

# Modeling the Effects of Variable Soil Moisture

**Piers Sellers**<sup>2</sup>, Ian Baker<sup>1</sup>, Scott Denning<sup>1</sup>, David Randall<sup>1</sup>,  
Isaac Medina<sup>1</sup>, Parker Kraus<sup>1</sup>

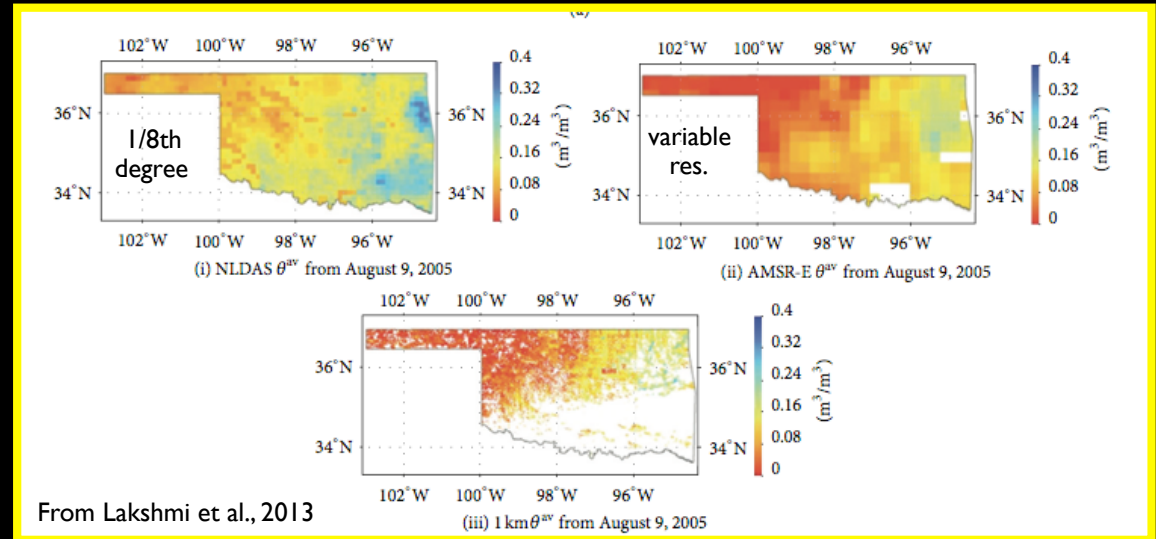
1: Atmospheric Science Department, Colorado State University, Fort Collins CO, USA

2: NASA GSFC, Greenbelt MD, USA



# Surface Heterogeneity

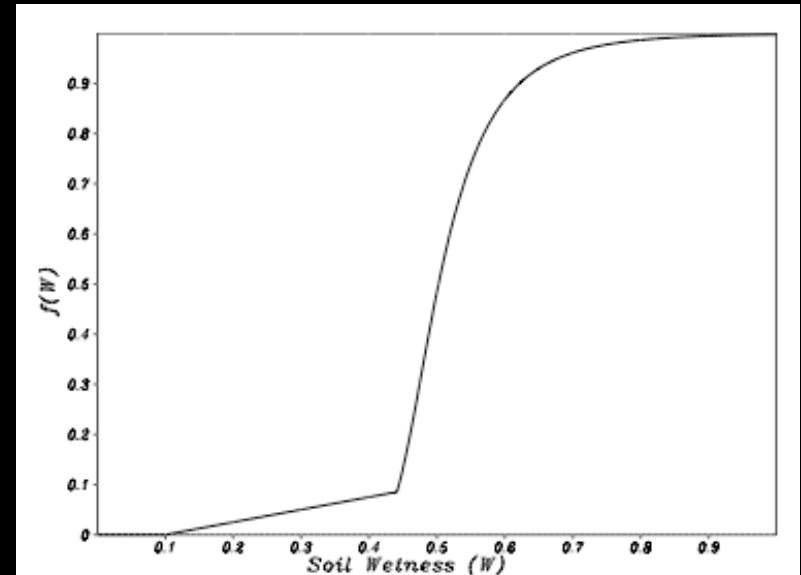
- Surface varies on scales smaller than gridcell
- Slope/aspect
- Vegetation
- Soil Moisture



# Soilwater Constraint on ET

$$E = E_p f(W)$$

- Roots and stomates link soil water to ET
- 1 = no stress; 0 = no transpiration
- Nonlinear; position depends on sand/clay %
- Scale Gap
- Problems when this  $f(W)$  is applied on GCM gridcells (Amazon)

