

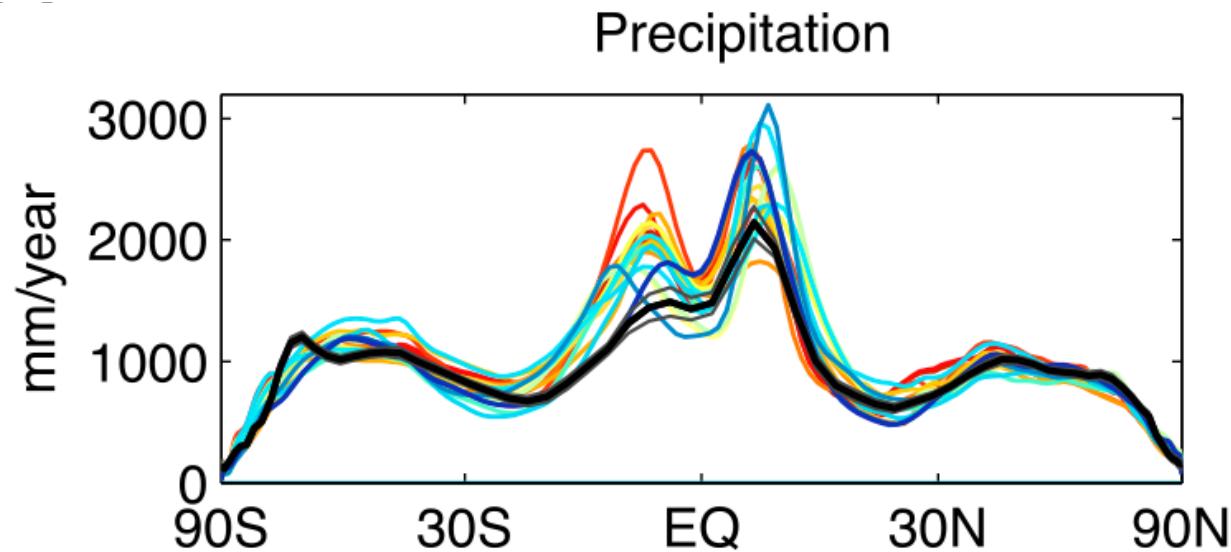
Convection, Tropical Waves, and Double ITCZs

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Thanks to John Chiang (UC Berkeley)
Mike Pritchard (UC Irvine)
Zhihong Tan (Caltech)

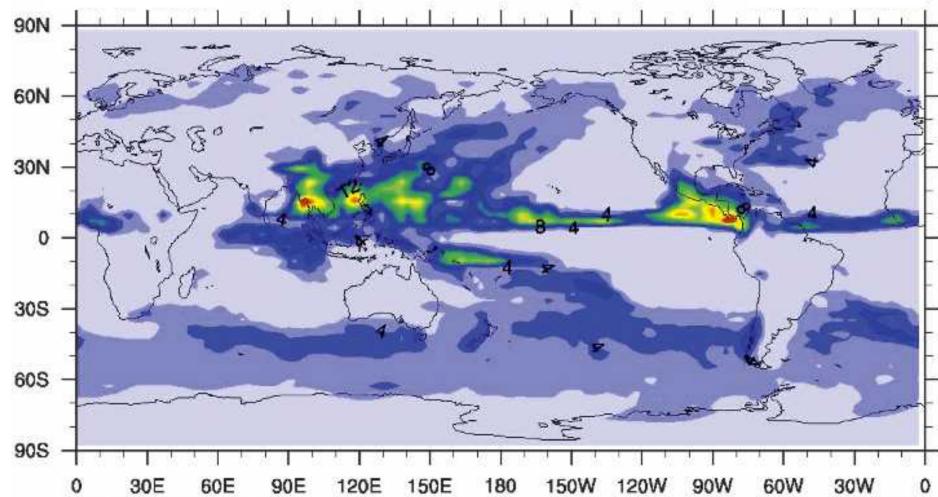
Double Inter-tropical Convergence Zones (ITCZs)



