

Z-grid unified model update and an issue

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Update

- **April 2009**
 - **Z-grid anelastic global model**
 - Based on Z-grid hydrostatic model of Ross
- **July 2009**
 - **Unified model**
 - Cold bubble test in a limited domain (Cartesian)
 - Z-grid unified global model: Jablonowski's test (It was just a lucky.)
- **January 2010**
 - **Tentative solution to a problem**
 - Model was unstable in a case including larger divergence.
 - **Implementing SAM's physics**
 - (finished) Surface fluxes
 - (finished) Cloud microphysics
 - (not yet) Turbulence
 - (not yet) Radiation
 - **Test simulations with "clouds"**

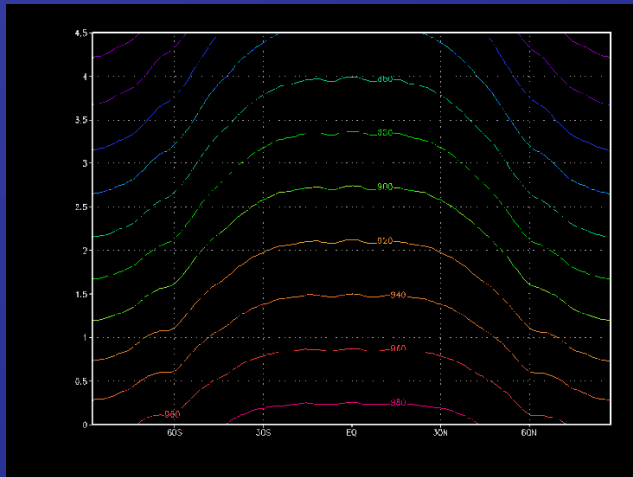
Surface fluxes + -2 K forcing

- SAM's surface flux parameterization
- -2 K forcing for the prognostic equation of potential temperature



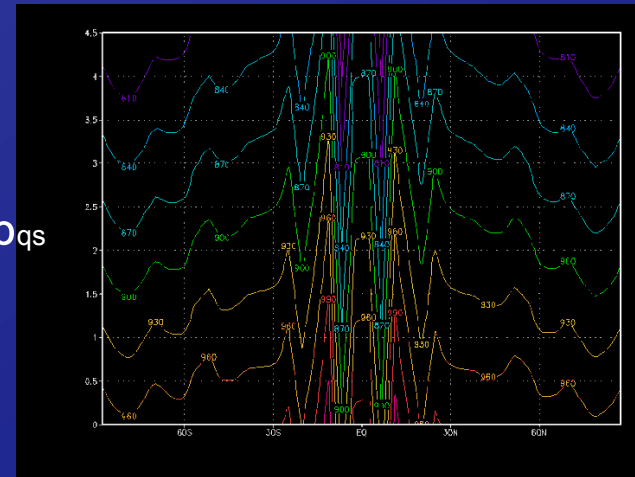
Model blows up...

p



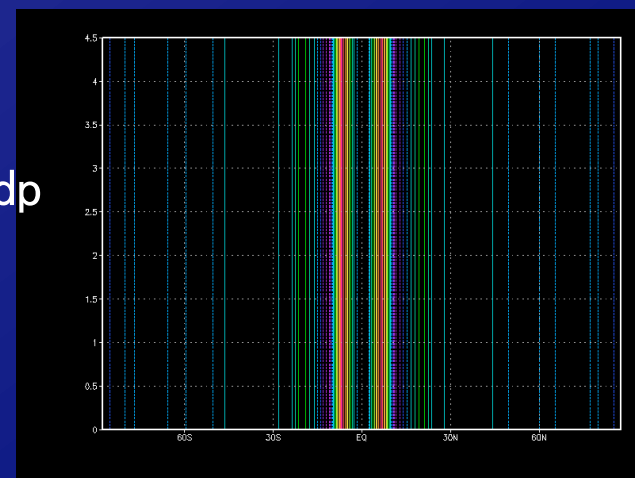
=

p_{qs}



+

dp



Cause

$$\frac{\partial(\pi_{qs})_{sf c}}{\partial t} = \frac{1}{(p_{qs}/\pi_{qs})_{sf c} - (p_{qs}/\pi_{qs})_{top}} \left[\left(\frac{p_{qs}}{\pi_{qs}} \right)_{top} \int_{z_{sf c}}^{z_{top}} \frac{g}{c_p \theta^2} \frac{\partial \theta}{\partial t} dz - \kappa g \int_{z_{sf c}}^{z_{top}} \nabla_h \cdot (\rho_{qs} \mathbf{v}_h) dz \right] \quad (2.16)$$