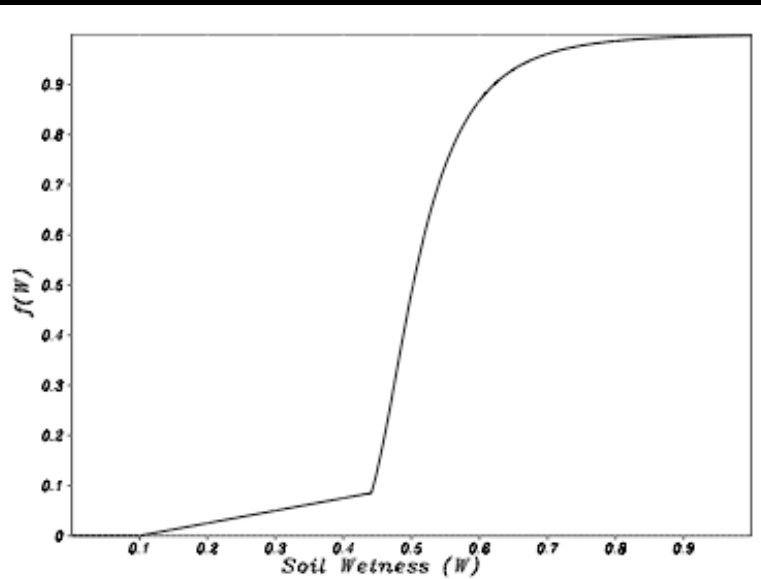


**Figure 1.** Relationship between the soil moisture stress function  $f(W)$  and soil moisture  $W$  as used in this study (see equation (2)). This function is based on data presented by *Colello et al. [1998]* for the FIFE prairie grassland site in Kansas, USA.

From Sellers et al., 2007

# Non-Linear Relationship

$$E = E_p f(W)$$



**Figure 1.** Relationship between the soil moisture stress function  $f(W)$  and soil moisture  $W$  as used in this study (see equation (2)). This function is based on data presented by *Colello et al.* [1998] for the FIFE prairie grassland site in Kansas, USA.



$$F(\bar{W}) \neq \bar{F}(W)$$

From Sellers et al., 2007

# Soil Wetness 'BINS'

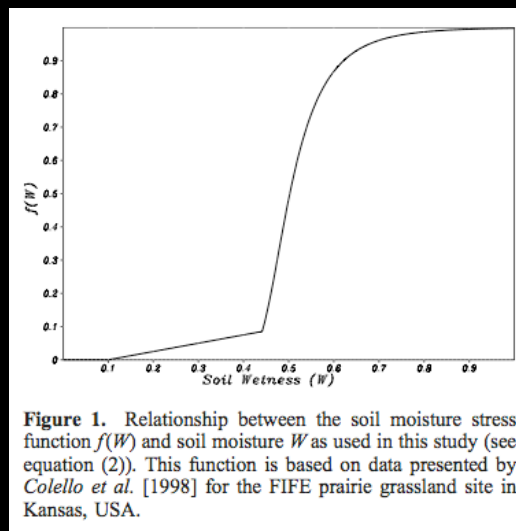
- Experiment: 3 methods to calculate ET in an idealized setting

1. Explicit,  $10^6$  gridcells

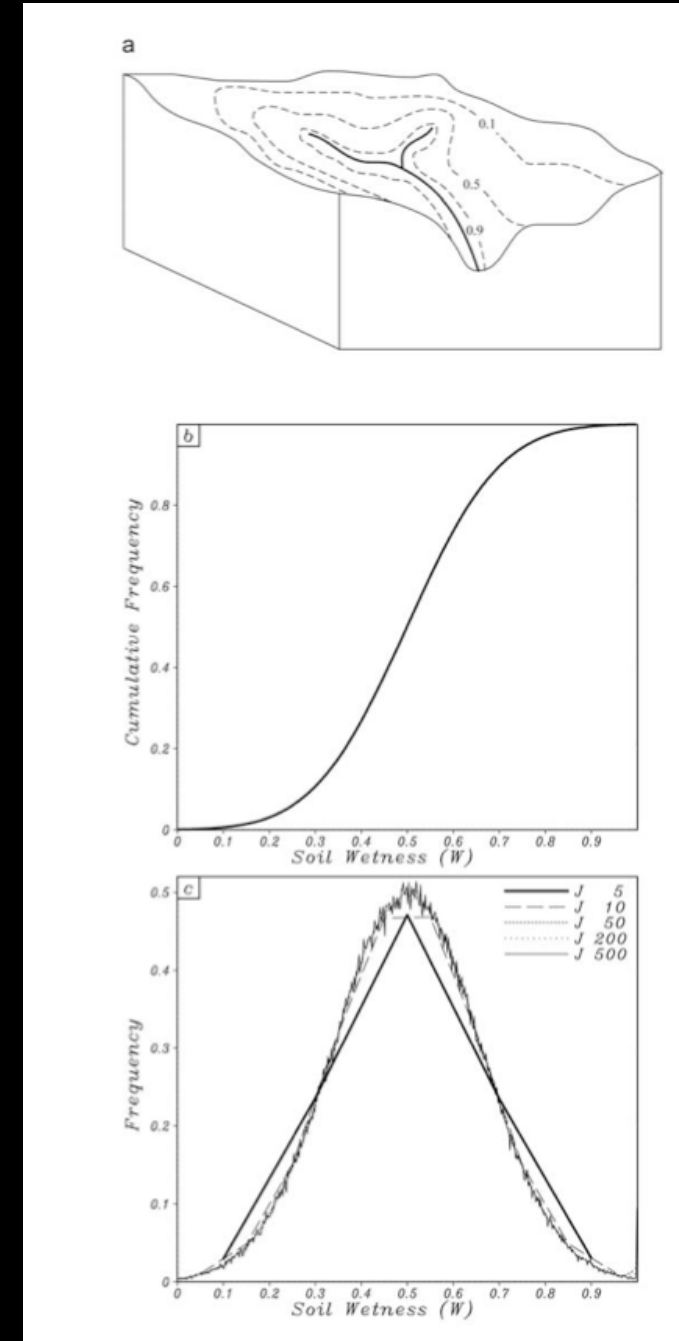
2. Single calculation of  $f(\overline{W})$