Black: Observation

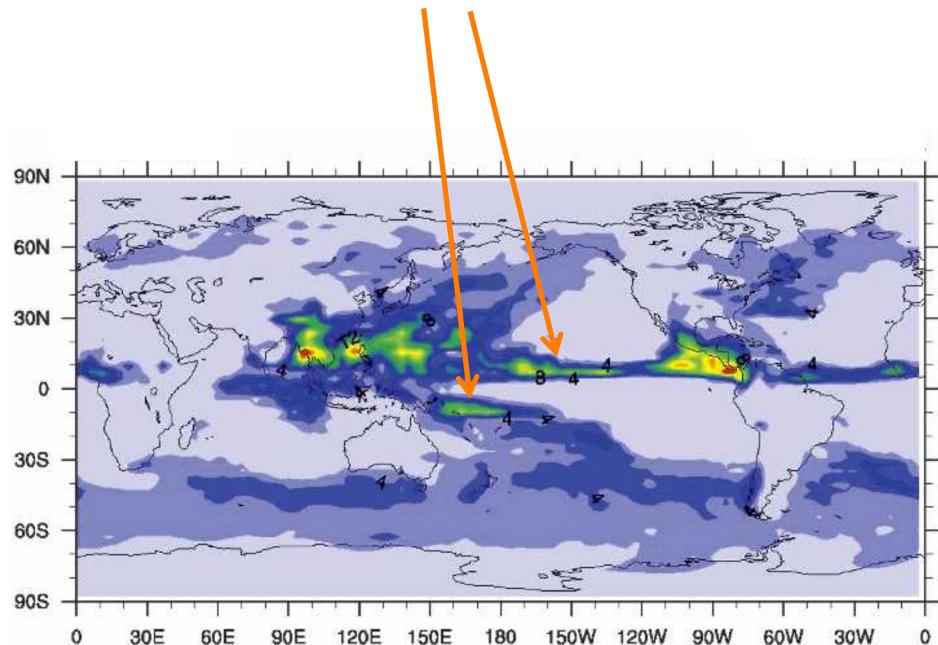
Others: CMIP simulation results

Double ITCZs are simulated by the SPCAM



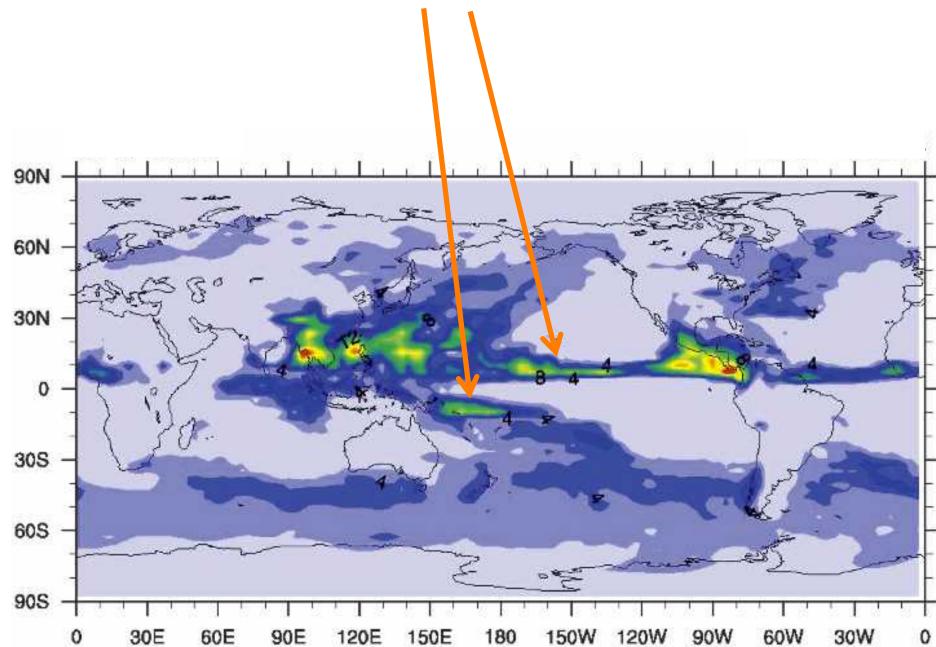
Khairoutdinov et al. 2005

Double ITCZs are simulated by the SPCAM

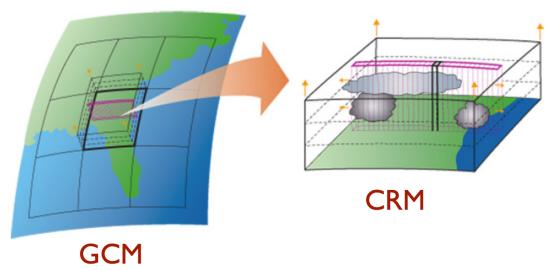


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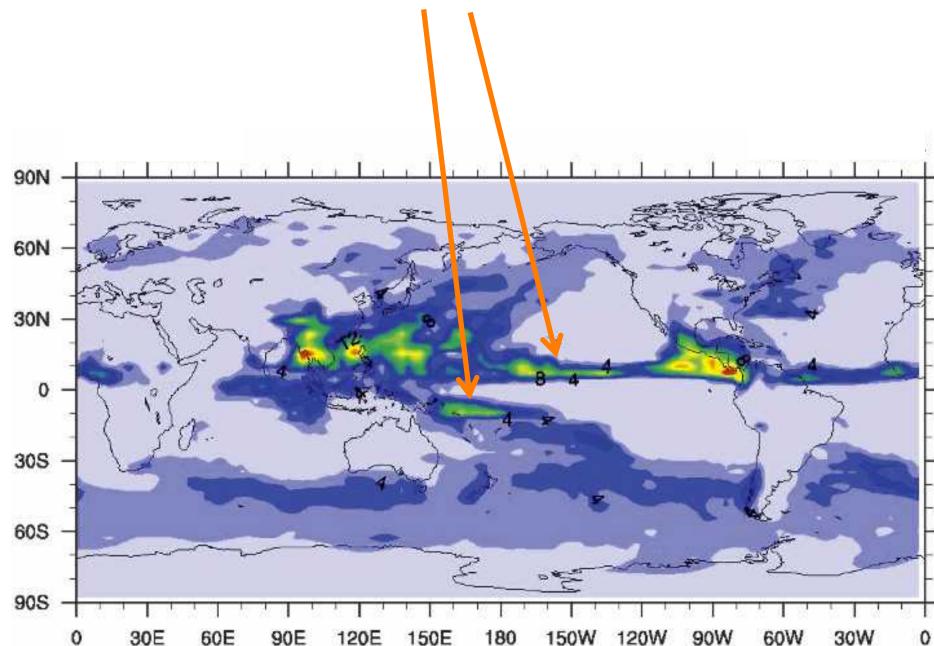
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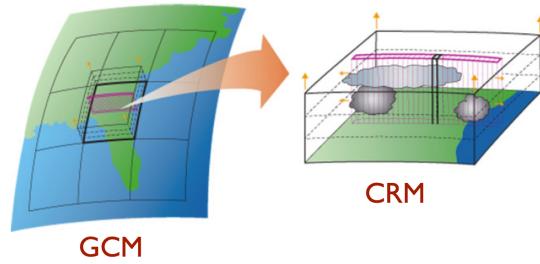
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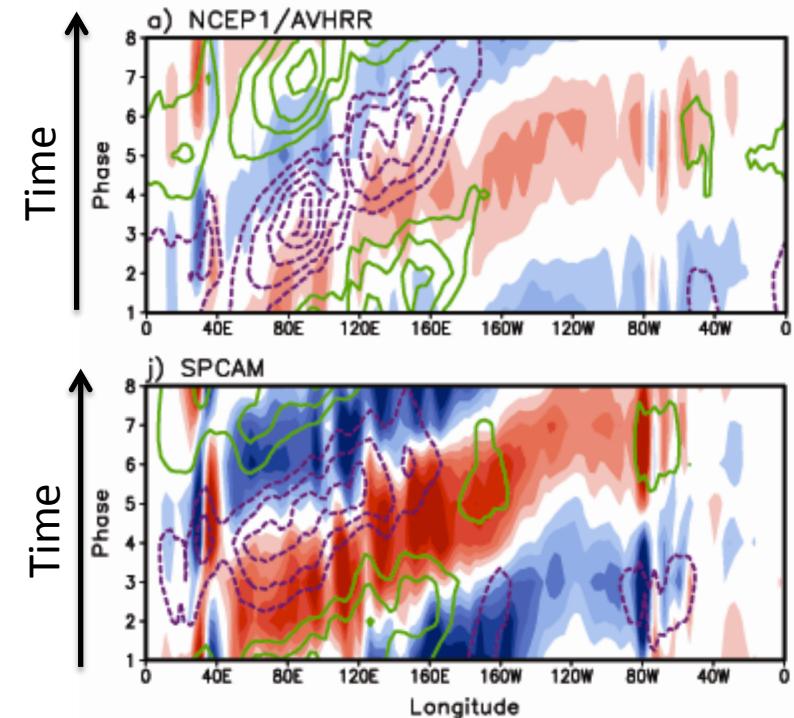
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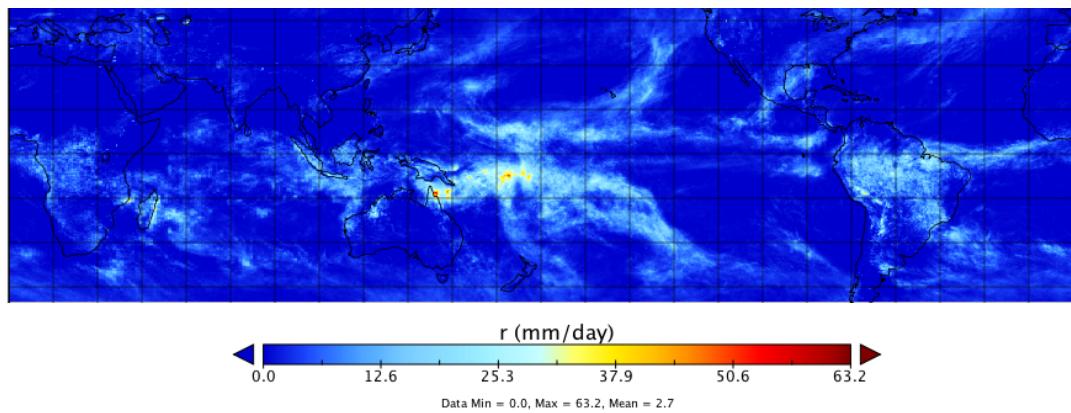
SPCAM simulates even
stronger MJOs than observation.



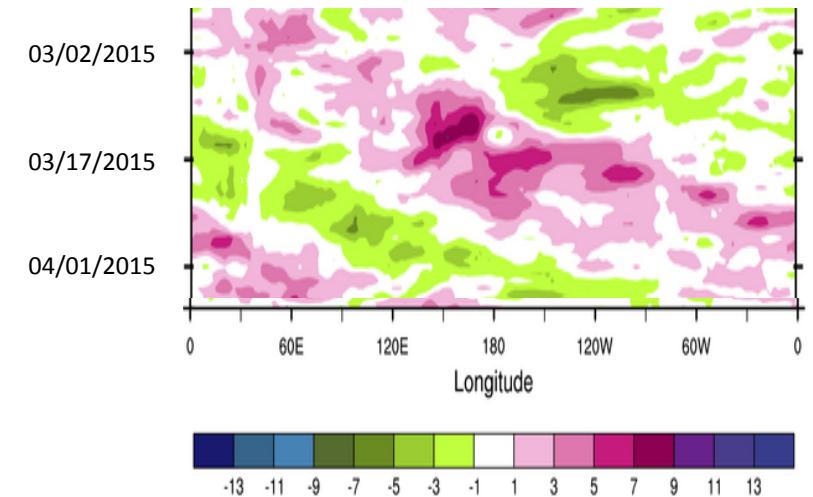
Kim et al. 2009

Double ITCZs emerged during a strong Madden-Julian Oscillation event in March 2015

Daily mean precipitation 03/05/2015–03/20/2015

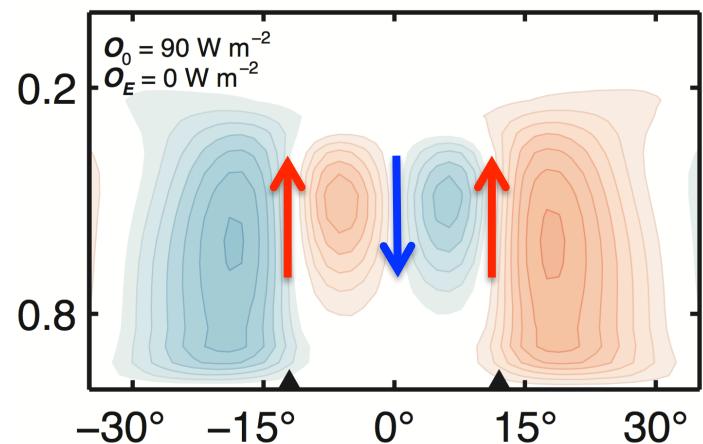


Westerly winds (m/s)



Why double ITCZs?