$(\pi_{qs})_{sf c} \theta \rightarrow \pi_{qs} \rightarrow p_{qs} \rightarrow \rho_{qs}$

Forward scheme is used to get quasi-static density of n+1.

$$\frac{\rho_{qs}^{n+1} - \rho_{qs}^n}{\Delta t} = -\nabla_h \cdot (\rho_{qs} \mathbf{v}_h)^{n+1} - \frac{\partial}{\partial z} (\rho_{qs} w)^{n+1} \quad (4.5)$$

Backward scheme is assumed to get a Poisson equation.

Assume $\text{del}(\theta)/\text{del}(t) = 0$.

- If horizontal momentum is convergent for a vertical column at time “n”, vertically integrated horizontal momentum at time “n+1” must be convergent to satisfy (4.5).

$$\int_{z_{sf c}}^{z_{top}} \nabla_h \cdot (\rho_{qs} \mathbf{v}_h) dz > 0 \rightarrow \int_{z_{sf c}}^{z_{top}} \nabla_h \cdot (\rho_{qs} \mathbf{v}_h)^{n+1} dz > 0$$

- Then, convergent momentum at “n+1” is used to get mass at “n+2”. That means vertically integrated momentum at “n+2” is also convergent.
 - ρ_{qs} tries to make divergence, but dp does not allow that.
- There is a positive feedback in the vertically integrated horizontal momentum.

Solution

$$\frac{\partial(\pi_{qs})_{sf c}}{\partial t} = \frac{1}{(\rho_{qs}/\pi_{qs})_{sf c} - (\rho_{qs}/\pi_{qs})_{top}} \left[\left(\frac{p_{qs}}{\pi_{qs}} \right)_{top} \int_{z_{sf c}}^{z_{top}} \frac{g}{c_p \theta^2} \frac{\partial \theta}{\partial t} dz - \kappa g \int_{z_{sf c}}^{z_{top}} \nabla_h \cdot (\rho_{qs} \mathbf{v}_h)^{n+1} dz \right] \quad (2.16)$$

Backward scheme for (2.16)?

$$\frac{(\rho_{qs} \mathbf{v}_h)^{n+1} - (\rho_{qs} \mathbf{v}_h)^n}{\Delta t} = -c_p \rho_{qs} \theta (\nabla_h \pi_{qs}^n + \nabla_h \delta \pi^n) + \mathbf{A}_{\mathbf{v}_h}^n$$

δp is diagnosed by a Poisson equation assuming (4.5).

$$\frac{\rho_{qs}^{n+1} - \rho_{qs}^n}{\Delta t} = -\nabla_h \cdot (\rho_{qs} \mathbf{v}_h)^{n+1} - \frac{\partial}{\partial z} (\rho_{qs} w)^{n+1} \quad (4.5)$$

$\rho_{qs}(n+1)$ is computed through (2.16).

We need to make implicit equations to avoid the positive feedback of divergence/convergence.

I tried. But, I have not succeeded yet.


I could not clear the complexity coming from $p_{qs} = p_{00} \pi^{1/\kappa}$.

Temporary solution

Is it possible to estimate dp or to avoid the use of dp ?

