# Chemical Transport in the MMF: Tests and Implications for Climate

Daniele Rosa, Wei-Chun Hsieh, and Bill Collins Earth and Planetary Science, UC Berkeley

Objectives
Tracers of interest
Experimental strategy
Data
Results of tests





# Objectives

Objectively test fidelity of vertical convective-scale transport comparing modeled vs. measured tracers

Quantify effects of vertical transport on the lifetime of dust and other radiative species





#### Residence time of aerosols increases with height



Residence time increases by factor of 4 from lower to upper troposphere.





### Tracers and species of interest

Chemical tracer species:

- Radon (Rn)
- Methyl Iodide (CH<sub>3</sub>I)
- Carbon Monoxide (CO)





### Tracers and species of interest

#### Chemical tracer species:

- Radon (Rn) Convection over lan
- Methyl Iodide (CH<sub>3</sub>I) -
- Carbon Monoxide (CO) Convection over fires

Convection over land Convection over oceans & rice





### Sources of Radon

Berkele



### Sources of CH<sub>3</sub>I and CO



### Tracers and species of interest

#### Chemical tracer species:

- Radon (Rn) -Convection over land
- Methyl Iodide (CH<sub>3</sub>I) Convection over oceans
- Carbon Monoxide (CO) Convection over fires

#### Soil dust:

- Dust is a significant natural radiatively active aerosol.
- Longwave dust forcing increases with altitude.





### Sources of dust



- Primary emissions are localized to northern subtropics and midlatitudes.
- Primary sinks are dry deposition and scavenging by precipitation.





### Characteristic lifetimes of tracers

#### CH<sub>3</sub>I: τ~2 days

- Useful for studying transport on short length/time scales.
- Rn: τ~5.5 days
  - Useful for studying transport on longer scales.
- CO: τ~40 days
  - UT gradient is balanced between convective divergence and photochemical processes.



Figure 11. An estimate of the tropical mean production of carbon monoxide, and its lifetime against photochemical loss. These estimates were used in the CONRAD (Emanuel) and TCM (Folkins and Martin) models to calculate the CO profiles shown in Figure 9. Their derivation is described in the text.

Folkins et al, 2006



#### Earlier studies of tracers for convective transport



Tracers help diagnose differences in mass fluxes and detrainment in alternative convective parameterizations.

Observational tracer comparisons help assess fidelity.





# **Experimental Strategy**

Control GCM: CAM
Experimental GCM: SPCAM

To isolate effects of cloud-system-scale velocities, GCMs are run as Chemical Transport Models.

> Large-scale lateral transports by  $\overline{v}$  are identical.

Small-scale vertical transports differ due to physics.





### Identical large-scale wind fields



# CTM mode simplifies model↔model tests



- Large scale meteorology is identical in two models.
- This eliminates feedbacks from physics to large scales, isolating signal from just convective-scale vertical motions.





### Sources of atmospheric moisture are identical



Moisture is controlled via surface water fluxes.





#### $\Delta RH$ due to convective-system scale physics.





MMF dries the PBL and UT and moistens the MT.



### CTM mode simplifies model⇔data tests

If reanalysis fields are used for the CTM mode in CAM and SPCAM, this minimizes errors in modeled vs. (one realization of) actual large-scale transport.

Differences between model and data should be dominated by model physics.





#### Impacts of convective transport on tracers







# Impact of convective transport on errors relative to observations





# Significant sub-grid dust transport?











### Satellite retrievals for carbon monoxide



# Initial Simulations for Dust



SPCAM transports less dust to free troposphere, and to high latitudes far from desert sources.





### Model differences in Carbon Monoxide





### Model differences in Methyl Iodide







### Model differences in Radon







### Radon sections: Lower troposphere



### Radon sections: Mid Troposphere



### Radon sections: Upper troposphere





#### SPCAM has smaller errors in MBL and lower troposphere.





### **Evaluation of Radon Profiles**



#### SPCAM has small errors in upper troposphere.





### Gradients as measures of transport



The gradient in fractional changes in concentration measures gain/loss in each layer.





### Metrics for fidelity of tracer gradients

**Table 5.** Marine Convection Index (MCI) Over the Pacific: Ratio of Upper Tropospheric (UT; 8-12 km) to Lower Tropospheric (LT; 0-2.5 km) CH<sub>3</sub>I Concentrations

Region	Campaign	Lifetime, <sup>a</sup> days	MCI (Observed)	MCI (Simulated)
5. North Pacific	PEM-WA	4.8	0.22	0.41
6. Hawaii	PEM-TB	4.4	0.26	0.18
	PEM-TA	4.7	0.20	0.39
8. Philippine Sea	PEM-WB	4.4	0.27	0.31
10. Christmas Island	PEM-TB	4.0	0.32	0.30
	PEM-TA	4.1	0.24	0.37
12. Fiji	PEM-TB	4.3	0.40	0.22
-	PEM-TA	4.7	0.16	0.14
13. Tahiti	PEM-TB	4.2	0.34	0.44
	PEM-TA	4.1	0.23	0.19
21. Easter Island	PEM-TB	4.4	0.11	0.11
<sup>a</sup> Moon model lifetime of CILL in the 0 to 12 km column (24 hour				

<sup>a</sup> Mean model lifetime of  $CH_3I$  in the 0- to 12-km column (24-hour average).

Bell et al, 2002





### **Convective indices**



• Cumulative convective index error =  $RMS[\delta dln(VMR)/dH]$ .

Ratio plotted here is CCIE(SPCAM) / CCIE(CAM).





# Conclusions

- Analysis of MMF in CTM mode reveals systematic differences in vertical transport due to convective-scale motions.
- These differences appear in all passive tracers simulated to date.
- Changes in model fidelity can be quantified using a convective index, a fractional measure of vertical gradients.
- Errors in the indices relative to observations systematically decreases in SPCAM.
- Next steps: Studies of dust and water vapor.



