

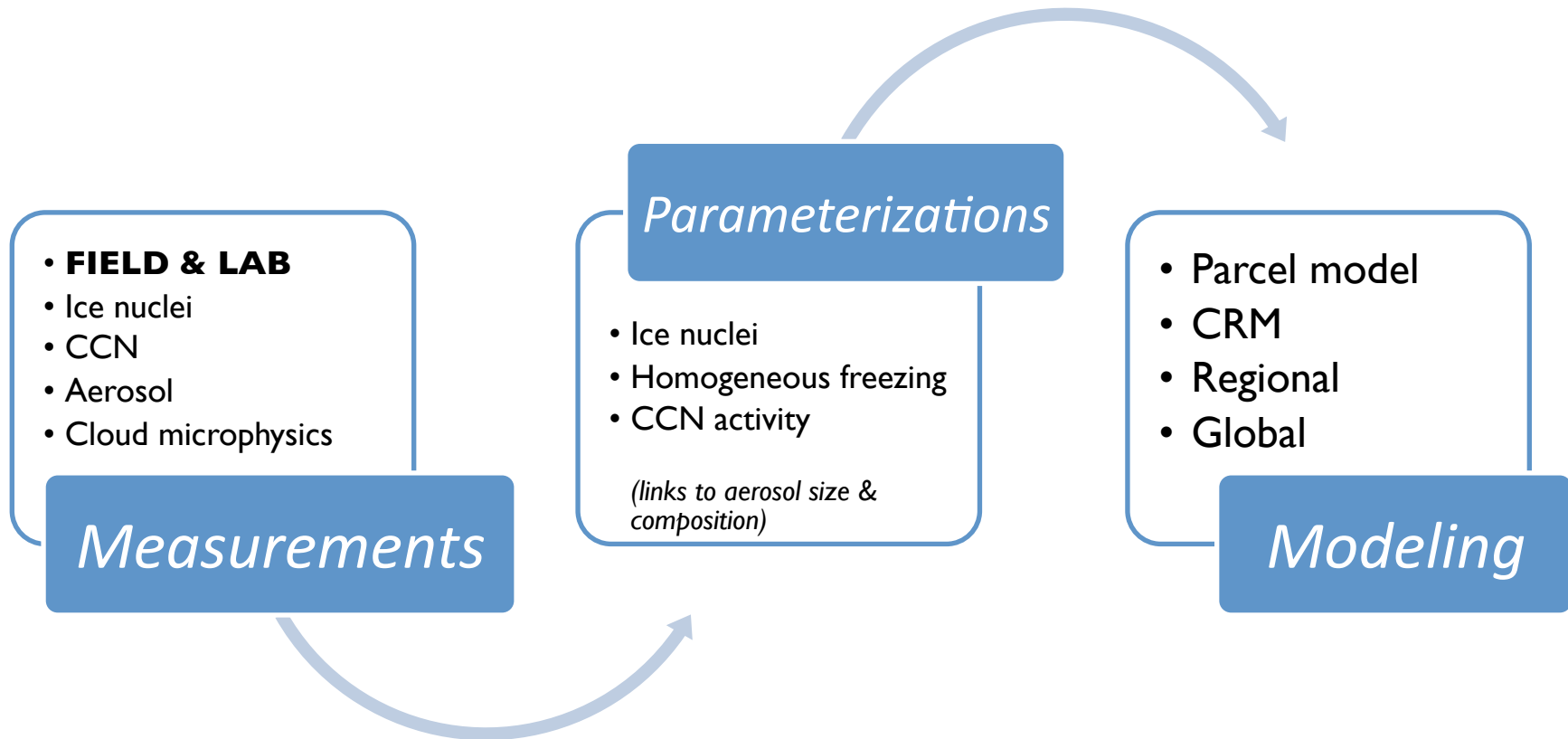
# **Predicting global atmospheric ice nuclei distributions and their impacts on climate**