1. Updating momentum without dp .
2. Estimating dp through iteration (Runge-Kutta).
3. Using dp of the previous time step.
4. Combination of 1, 2 and 3.
5. Prognosing vertical mean part of dp .
6. Horizontally explicit and vertically implicit computation of momentum equations.
7. Damping the Lamb wave part in 4.

$$\frac{\partial(\pi_{qs})_{sf c}}{\partial t} = \frac{1}{(p_{qs}/\pi_{qs})_{sf c} - (p_{qs}/\pi_{qs})_{top}} \left[\left(\frac{p_{qs}}{\pi_{qs}} \right)_{top} \int_{z_{sf c}}^{z_{top}} \frac{g}{c_p \theta^2} \frac{\partial \theta}{\partial t} dz - \kappa g \int_{z_{sf c}}^{z_{top}} \nabla_h \cdot (\rho_{qs} \mathbf{v}_h) dz \right] \quad (2.16)$$


$$\frac{\partial(\pi_{qs})_{sf c}}{\partial t} = 0$$

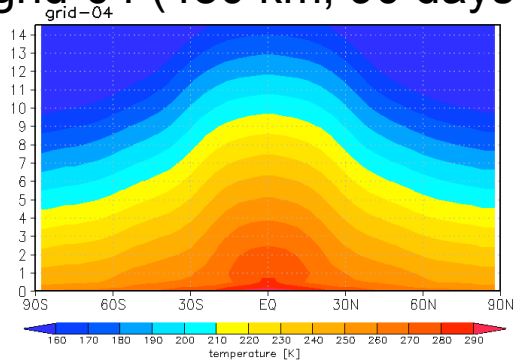
Experimental setup

- Z-grid unified dynamical core
 - Fixed $(p_{qs})_{sfc}$, 2nd-order Runge-Kutta
- SAM's physics
 - w/ surface fluxes, cloud microphysics
 - w/o turbulence, radiation, ice sedimentation
- Horizontal grids and time integrations
 - grid-04 (dx~480 km): 120 days (dt=360 s)
 - grid-05 (dx~240 km): 120 days (dt=150 s)
 - grid-06 (dx~120 km): 59 days (dt=60 s)
 - grid-09 (dx~ 15 km): 5 days (dt=10 s, 24 hr by 2560 PEs)
- Vertical grid
 - dz=500 m, 30 levels (0-15000 m), No sponge layer
- Aquaplanet
 - Control SST (http://www.atmos.ucla.edu/~brianpm/cfmip2_aqua.html)
 - No initial wind
 - temperature lapse rate of -6 K/km (RH=80%)
 - -2 K/day forcing

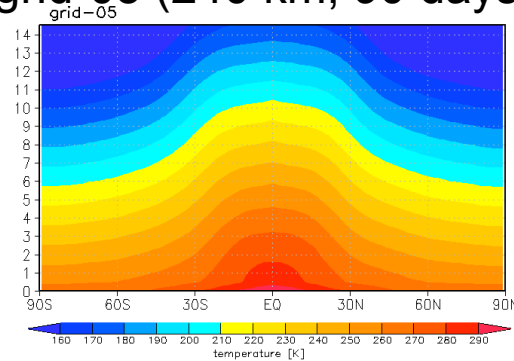
Zonal mean

Temperature

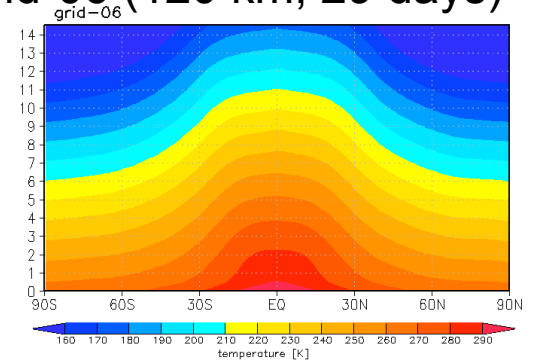
grid-04 (480 km, 90 days)



grid-05 (240 km, 90 days)

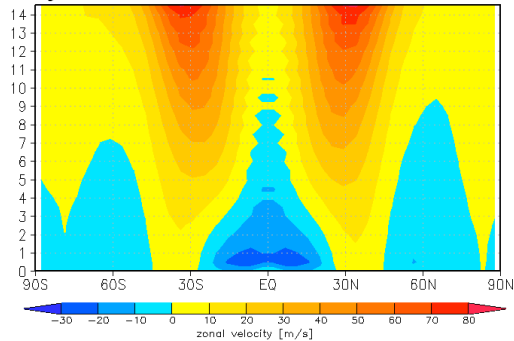


grid-06 (120 km, 29 days)