3. BINS,  $J=5$  to 500

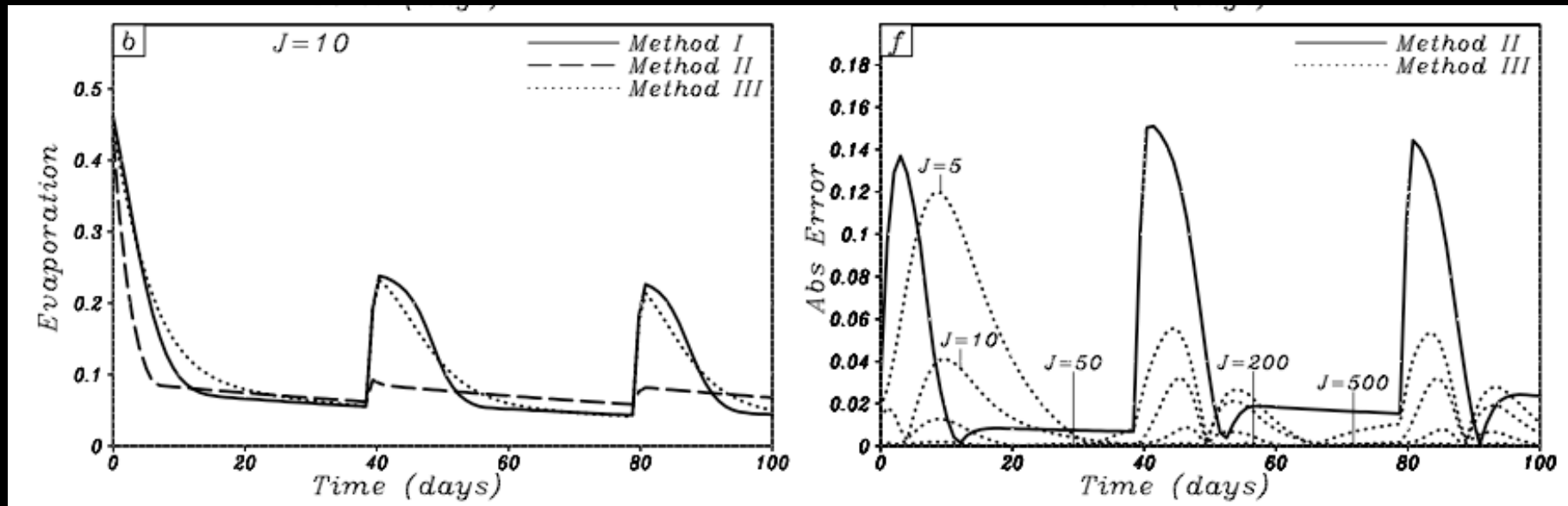
$$E = E_p \sum_{i=1}^n a_i f(W_i)$$



From Sellers *et al.*, 2007



# How Well Does the Conceptual Model Work?

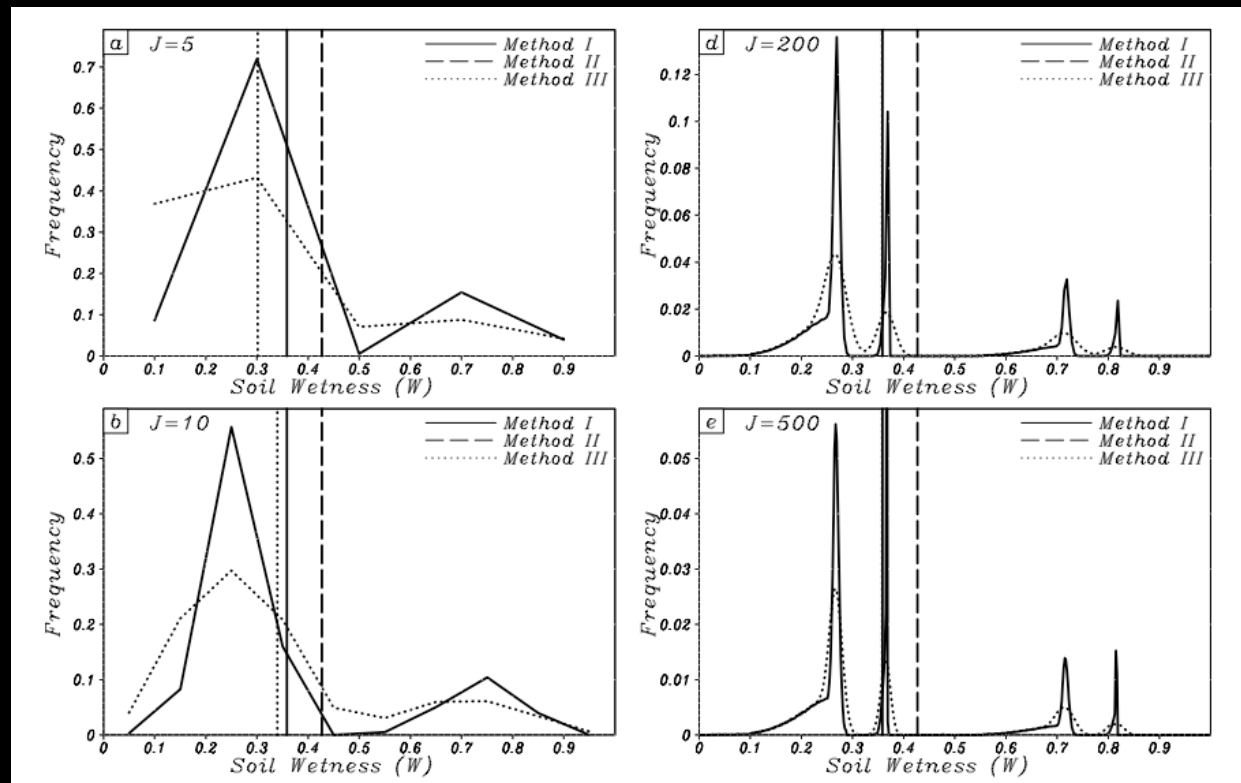


From Sellers et al., 2007

- 500 bins is indistinguishable from explicit method
- 10 bins is a reasonable approximation
- Even 5 bins is better than **single  $\bar{W}$**



# BIN distribution over time



From Sellers et al., 2007