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S. M Kreidenweis, T. Eidhammer, C. H. Twohy,  
M. S. Richardson, and D. C. Rogers

*Acknowledgments:*

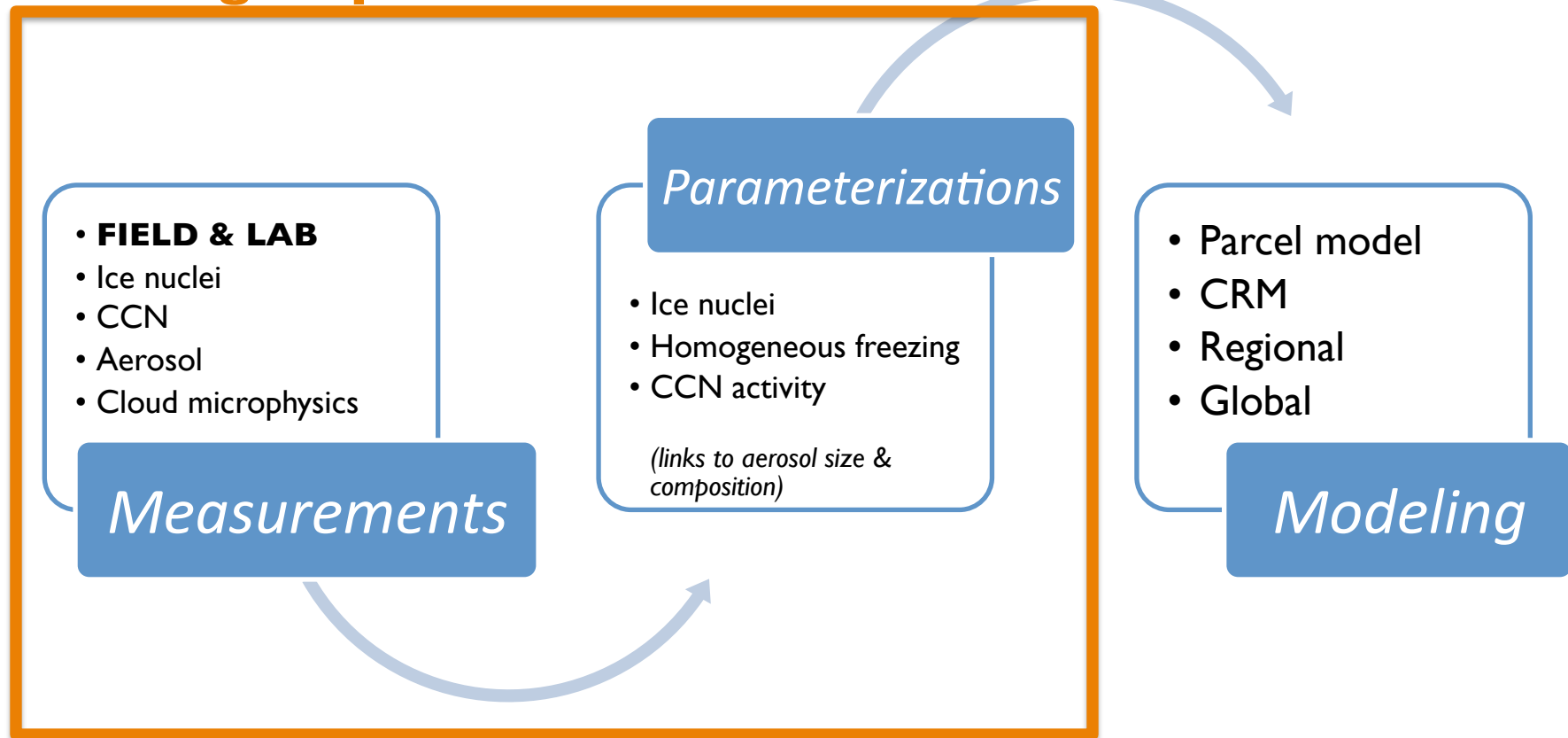
DOE-ARM (Grant No. DE-FG02-09ER64772), NASA-MAP (Grant NNG06GB60G),  
NASA New Investigator Program, DOE Climate Change Prediction Program,  
NSF (various), CIRA (CLEX-10)