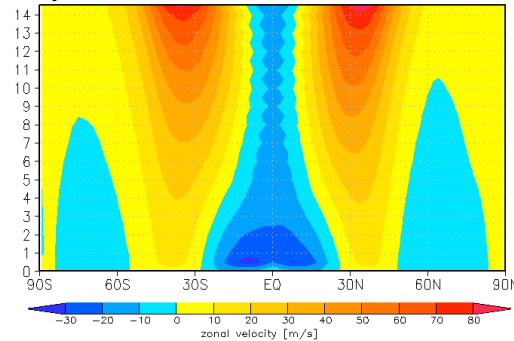


zonal velocity

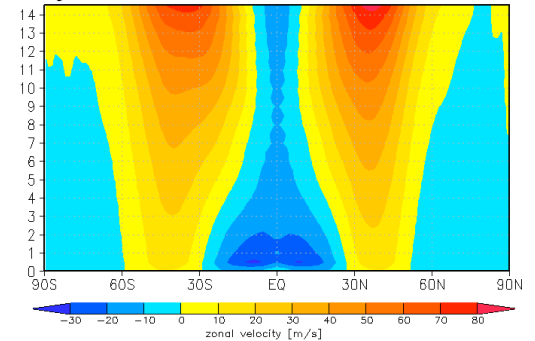
grid-04



grid-05

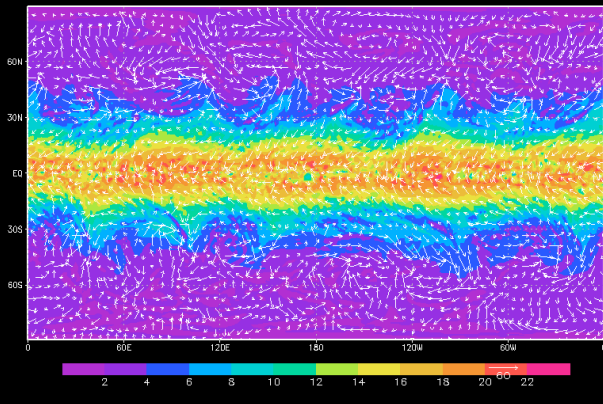


grid-06

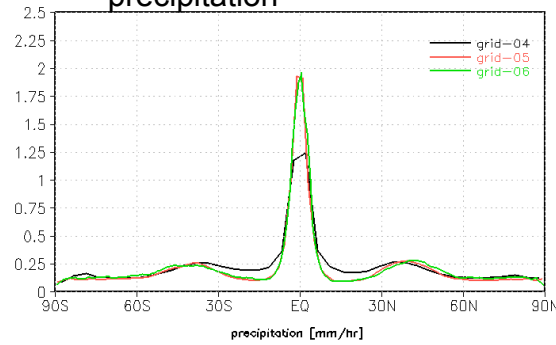


Moisture

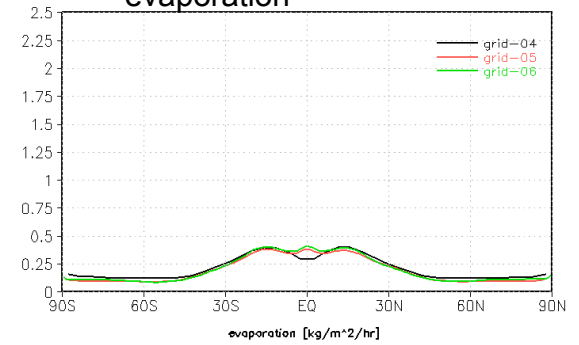
qv [g/kg] & u,v, grid-06, day-59



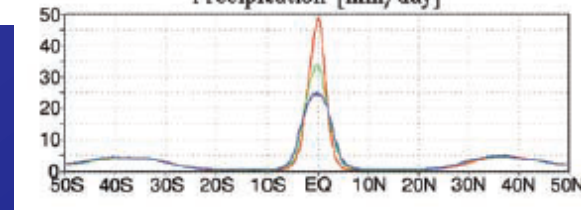
Time- and zonal-mean precipitation



Time- and zonal-mean evaporation



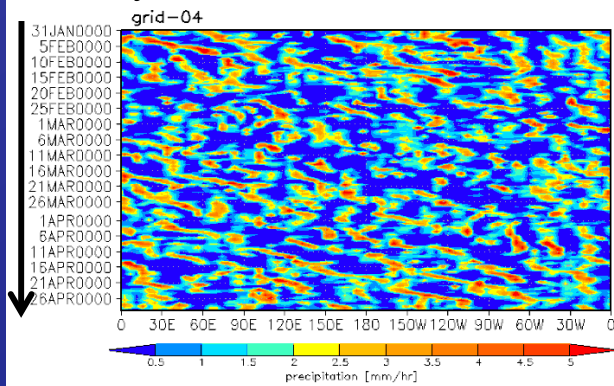
Precipitation [mm/day]



