_SUBW 200401_UTPA 200407_EUS_ 200601_UTPA 200603_GMEX 200604_SUBT 200605_GALA 200806_SFOB 200807_CANA 198705_AMA2 199407_COAS ALLY07_CON

Simulated and observed vertical profiles and LCI anomalies in the PBL, LTL, MTL, and UTL for Rn. Climatological values are shown in red for SP-CAM and blue for CAM. Observed profiles are shown in black.







Case Studies

	Tracer	Pts	Period	Lat×Lon [°N×°E]	Alt. [Km]	Mission	Reference
PMT	CH ₃ I	2242	(1996-08-31,1996-10-05)	$(-50,15) \times (102,310)$	(0,11)	PEM-Tropics A	Hoell et al. [1999]
PMT	CH3I	533	(1996-08-21,1996-10-05)	$(-15,35) \times (137,277)$	(0,10)	PEM-Tropics A	Hoell et al. [1999]
P_PAC_PMB	CH ₃ I	1612	(1999-03-06,1999-04-13)	$(-49,35) \times (144,275)$	(0, 12)	PEM-Tropics B	Raper et al. [2001]
MEX_ACCENT	CH ₃ I	135	(1999-09-03,1999-09-20)	$(10,54) \times (212,331)$	(10,19)	ACCENT	Murphy et al. [1998]