Anomalous anti-Hadley circulations



Arguments about ITCZ location:

- A1: **SST distribution**: e.g., Lindzen and Nigam 1987; Neelin 1989; Wang and Li 1993; Chiang et al. 2001
- A2: **Quasi-equilibrium**: e.g., Emanuel 1995; Prive and Plumb 2007; Boos and Kuang 2010
- A3: **Energy transport**: e.g., Kang et al. 2008, 2009; Frierson and Hwang 2012; Donohoe et al. 2013; Bischoff and Schneider 2015

A thought experiment

- Imagine ...
 - Aquaplanet
 - Forced by uniformly distributed sea surface temperatures
(No baroclinic waves)
 - No Sunlight (only longwave radiation)

A thought experiment

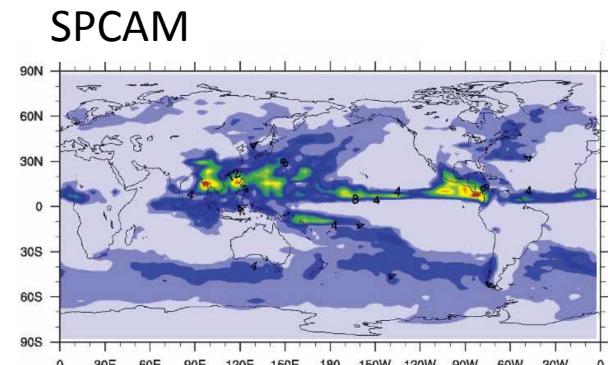
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 - Are there still tropical rainfall peaks?
 - Where are they?

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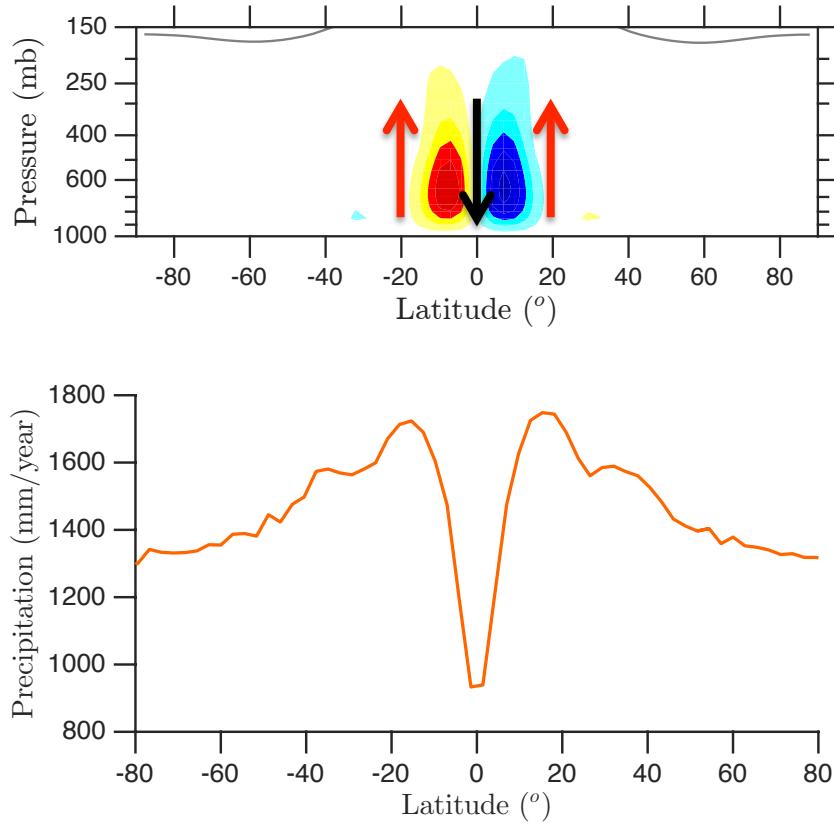
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- An educated guess based on:
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Double ITCZs emerge over uniform SSTs



This result suggests:

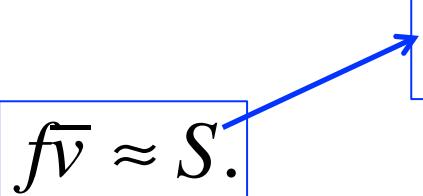
- Thermodynamic constraints cannot predict this behavior.
- We need other constraints.

Proposed mechanism:
**Tropical waves can drive double ITCZs
by transporting angular momentum**

At steady state, when $R_o \ll 1$,

$$\bar{fv} \approx S.$$

eddy momentum
flux divergence

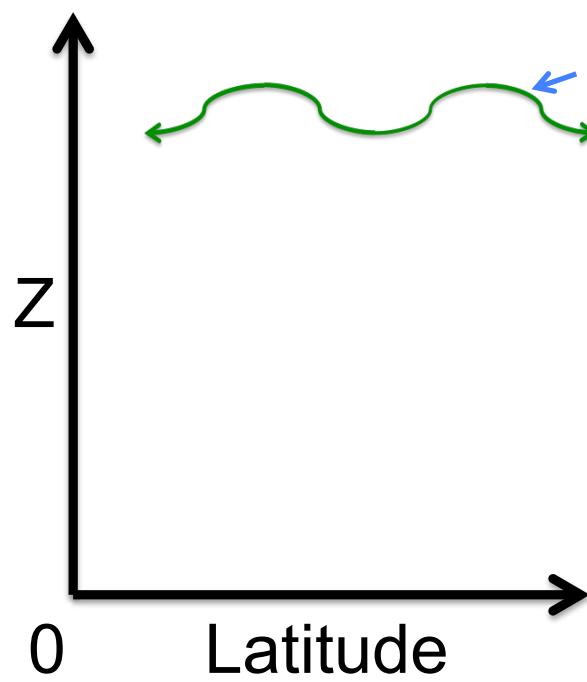


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Tropical waves converge angular momentum,
i.e., $S < 0$,

$$\bar{f}\bar{v} < 0$$

\Leftrightarrow Equatorward flow

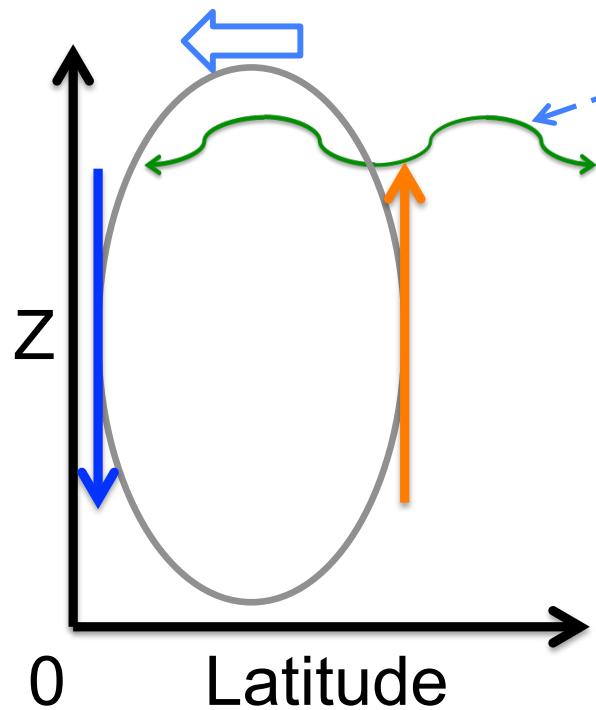
Schneider 2006; Bordoni and Schneider 2008; Shaw 2014;