# Approach

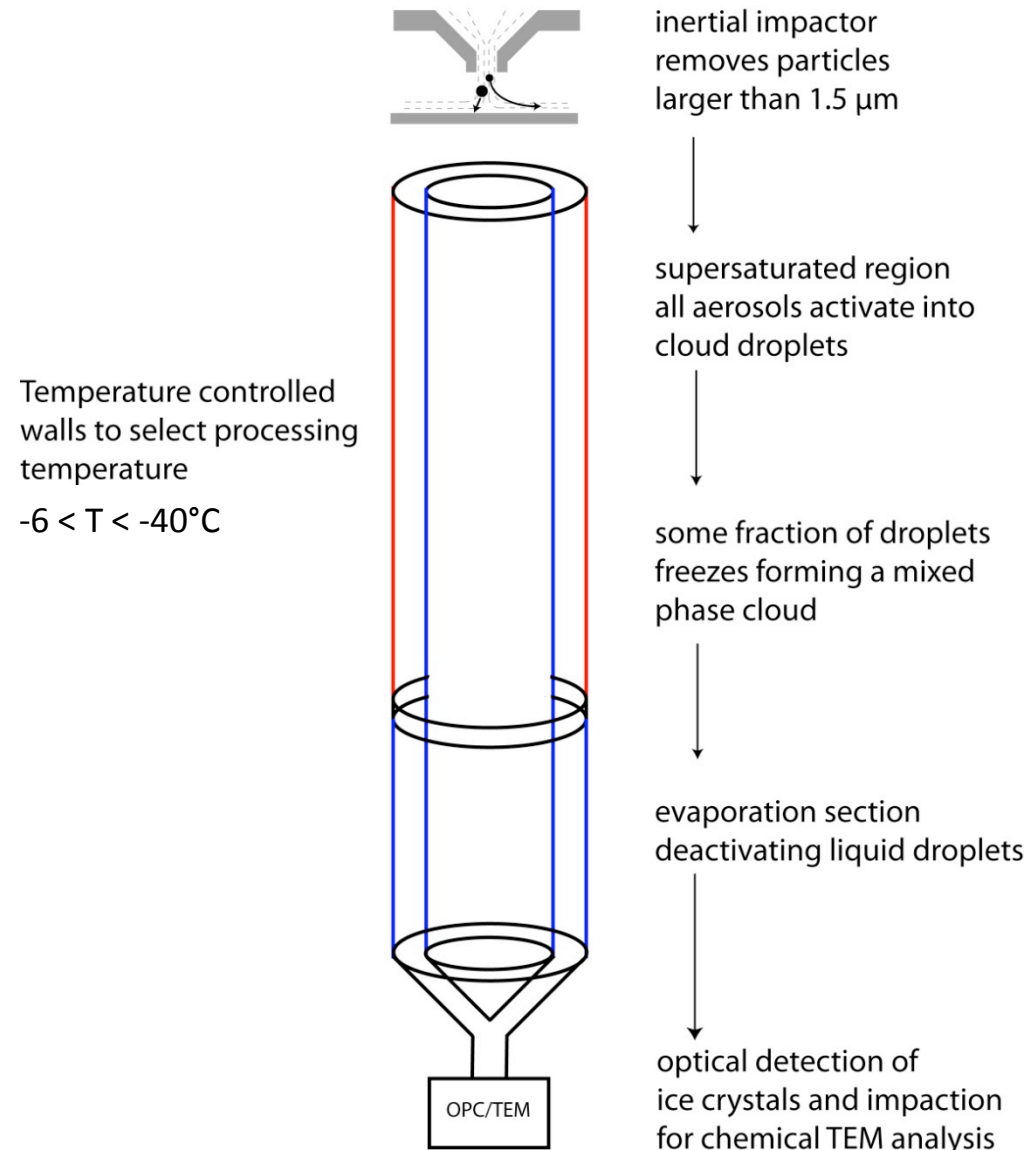


# Approach

## Our group

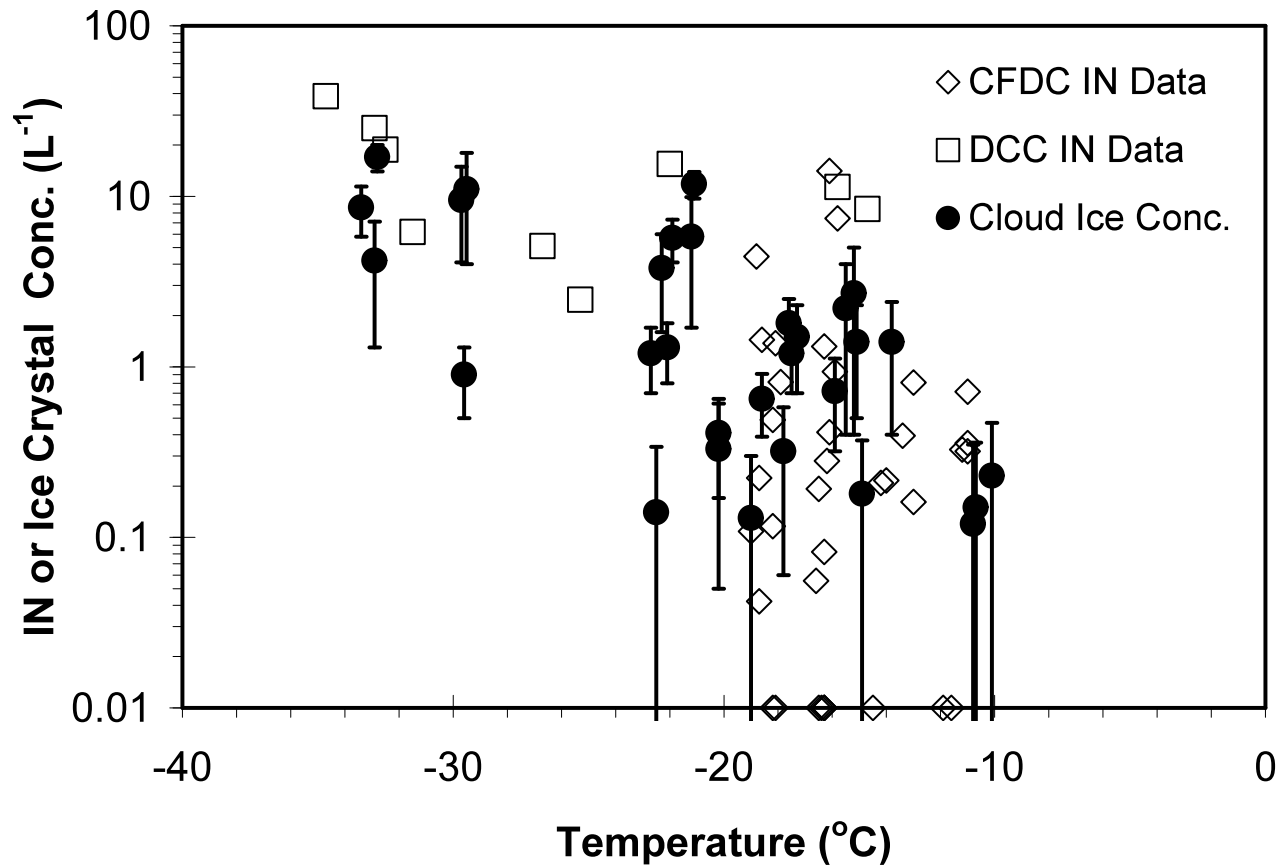