PAC_TRP	CH ₃ I	1383	(2001-02-26,2001-04-09)	$(1,56) \times (110,288)$	(0,12)	TRACE-P	Jacob et al. [2003]
VTPAC_TRP	CH ₃ I	3721	(2001 - 02 - 27, 2001 - 04 - 03)	$(-8,61) \times (62,210)$	(0, 12)	TRACE-P	Jacob et al. [2003]
NAGMEX_PRE_AVE	CH ₃ I	138	(2004-01-16, 2004-02-02)	$(-18,53) \times (213,329)$	(8,19)	Pre-AVE	N/A
INTEXA	CH ₃ I	2654	(2004-07-06, 2004-08-14)	$(12,69) \times (200,360)$	(0,13)	INTEX-A	Singh et al. [2006]
NA_CR_AVE	CH ₃ I	151	(2006-01-17, 2006-02-09)	$(-17,26) \times (219,333)$	(12, 19)	CR_AVE	N/A
X_INTEXB	CH ₃ I	893	(2006-02-24, 2006-03-22)	$(-1,50) \times (205,324)$	(0,12)	INTEX-B	Singh et al. [2009]
EPAC_INTEXB	CH ₃ I	1058	(2006-04-17, 2006-05-15)	$(4,65) \times (134,285)$	(0, 12)	INTEX-B	Singh et al. [2009]
ASKA_INTEXB	CH ₃ I	369	(2006-05-01,2006-05-15)	$(35,78) \times (140,275)$	(0, 12)	INTEX-B	Singh et al. [2009]