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$$\begin{aligned} \bar{f}\bar{v} &< 0 \\ \Leftrightarrow & \text{ Equatorward flow} \end{aligned}$$

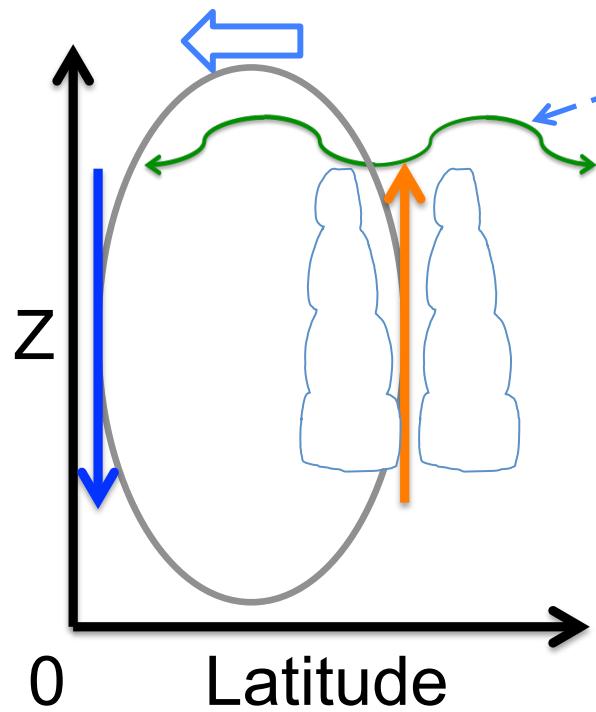
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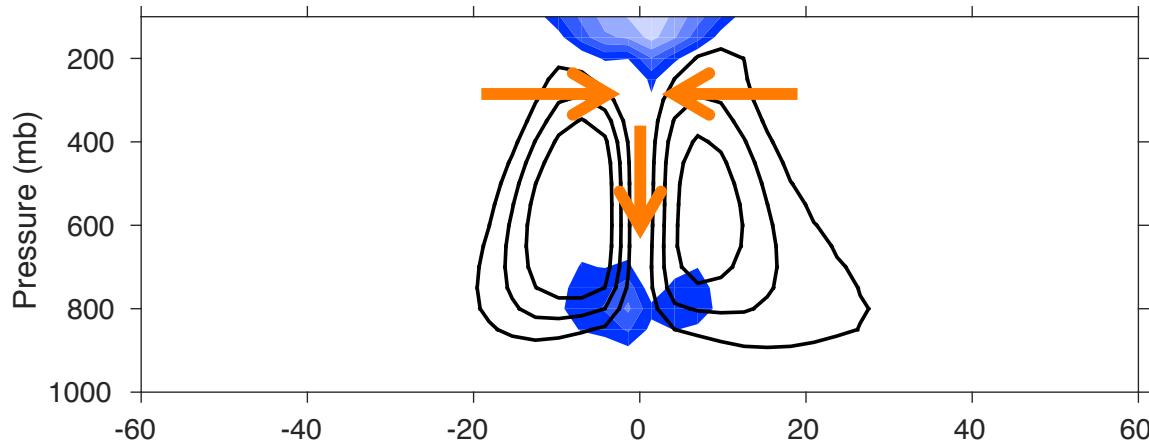
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In the upper troposphere, $\bar{fv} \sim S < 0$.

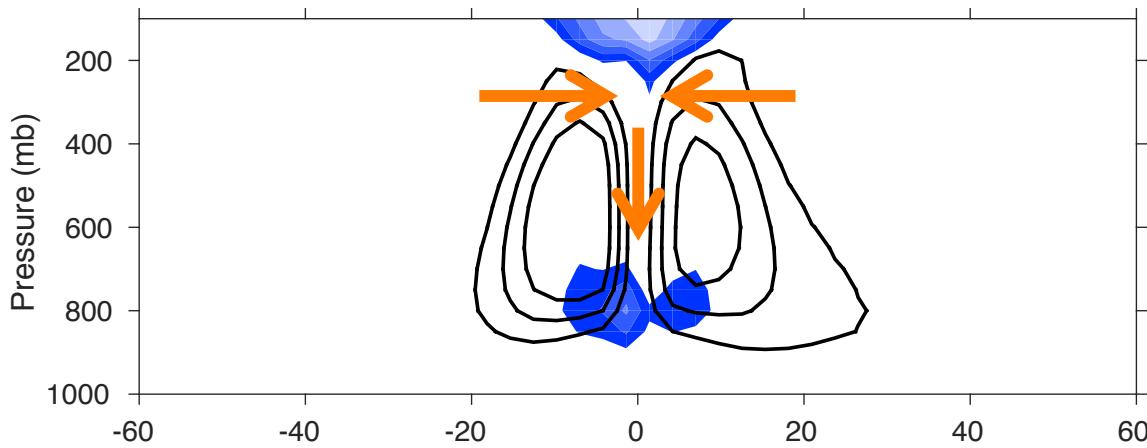
Angular momentum argument should work



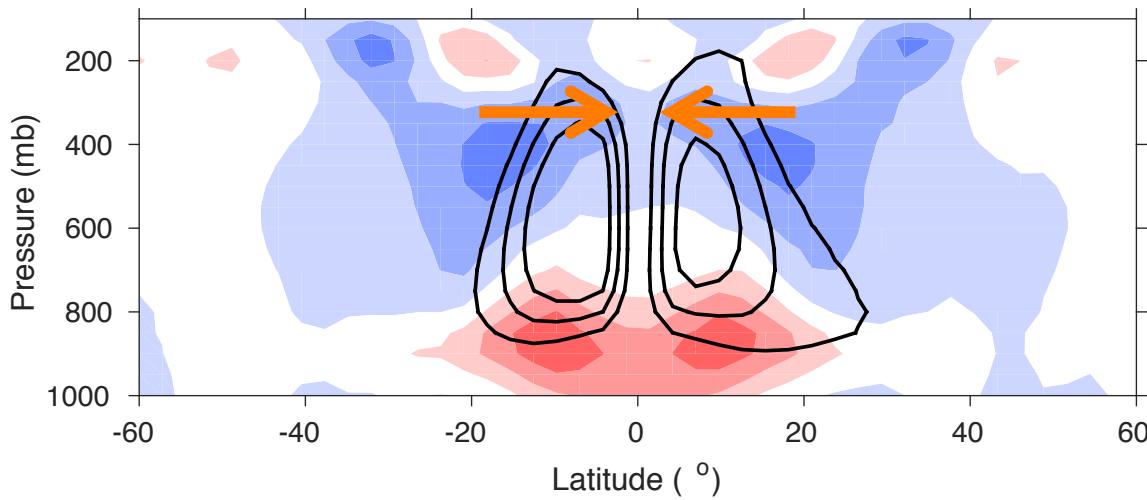
color shading: $R_o > 0.2$

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Angular momentum argument should work



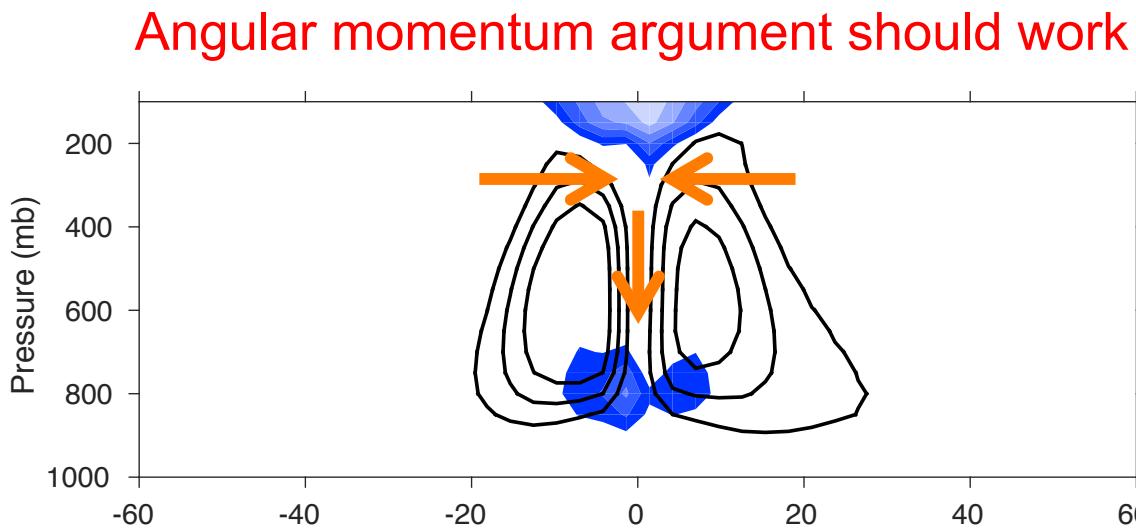
AM flux convergence requires equatorward flow