- Wetting events (precipitation) manifest as positive area on the right (wet) of the distribution
- Drying moves area to the left, decreases variability (narrower distribution)

# Can We Implement BINS in a 'Real' Model?

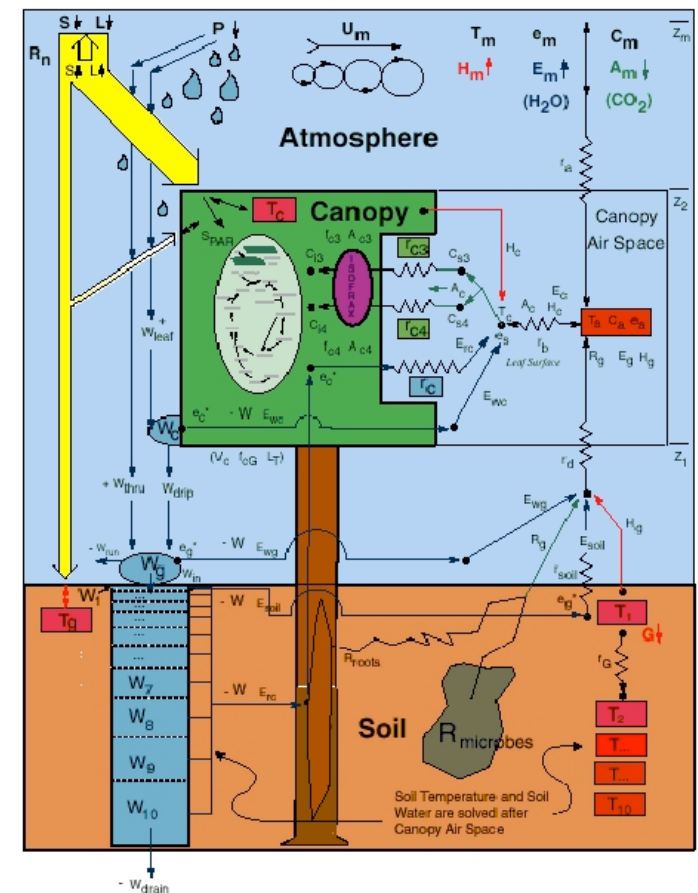
## Toy Model

- Single Soil layer
- Single control -  $f(W)$  - on ET
- no diurnal cycle

## Real Model

- Multiple Soil layers - with roots
- ET responds to  $W$ , temp., RH,  $R_{net}$ , phenology
- diurnal, synoptic, seasonal cycles, IAV (10-minute timestep)
- tested using 10 BINS

