Verification of Various formulations of Coupling Using a 3D MMF *(ongoing work)*

- Construct a 3D MMF in which a 3D CRM is coupled with an idealized GCM.
- Compare the results of 3D MMF with those of the straightforward application of the 3D CRM.

Vector Vorticity Cloud Model (VVCM)

- A Three-Dimensional Anelastic Model Based on the Vorticity Equation Joon-Hee Jung and Akio Arakawa, 2008, *Mon. Wea. Rev.*, **136**, 276-294.
- Nonhydrostatic anelastic 3D model
- Prognostic variables:
	- Horizontal components of vorticity
	- Vertical component of vorticity (*at a certain height*)
	- *Horizontally uniform* part of horizontal velocity (*at a certain height*)
	- Potential temperature
	- Mixing ratios of various phases of water
- 3D elliptic (or parabolic) equation is solved for vertical velocity
- Physics:
	- Bulk ice-phase microphysical parameterization
	- Radiation parameterization
	- Turbulence parameterization (1st-order closure)

Benchmark Simulation with VVCM

- Domain size: 384 km x 384 km x 18 km
- Horizontal resolution: 3 km
- Vertical resolution: 34 layers with a stretched vertical grid
- Lower boundary: ocean surface with a fixed temperature
- Idealized tropical condition: based on a GATE Phase-III mean sounding and a wind profile during TOGA COARE
- Large-scale forcing: prescribed cooling and moistening tendencies
- Perturbation: random temperature perturbations into the lowest layer

Approach A: Explicit formulation of GCM/CRM effects Approach B: Mutual adjustments of prognostic variables Approach C: Hybrid of A and B

- Approach A includes an ad hoc way of eliminating the double counting.
- Horizontal resolution of GCM: 96 km

EXPERIMENTS 1 AND 2

For thermodynamic variables

Approach A

$$
\begin{cases} \frac{\partial q_C}{\partial t} = S_C + \hat{S}_G\\ \frac{\partial q_G}{\partial t} = S_G + \langle S_C \rangle \end{cases}
$$

• For vorticity components

 $\left\langle S_{C}\right\rangle ^{*}$ includes only the tendency due to turbulence and surface flux.

Relative Error of a 3D-MMF Simulation (EXP 1)

Relative Error of a 3D-MMF Simulation (EXP 1)

Sensible heat

-29 % -25 %

Errors are quantitative rather than qualitative.

Relative Error of a 3D-MMF Simulation (EXP 2 and 2a)

Mass-weighted vertical mean Mass-Weighted Vertical mean of thermodynamic variables of variances

Evaporation

Sensible heat

is the cause for this.

EXP2a is practically the same.

EXPERIMENT 3

For both thermodynamic variables and vorticity components

Approach B

$$
\begin{cases} \frac{\partial q_C}{\partial t} = S_C - \frac{1}{\tau_C} \left(q_C - \hat{q}_G \right) \\ \frac{\partial q_G}{\partial t} = S_G - \frac{1}{\tau_G} \left(q_G - \langle q_C \rangle \right) \end{cases}
$$

 τ_c : time scale for the response of convection to large-scale forcing τ_G : time scale for the adjustment of the large-scale fields by convection

Ambiguities in formulating forcing and feedback do not exist.

• Since q_c tends to be adjusted to \hat{q}_G , q_c may be excessively damped.

Relative Error of a 3D-MMF Simulation (EXP 3)

Variance of w

Convective activity is sensitive to the choice of the time scales.

Over-prediction when τ_{G} τ_{C} *Under-prediction when* $\tau_{G} \leq \tau_{C}$

Relative Error of a 3D-MMF Simulation (EXP 3)

Mass-weighted vertical mean Mass-Weighted Vertical mean of thermodynamic variables of variances

Evaporation Sensible heat -36 % -26 % -12 % -10 %

EXPERIMENT 4

Approach C

For thermodynamic variables

$$
\begin{cases}\n\frac{\partial q_C}{\partial t} = S_C + \hat{S}_G \\
\frac{\partial q_G}{\partial t} = \left\langle \hat{S}_G \right\rangle + \left\langle S_C \right\rangle - \frac{1}{\tau_T} \left(q_G - \left\langle q_C \right\rangle \right)\n\end{cases}
$$

• For vorticity components

$$
\begin{cases}\n\frac{\partial q_C}{\partial t} = S_C + \hat{S}_G \\
\frac{\partial q_G}{\partial t} = \left\langle \hat{S}_G \right\rangle + \left\langle S_C \right\rangle - \frac{1}{\tau_V} \left(q_G - \left\langle q_C \right\rangle \right)\n\end{cases}
$$

Relative Error of a 3D-MMF Simulation (EXP 4) $\tau_{\mathbf{v}} = \mathbf{I}$ hr

Mass-weighted vertical mean Mass-Weighted Vertical mean of thermodynamic variables of variances

Summary and Conclusion

- Errors of the MMF are sensitive to formulation of the coupling.
- The variances of horizontal winds are under-predicted in all experiments.
- Consequently, the surface heat fluxes are under-predicted. The low-level temperature and humidity tend to be low.
- Finding a proper way of the coupling should be one of the central problems in MMF development. This is especially true for horizontal momentum and cloud-microphysical variables.

Revised Approach B (to be tested)

Predictor Step:

$$
\begin{cases}\n\frac{q_G^* - q_G^n}{\Delta t} = S_G + \langle P_C \rangle \\
\frac{q_C^* - q_C^n}{\Delta t} = S_C + \hat{P}_G\n\end{cases}
$$

P represents sources/sinks *due to physics*.

$$
\begin{aligned}\n\text{Adjustment Step:} \\
\begin{cases}\n\frac{q_{\text{G}}^{n+1} - q_{\text{G}}^{*}}{\Delta t} = -\frac{1}{\tau} \left(q_{\text{G}} - \langle q_{\text{C}} \rangle \right) \\
\frac{q_{\text{C}}^{n+1} - q_{\text{C}}^{*}}{\Delta t} = -\frac{1}{\tau} \left(\langle q_{\text{C}} \rangle - \langle \hat{q}_{\text{G}} \rangle \right)\n\end{cases}\n\end{aligned}
$$

- Implementation of the source/sink terms and mutual adjustment are performed sequentially.
- No local damping of q_c .
- Sources/sinks due to physics are compatible between the GCM and CRM.