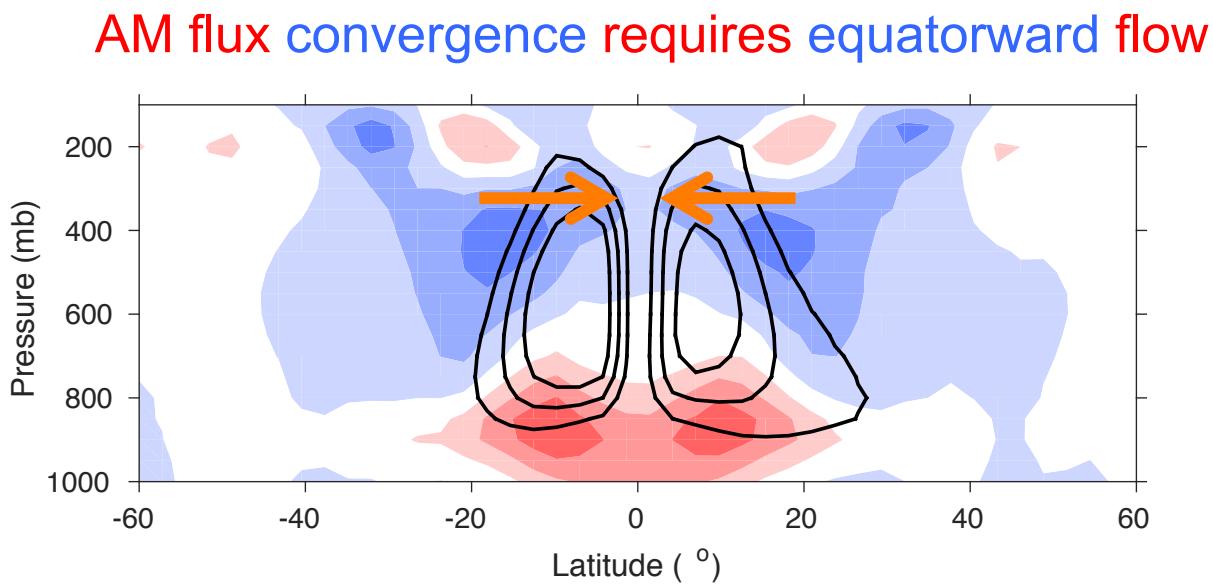
In the upper troposphere,

$$f\bar{v} \sim S < 0.$$

Where is this from?



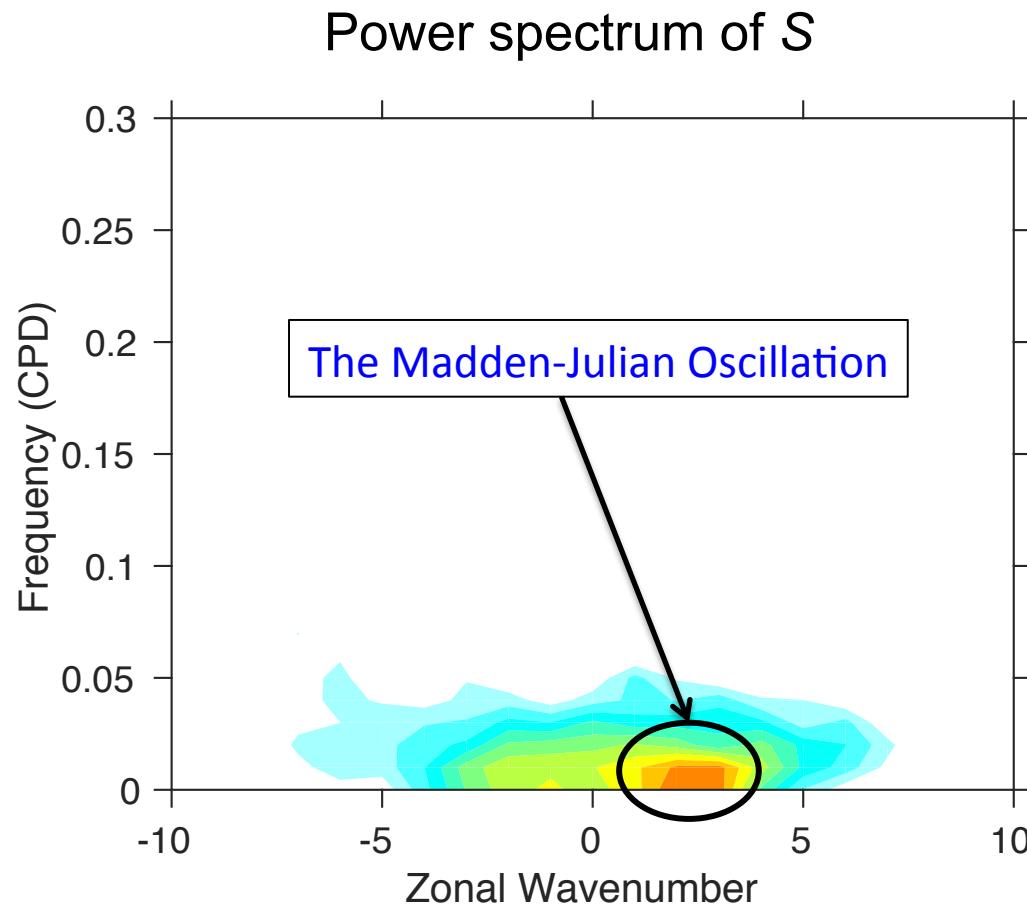
color shading: $R_o > 0.2$



color shading: S

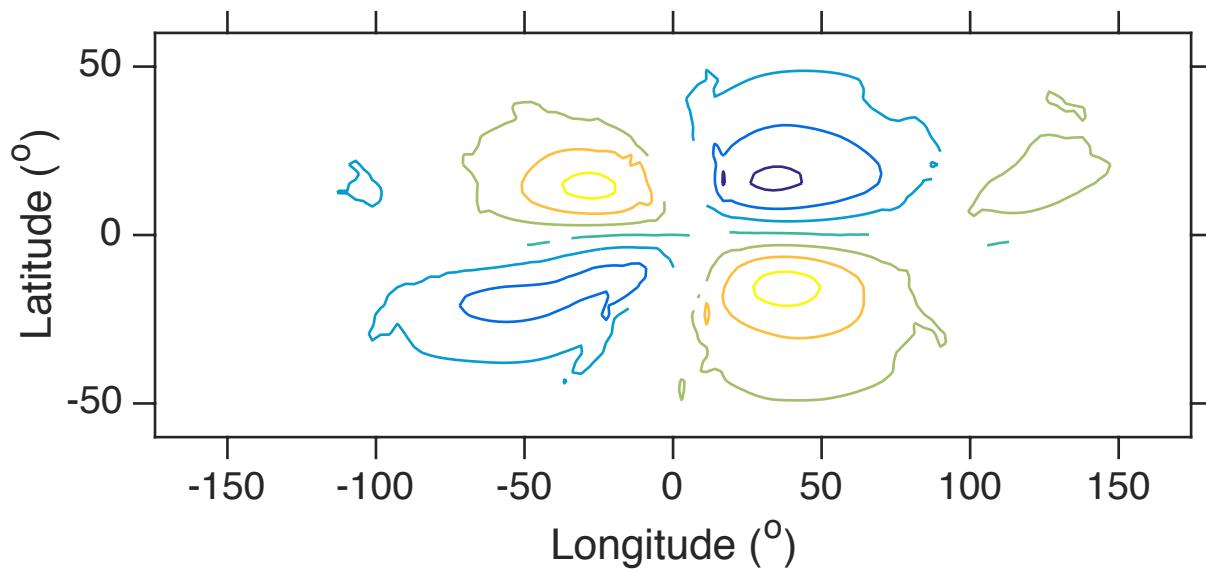
Blue: convergence, $S < 0$
Red: divergence, $S > 0$

The Madden-Julian Oscillation dominates the momentum transport.