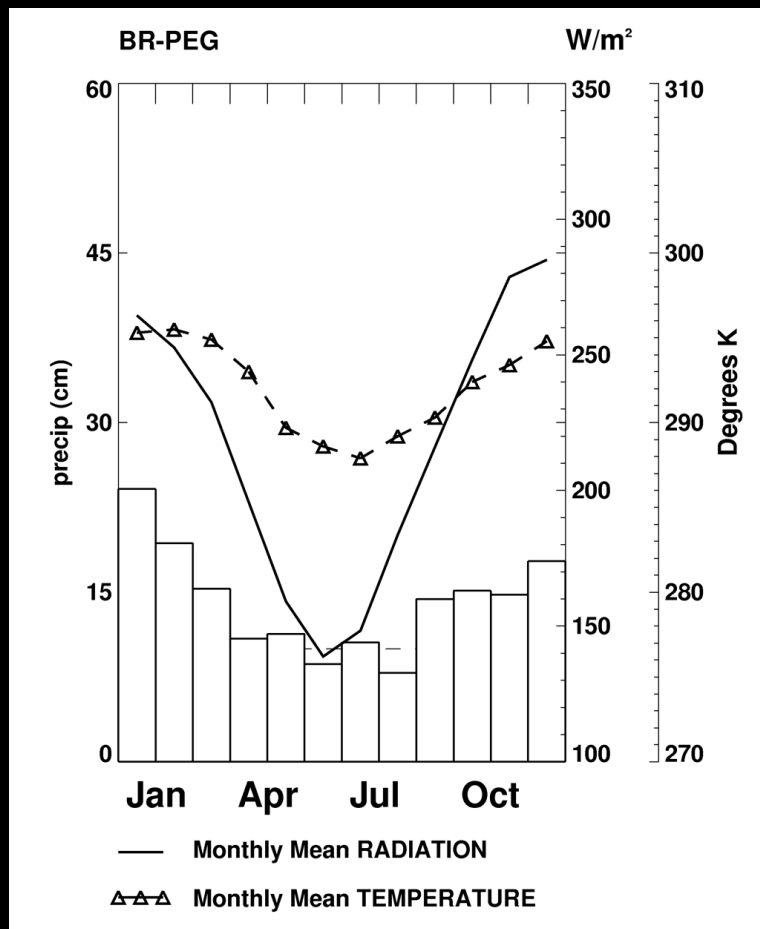
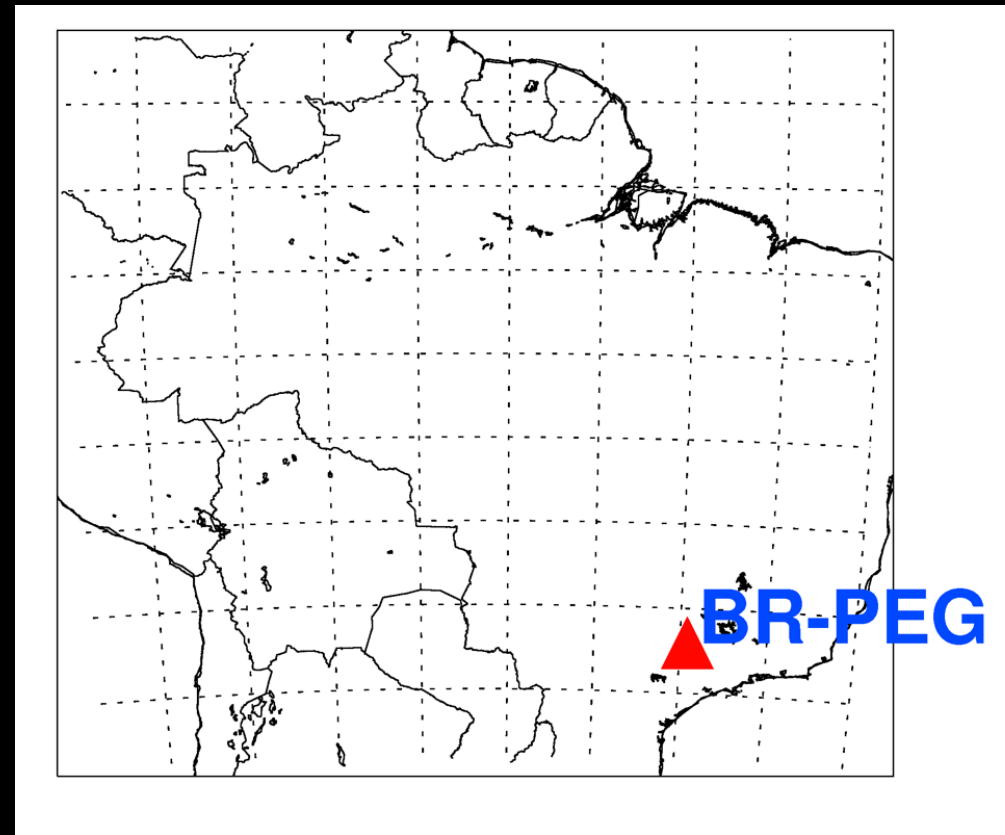
Simple Biosphere Model, version 3.0



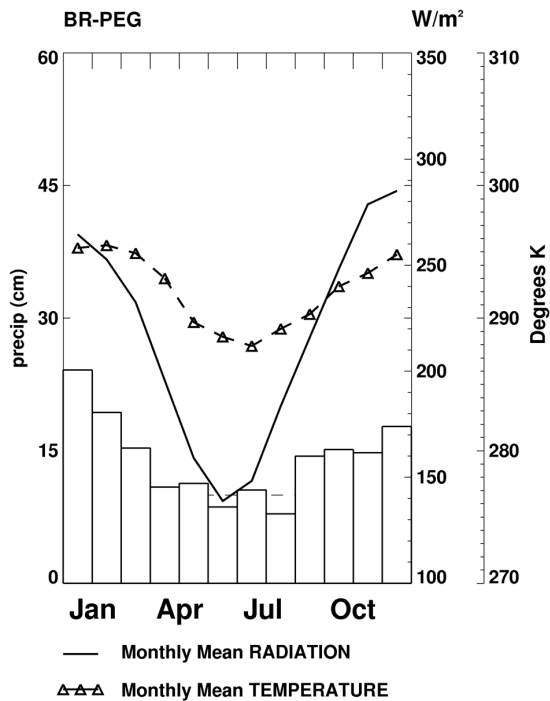
SiB3; Baker et al., 2003, 2008, many previous back to Sellers et al., 1986