# Real-time atmospheric measurement of IN - Continuous flow diffusion chamber (CFDC)



Total residence time  $\sim 6\text{s}$

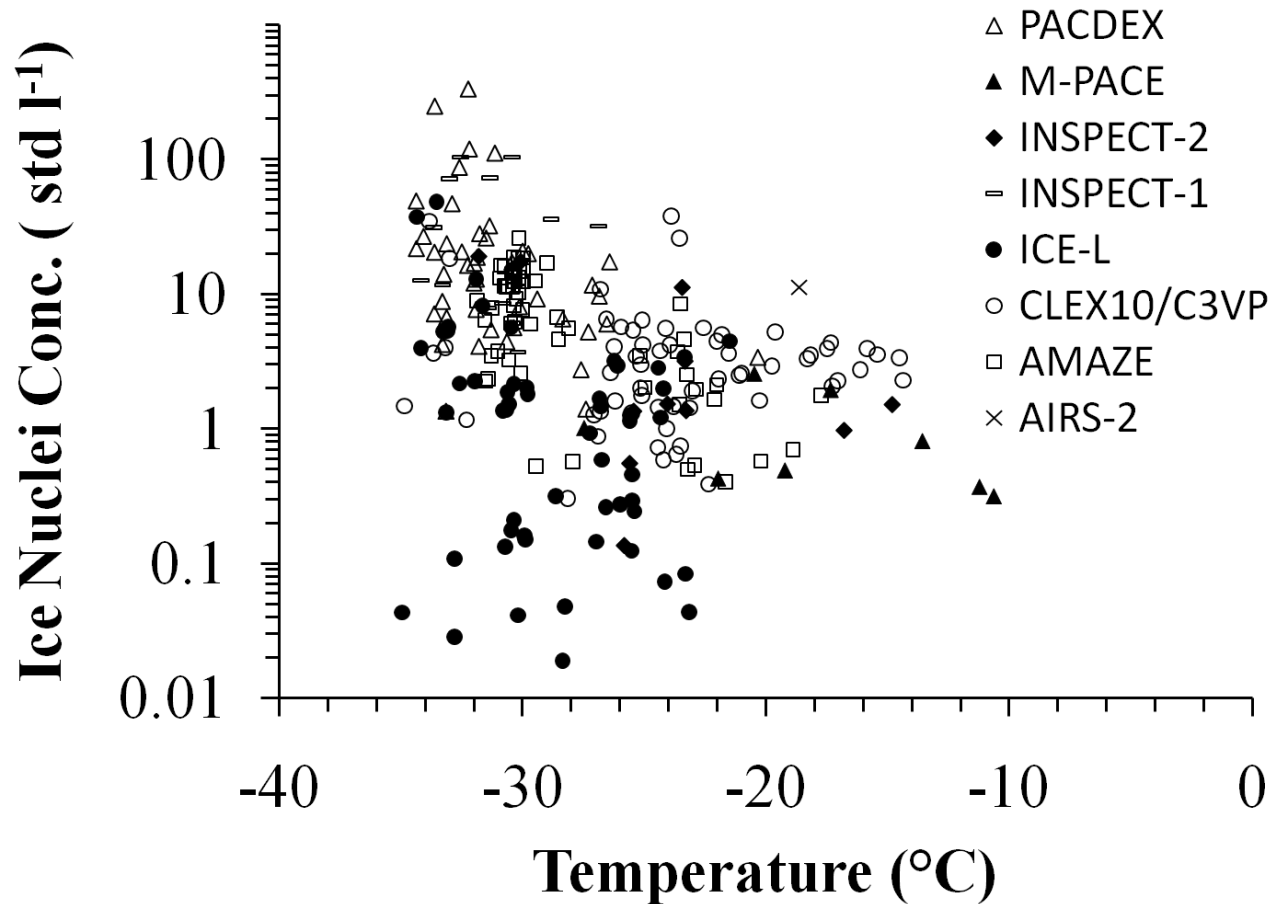
IN that we measure **DO** represent  
(primary) ice concentrations in clouds