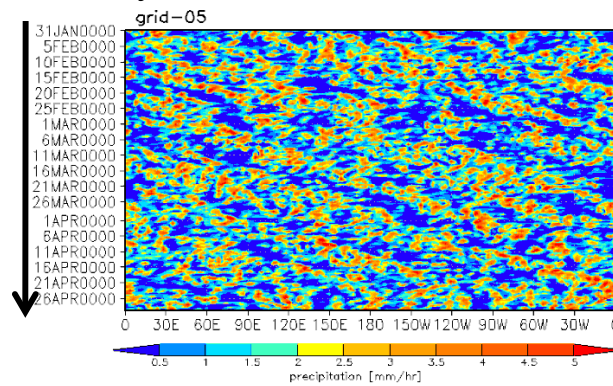
Tomita et al. (2005)

Hovmöller diagram (precipitation, 5S-5N mean)

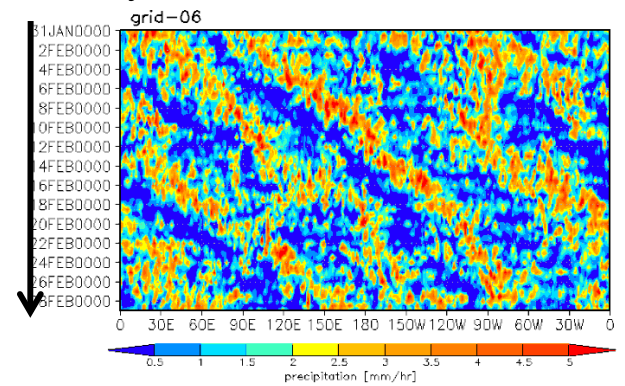
90 days



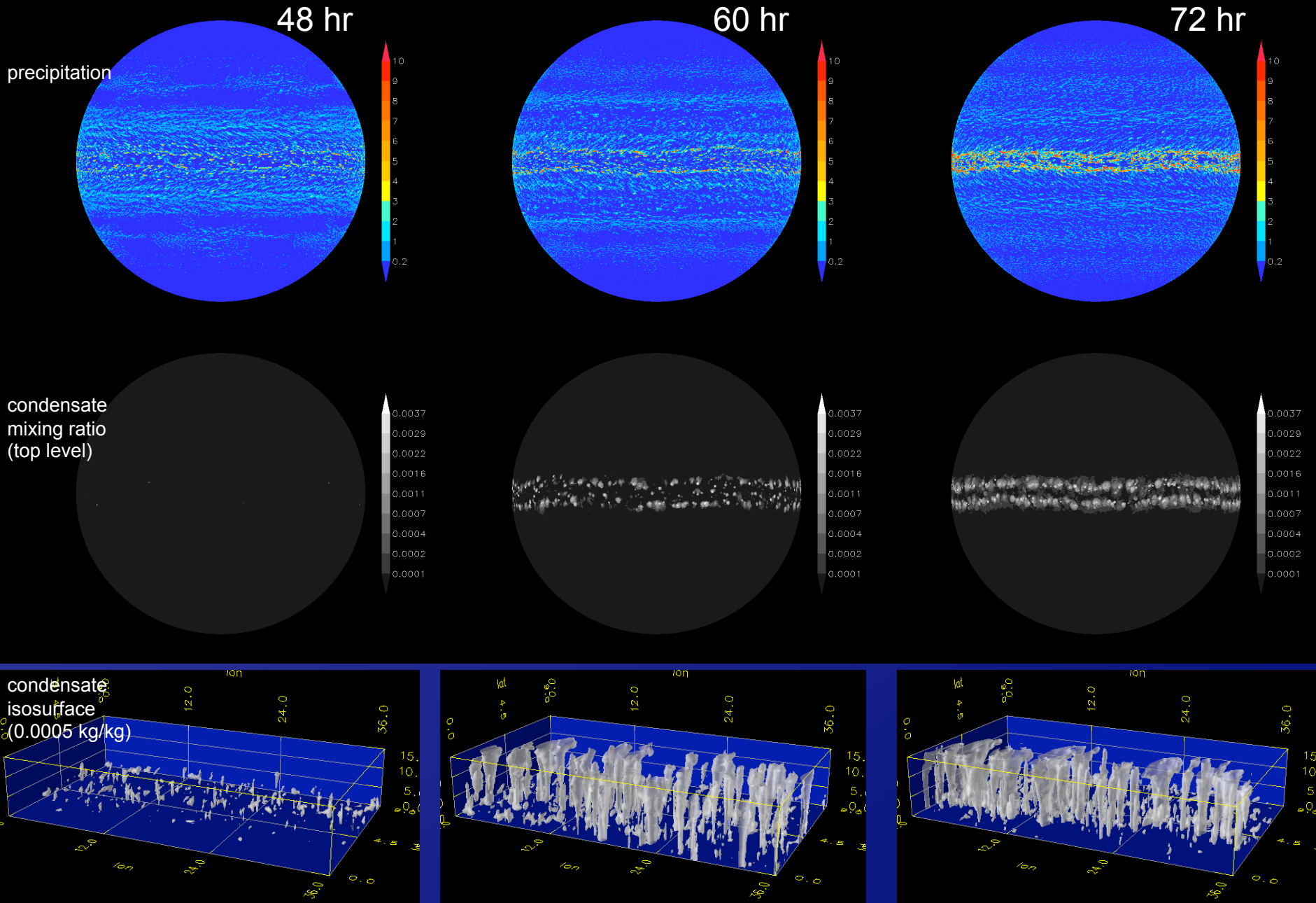
90 days