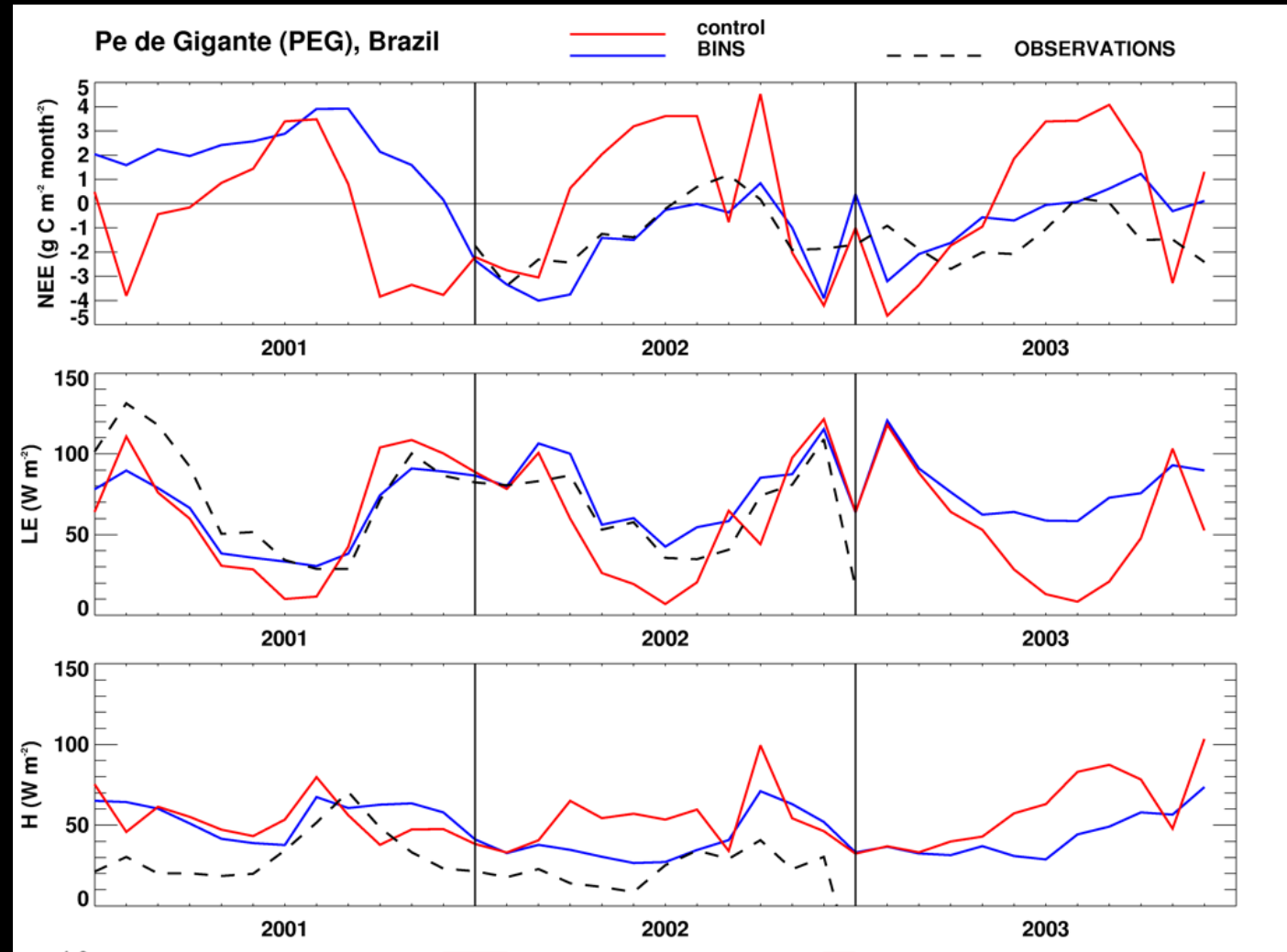
# Does It Work?