AY_ARCTAS	CH ₃ I	807	(2008-06-18, 2008-07-13)	$(17,57) \times (182,295)$	(0,10)	ARCTAS	Jacob et al. [2010]
ADA_ARCTAS	CH ₃ I	699	(2008-06-26, 2008-07-13)	$(30,75) \times (190,328)$	(0,12)	ARCTAS	Jacob et al. [2010]
ZON_A2B	Rn	61	(1987-04-20,1987-05-08)	$(-19,15) \times (249,360)$	(0,6)	ABLE-2B	Peretra et al. [1991]
SNEPAC_KRITZ	Rn	127	(1994-06-03,1994-08-16)	$(19,56) \times (184,296)$	(0,13)	N/A	Kritz et al. [1998]
VT_NH_LIU	Rn	150	(1970-07-01, 1970-08-31)	$(19,64) \times (0,321)$	(0, 12)	N/A	Ltu et al. [1984]

Radon Profiles & Relative VMR Change with Height

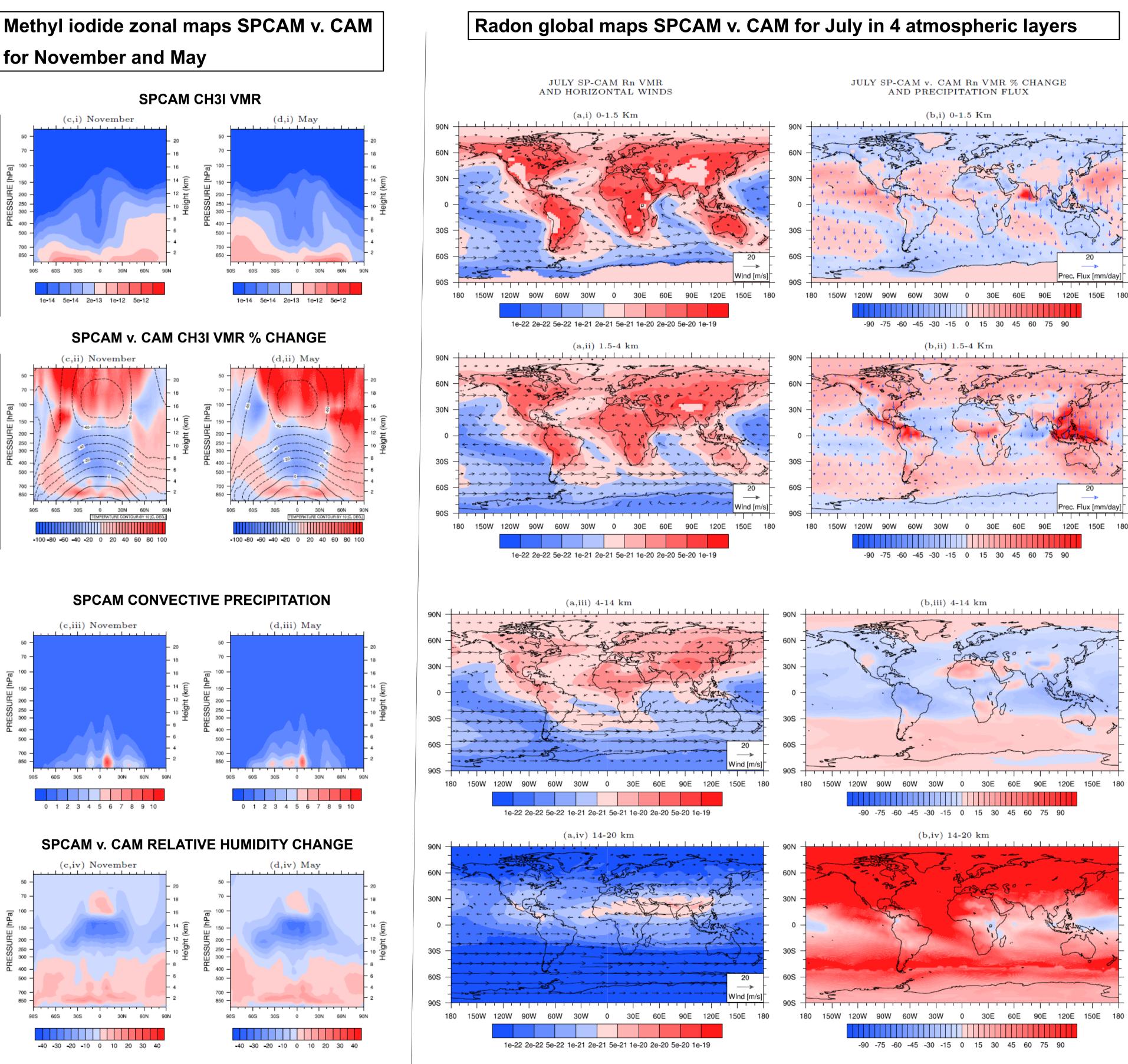
Methyl iodide Profiles & Relative VMR Change with Height

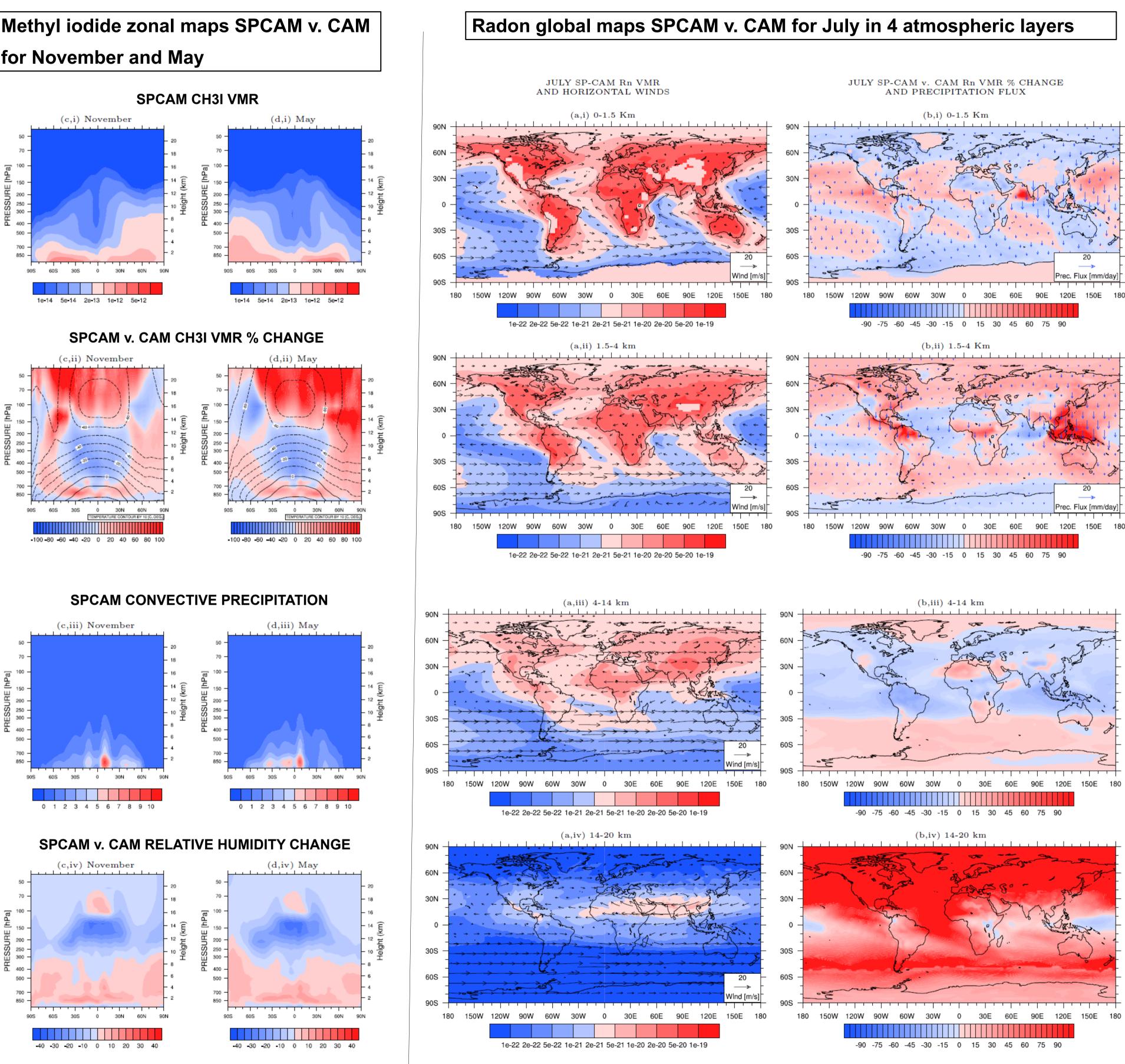
-100

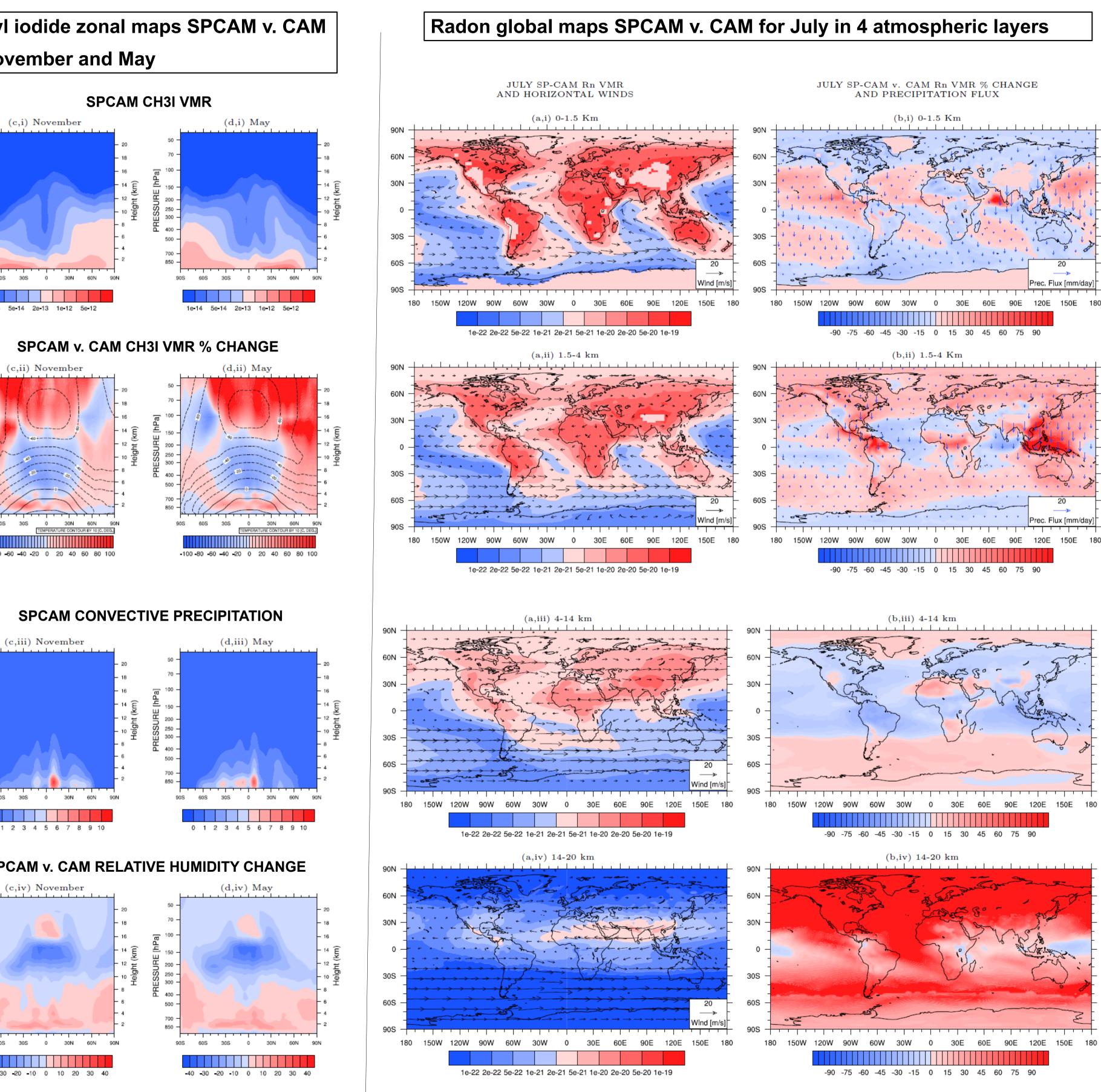
PBL

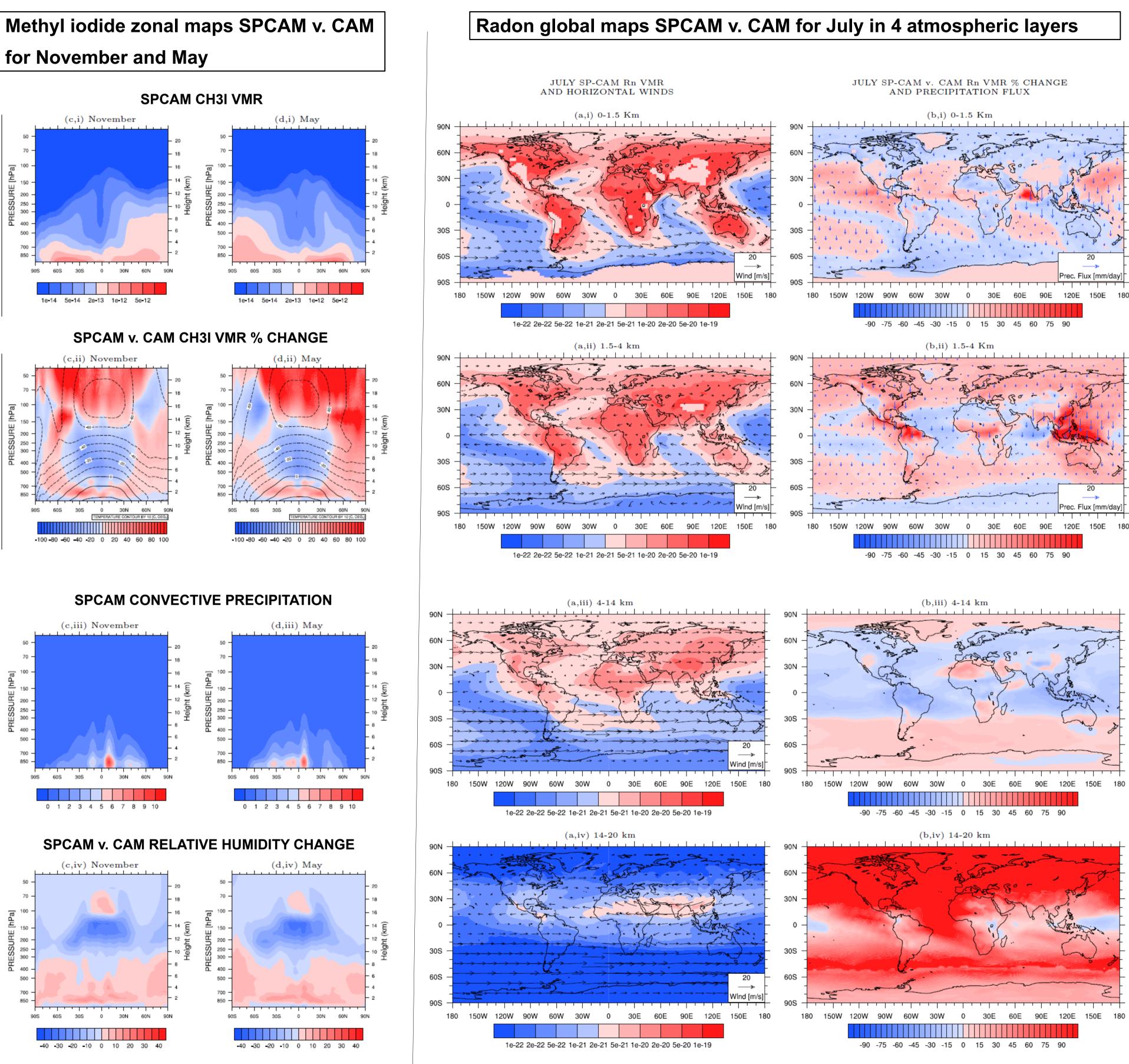
LTL MTL UTL