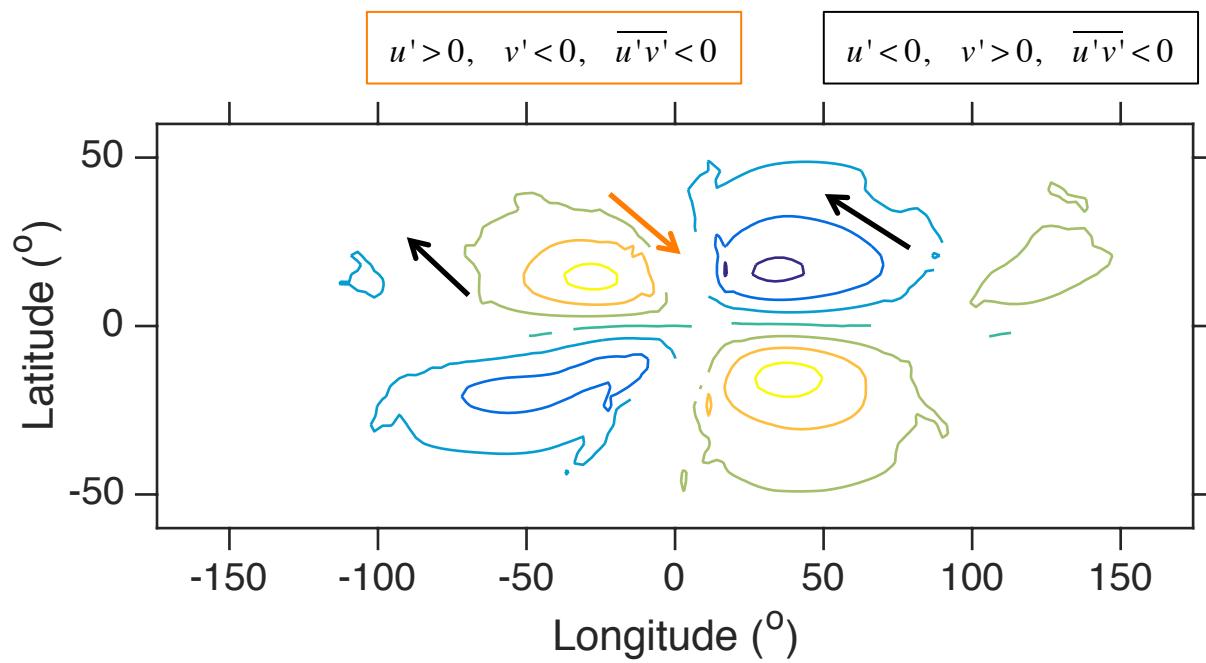
Consistent with studies under more realistic setup,
e.g., Lee 1999, Caballero and Huber 2010, Arnold et al. 2012