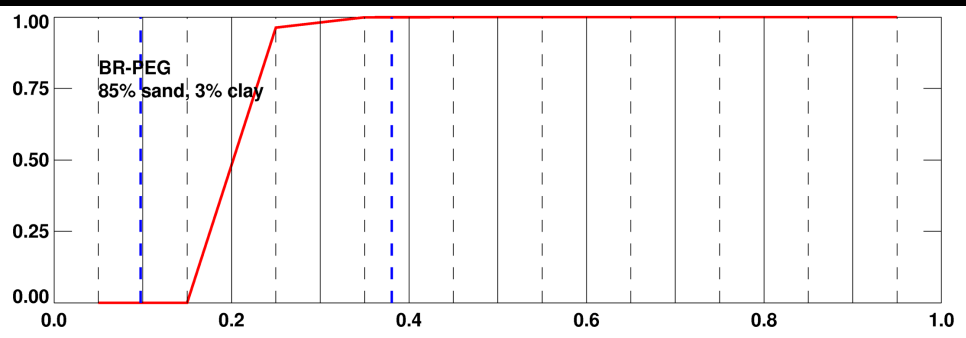
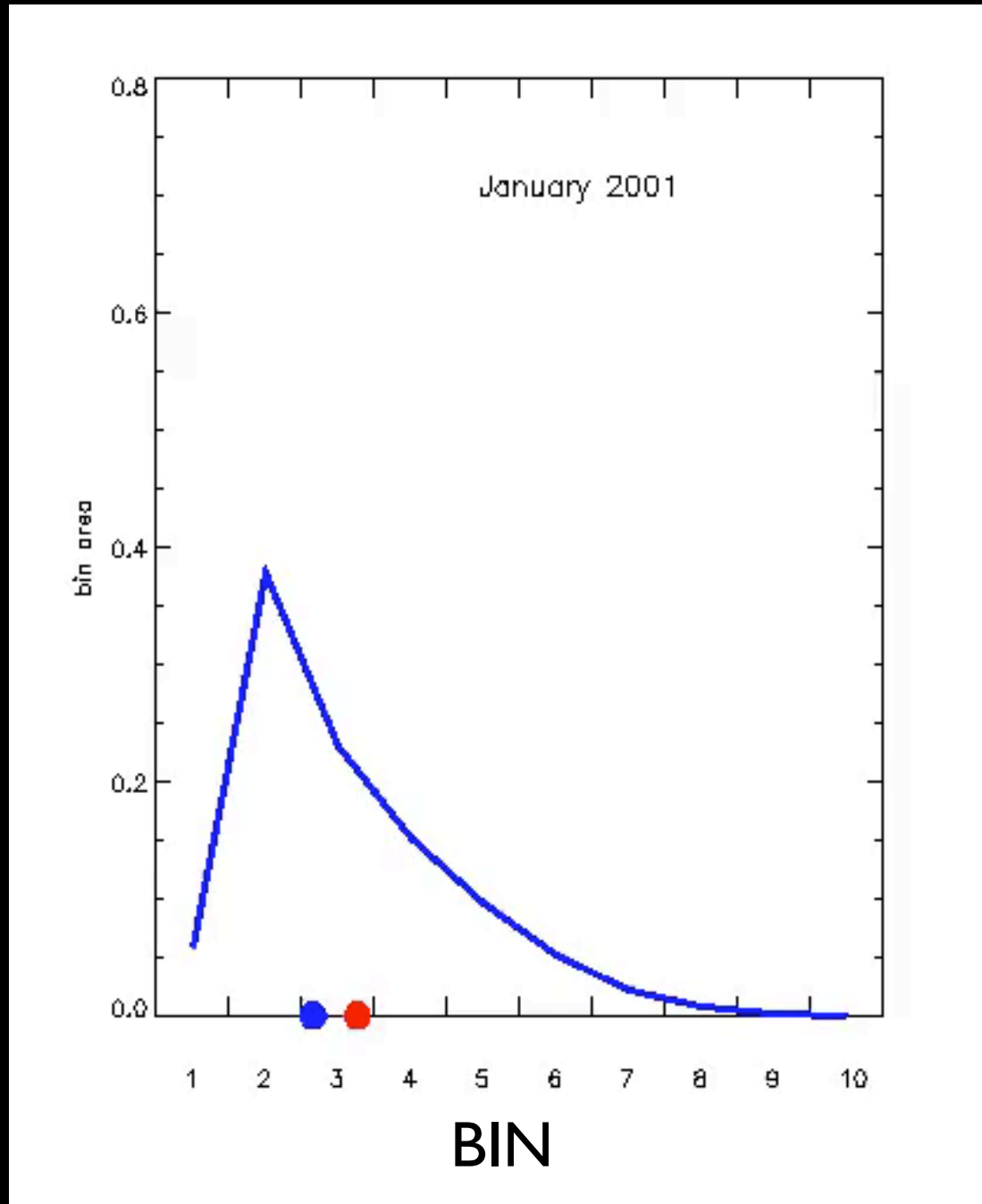
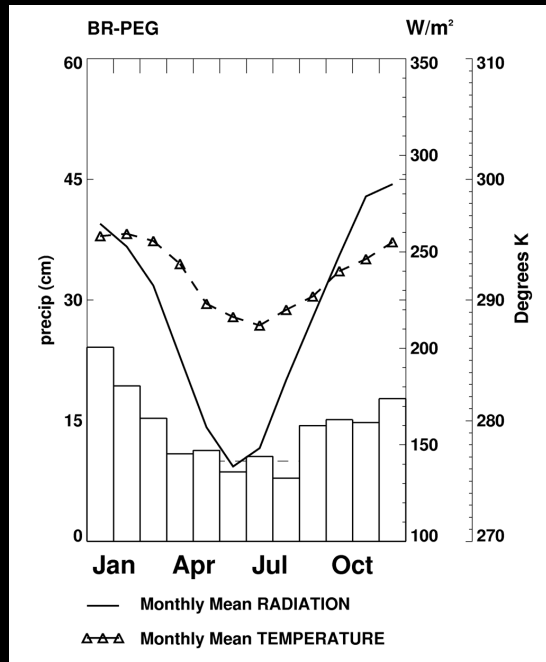
# Does It Work?