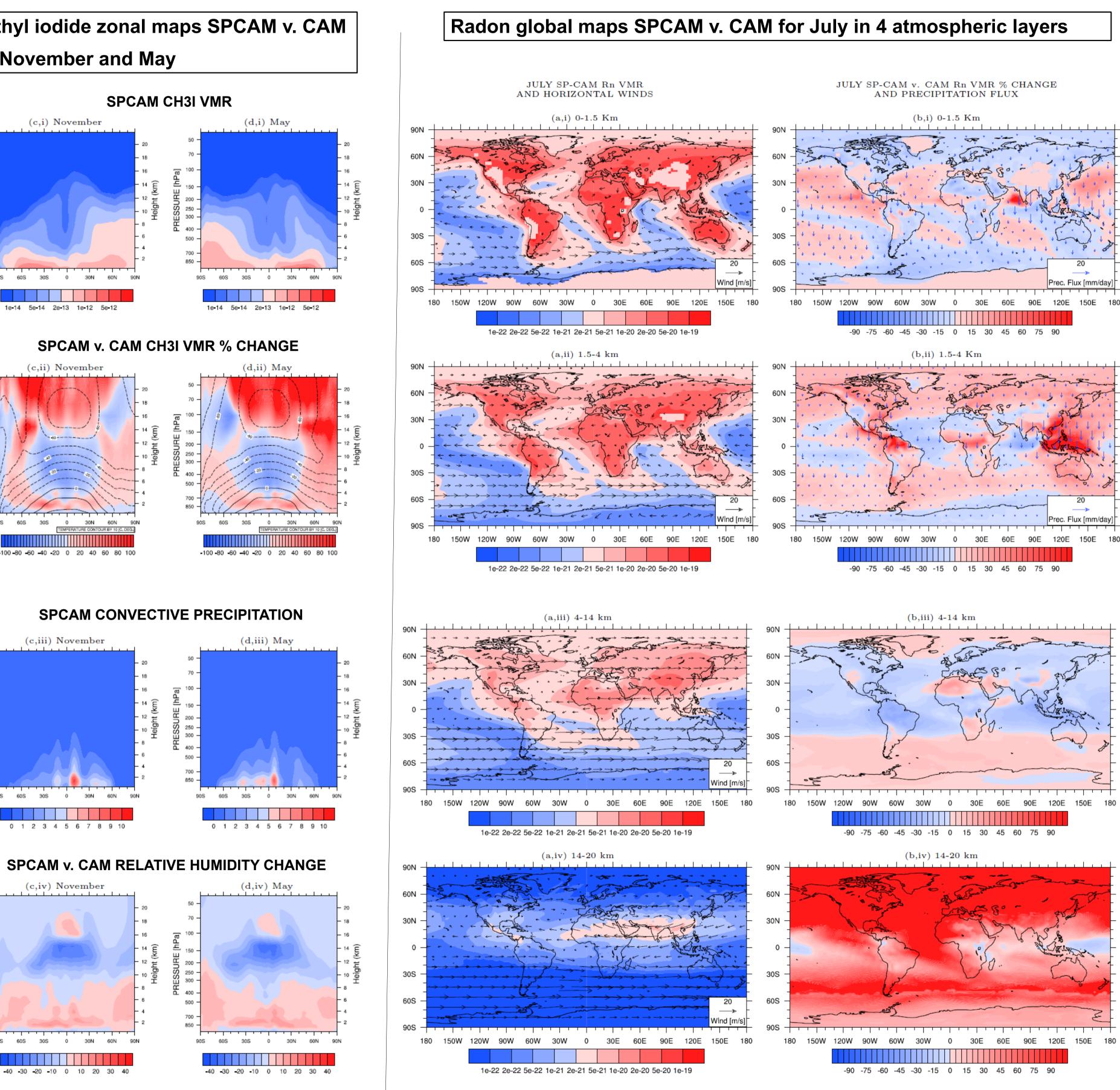
SUMMARY SPCAM v. CAM DIFFERENCES FROM SIMULATIONS ONLY:

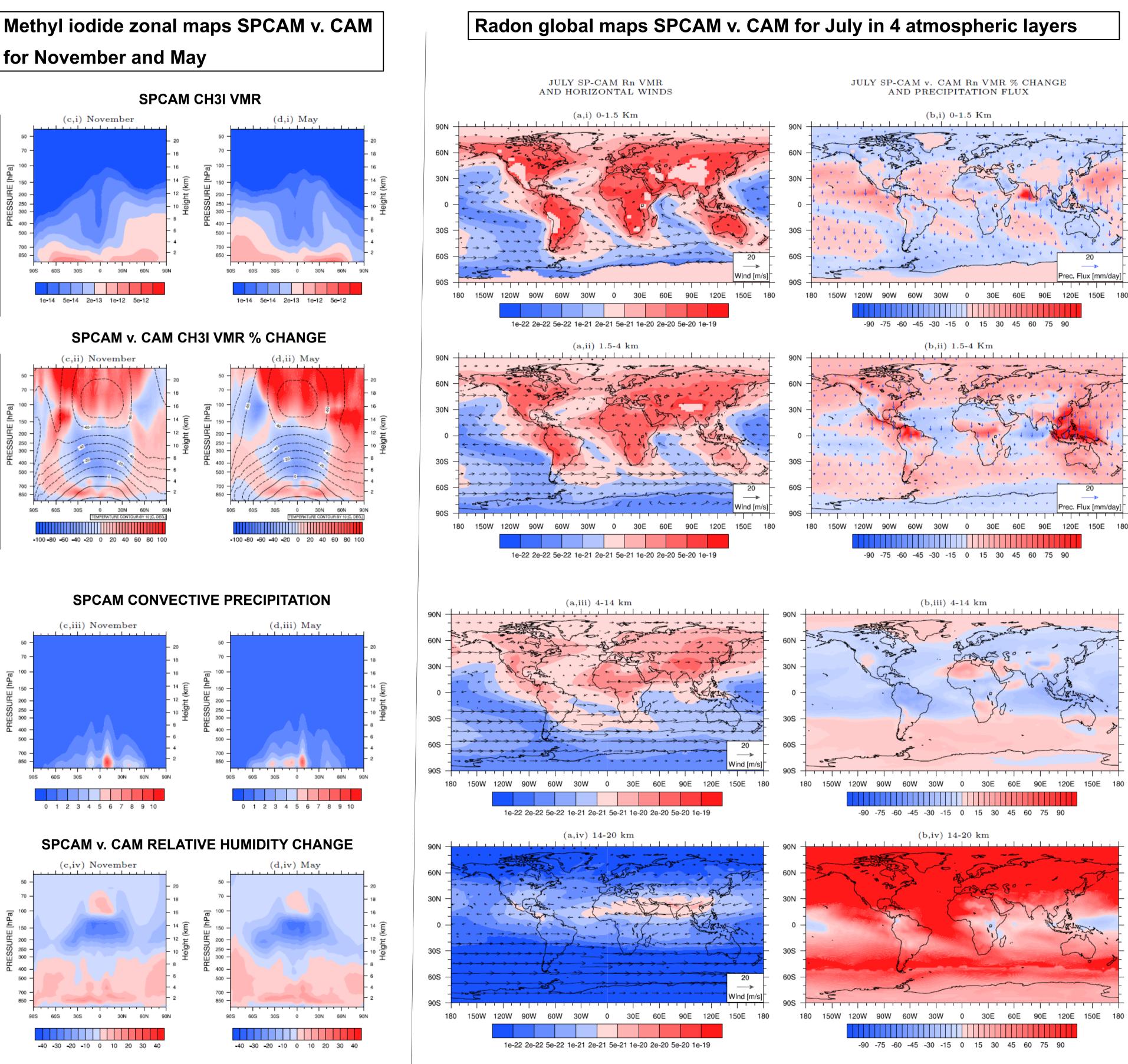












At latitudes where most of the convective precipitation occurs, with respect to CAM, SP-CAM depletes the PBL, enriches the layers of shallow convection, and depletes the layer of deep convection up to approximately 2 Km below the tropopause above which it enriches the atmosphere.