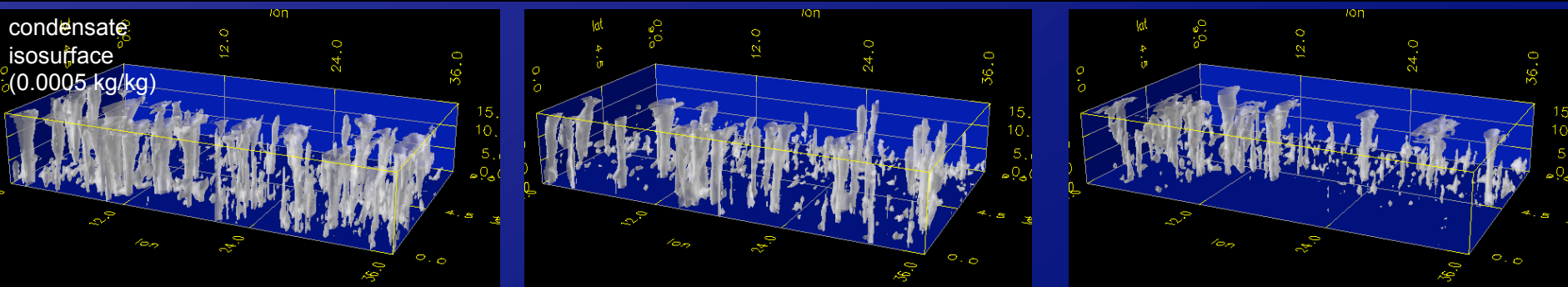
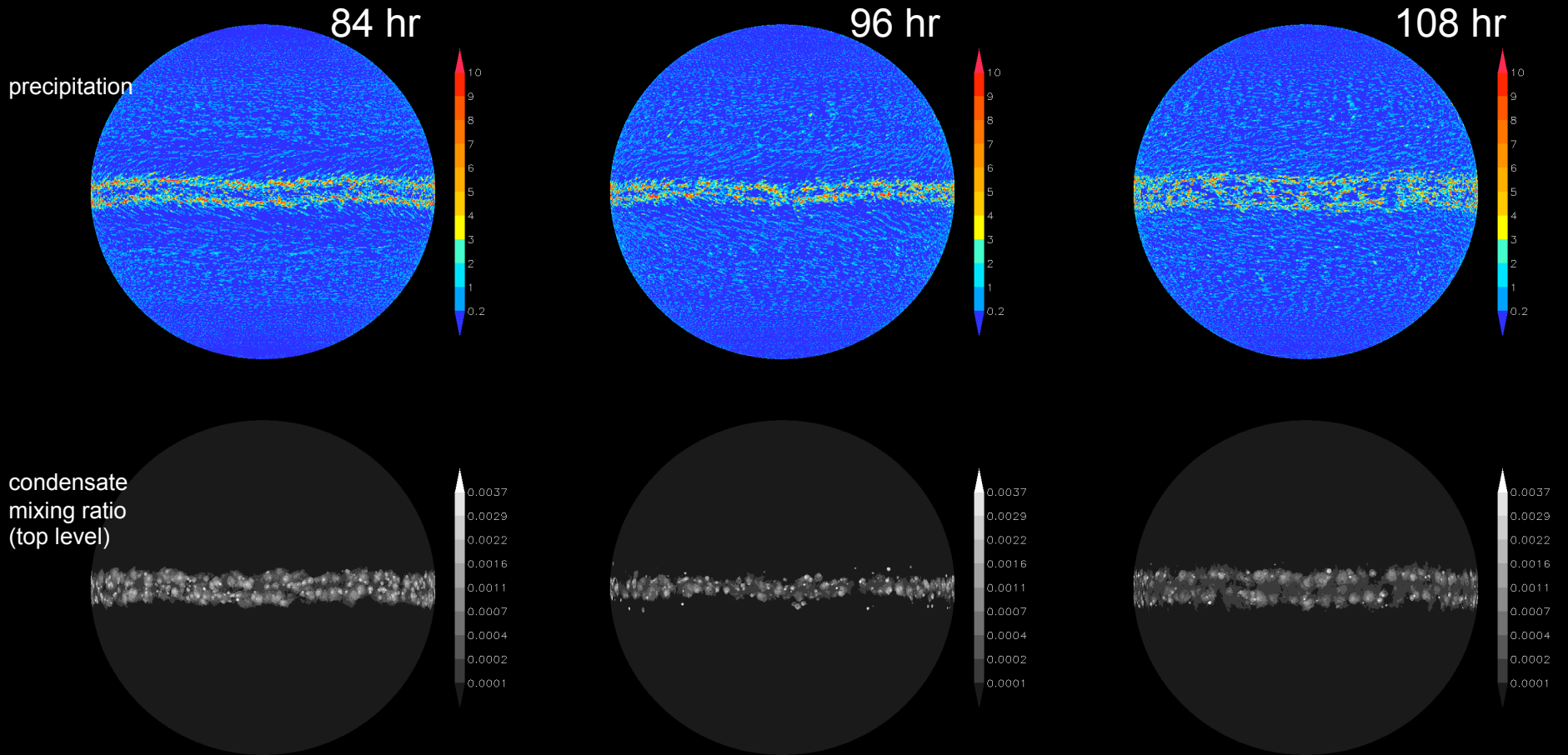
29 days



grid-09 (15 km)