- Comparison to observations improves with BINS
- Control run (Method II) has large excursions when  $W$  is on the steep segment of  $f(W)$  curve



# BIN Area Evolution



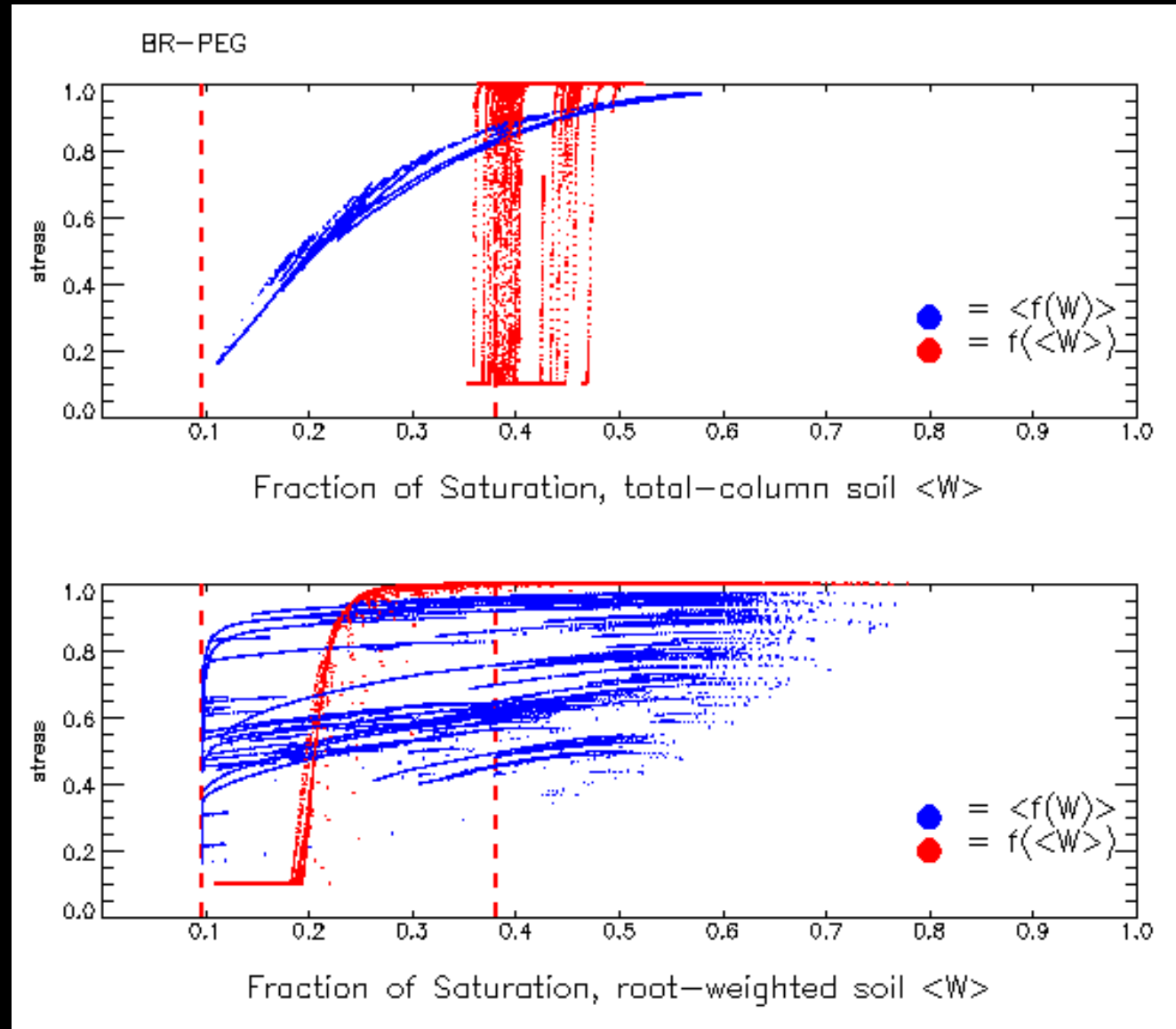


# Summary

- The BINS scheme improves simulated surface fluxes when compared to control runs using a single soil wetness value to predict stress
- Numerical scheme is stable: BINS balances energy/water to machine precision
- BINS: convective precipitation
- High potential for grassland (semi-arid to arid)

# Total Soil vs. Root-Weighted Column

- Dots represent column-mean wetness for control (method II-**RED**) and BINS (**BLUE**)
- Each dot corresponds to a single hour during a 3-year simulation
- **BINS** behaves smoothly when the whole column is considered
- **Method II** follows the  $f(W)$  curve in the root-weighted soil





# Soilmoisture Stress Calculation