Momentum transport by the MJO



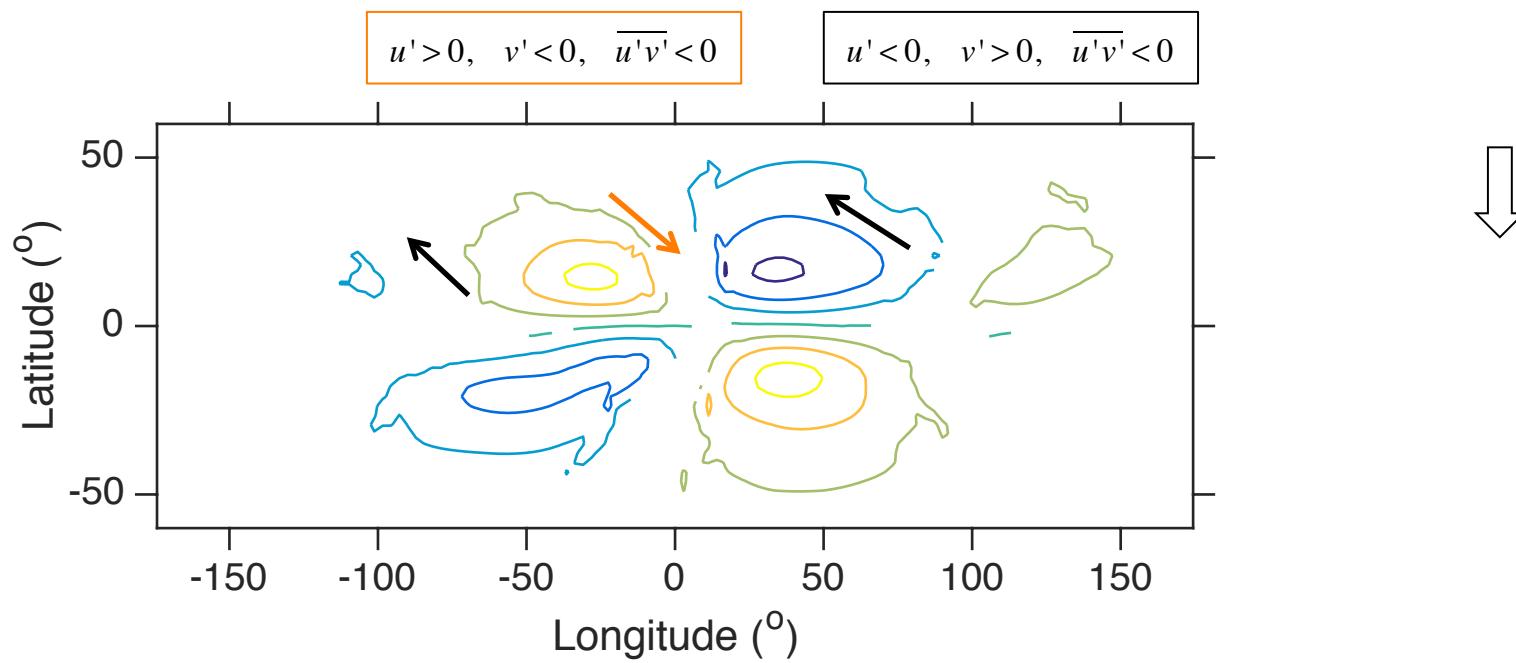
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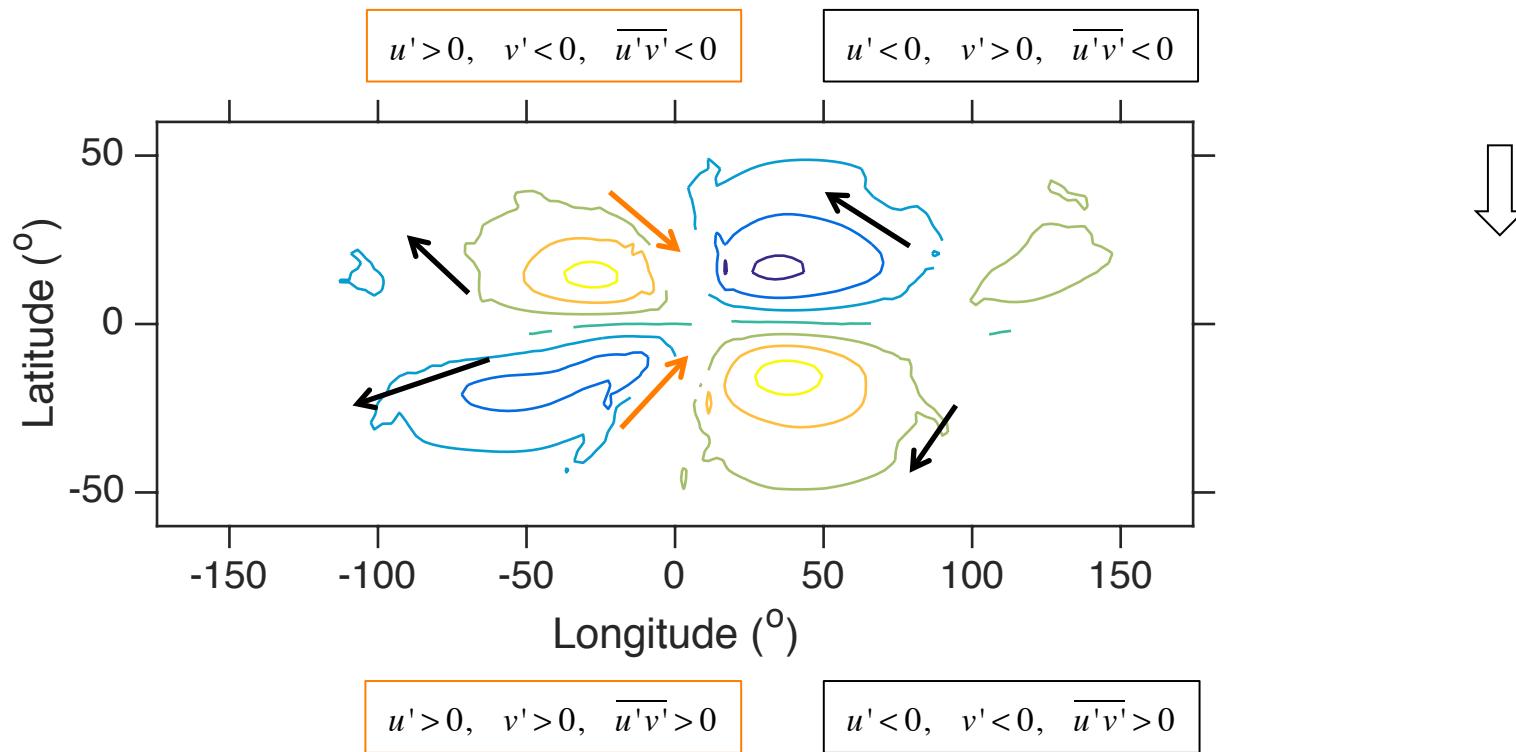
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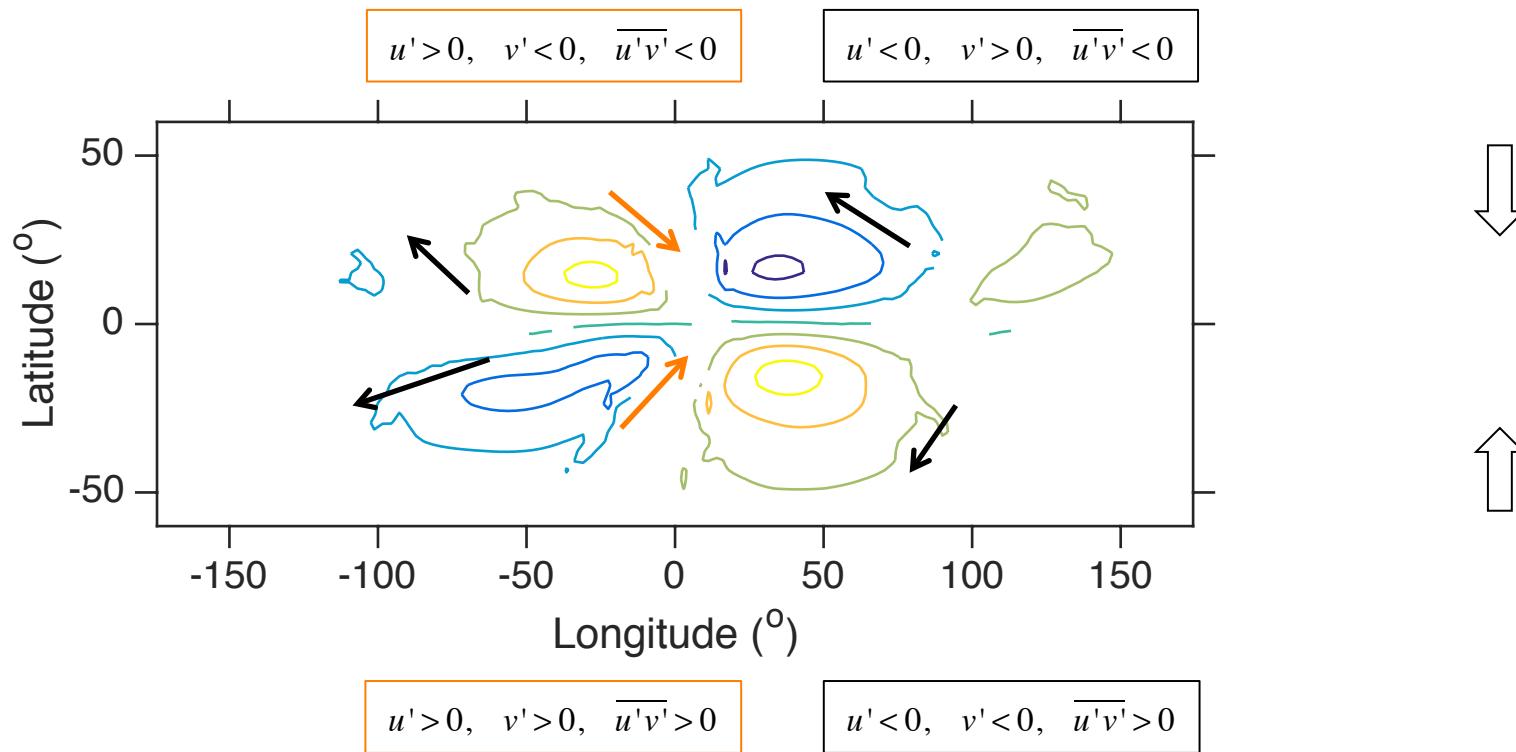
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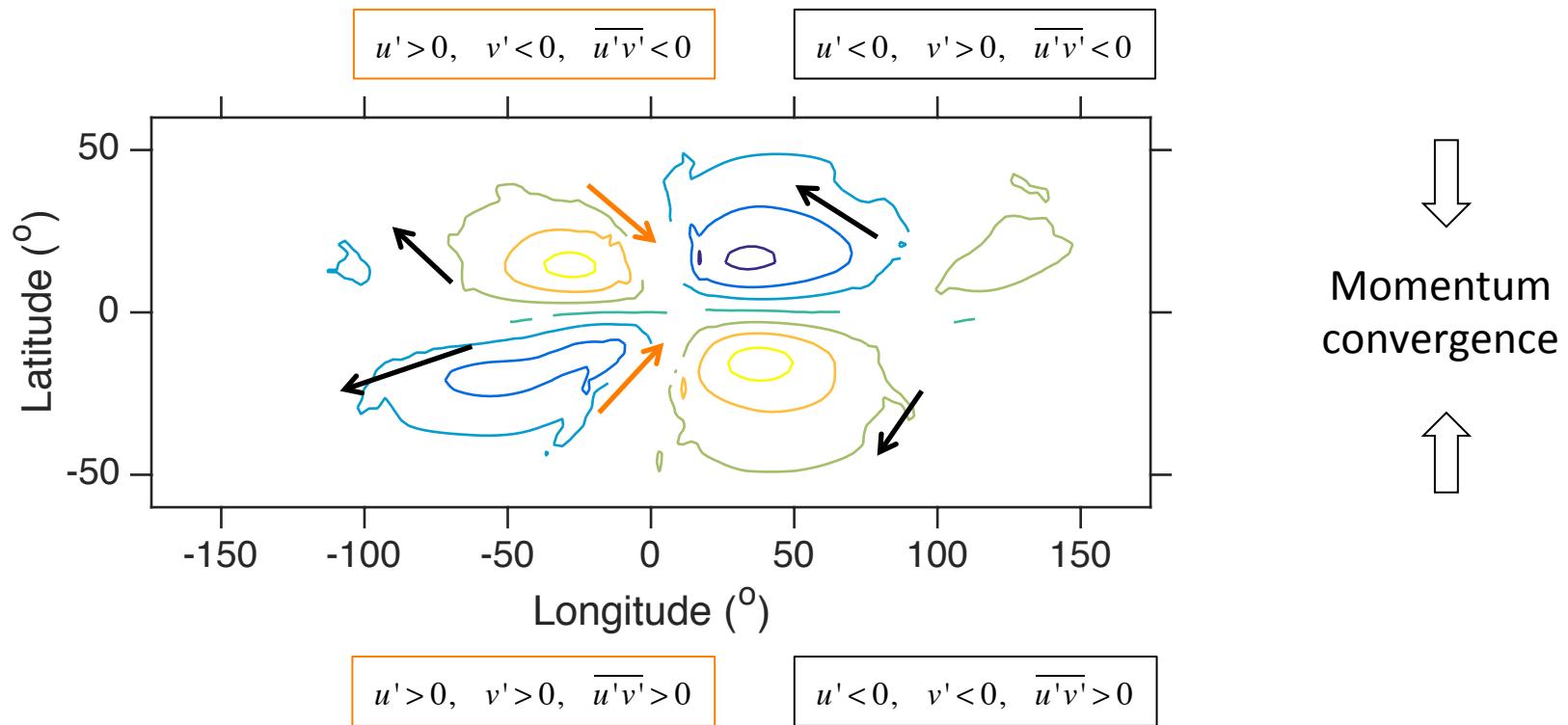
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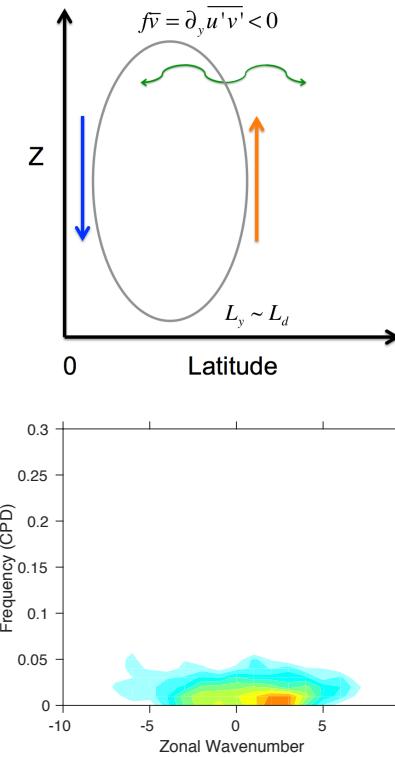
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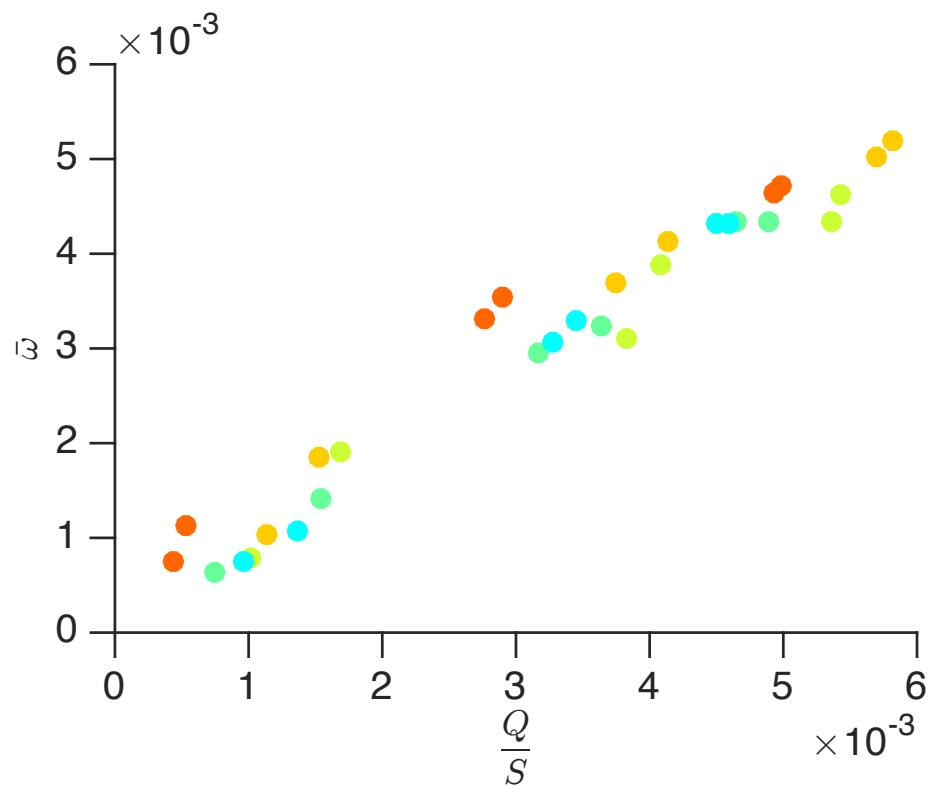
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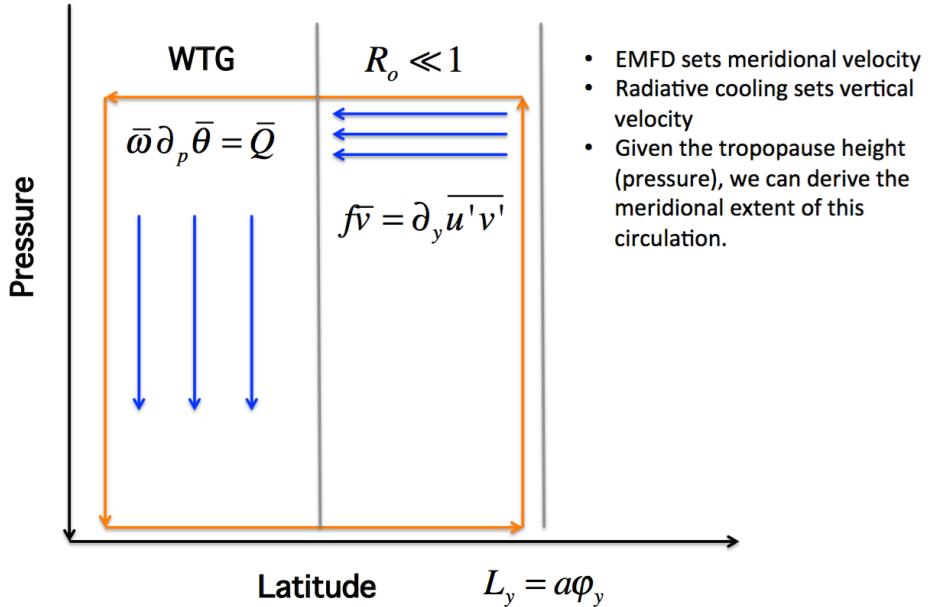
Summary