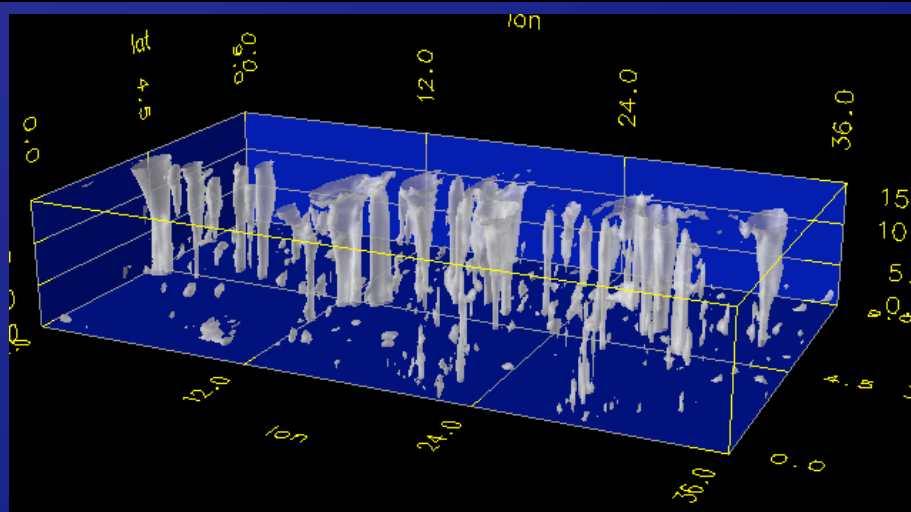
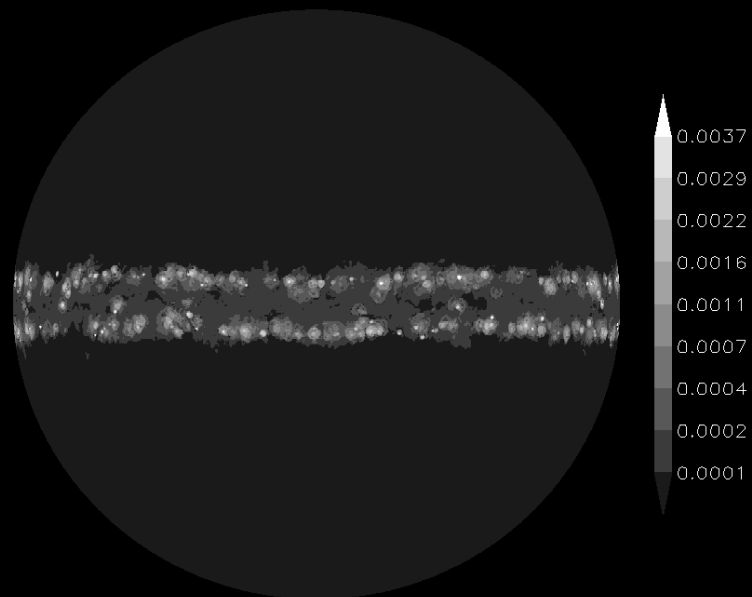
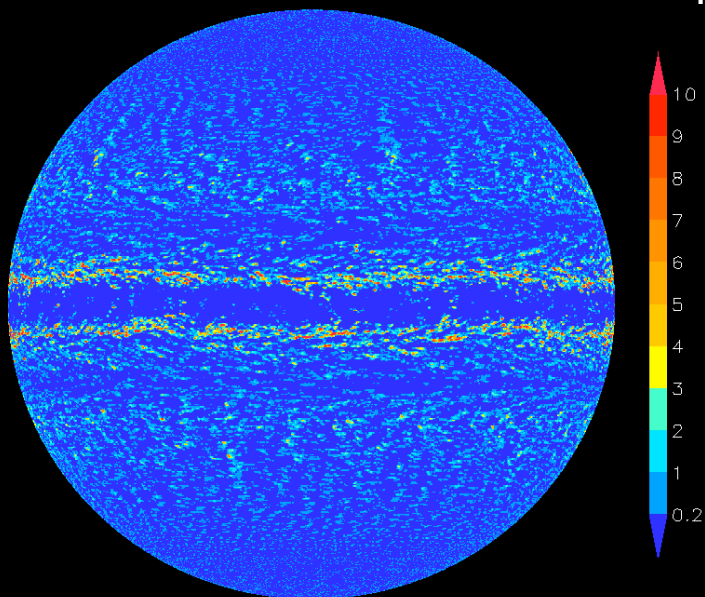


grid-09 (15 km)



grid-09 (15 km)

120 hr



Summary

- Surface fluxes and a cloud microphysics scheme have been transplanted from SAM.
 - I would like to thank Marat.
 - I had no time to test turbulence and radiation...
- Test simulations
 - Low-resolution simulations appear to reproduce “reasonable” atmosphere.
 - It is of course physically unreasonable to use cloud microphysics scheme.
 - 15-km simulation run for 5 days in its first attempt.
- There is an issue relating the prediction of the exner function of the surface.
 - We might need a magic!
 - I need to get a fresh start.