- Double ITCZs are simulated over uniform sea surface temperatures in SPCAM.
 - NOT expected from the thermodynamic arguments.
- The angular momentum argument can explain this result.
- The Madden-Julian Oscillation dominates the meridional eddy momentum transport.
- Double ITCZs in SPCAM might be due to strong MJO signals.
 - When the equatorial wave activity (e.g., the MJO and Rossby waves) is strong, this proposed mechanism can produce double ITCZs.



Weak temperature gradient approximation holds globally





$$L_y \sim \left(\frac{\Delta p \partial_p \bar{\theta} \partial_y \bar{u}' v'}{\beta \bar{Q}} \right)^{1/2}$$

Figure 1: Schematics of the proposed mechanism of the anti-Hadley circulation. The left and right boxes are connected by continuity.

$$f\bar{v} = \partial_y \bar{u}' v' \quad (2)$$

$$\bar{\omega} \partial_p \bar{\theta} = \bar{Q} \quad (3)$$

$$\partial_y \bar{v} + \partial_p \bar{\omega} = 0 \quad (4)$$

From (4), we know

$$\frac{|\omega|}{|v|} \sim \frac{\Delta p}{L_y} \quad (5)$$

Here L_y is the meridional extent of this tropic cell, and $\Delta p = p_s - p_t$. Combining (2) & (3), we get

$$\frac{\omega}{v} \sim \frac{fQ}{\partial_y u' v' \partial_p \bar{\theta}} \sim \frac{2\Omega \sin \varphi_c \bar{Q}}{\partial_y u' v' \partial_p \bar{\theta}} \sim \frac{2\Omega \varphi_c \bar{Q}}{\partial_y u' v' \partial_p \bar{\theta}} \quad (6)$$

MJO in SAM

SST = 290 K, 1/3 of Earth's circumference