WISP-1994  
data shown  
as example

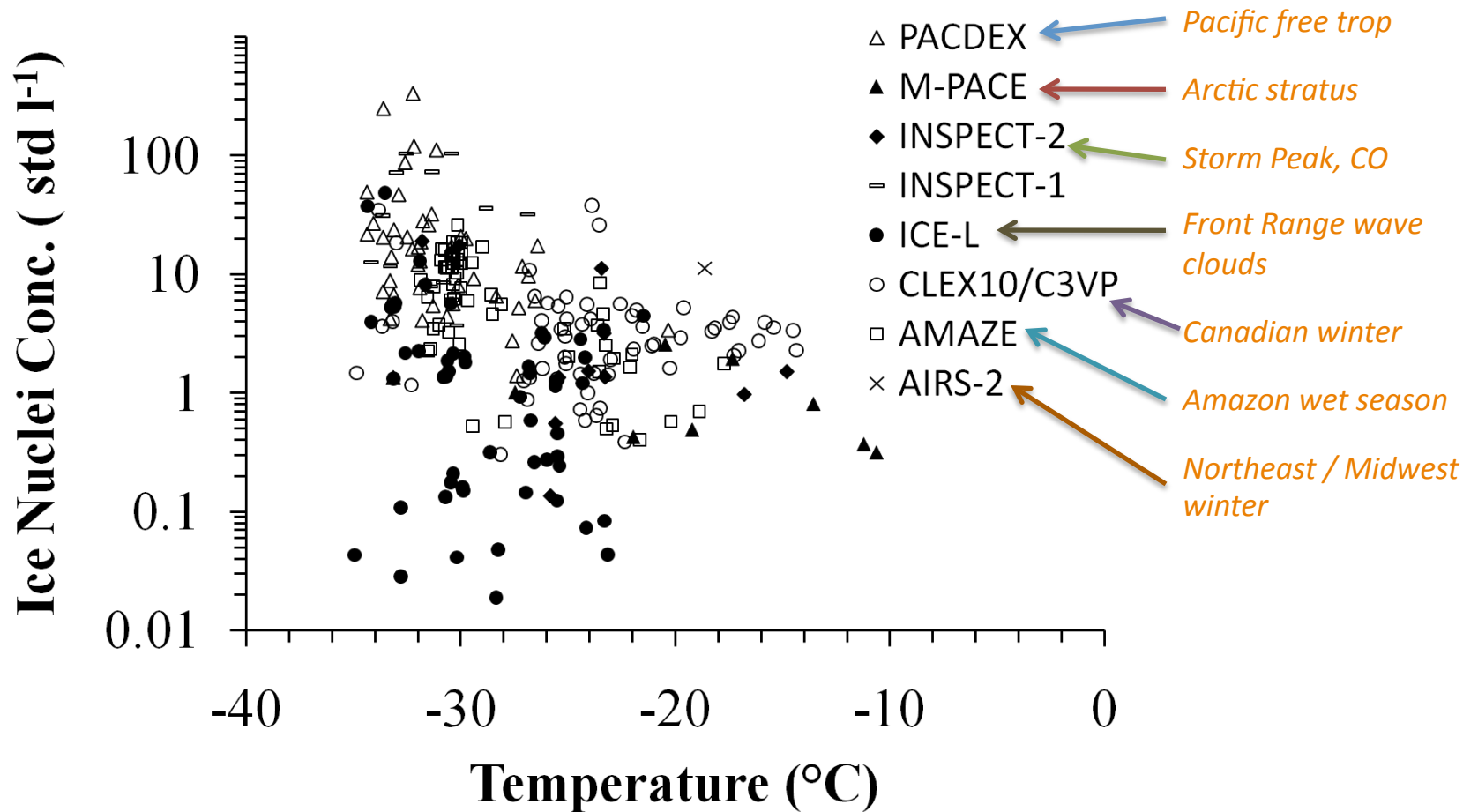
**Cloud ice based on 2D-C probe > 50 μm**

# Ice nuclei concentrations over several projects (10-30 min. averages)



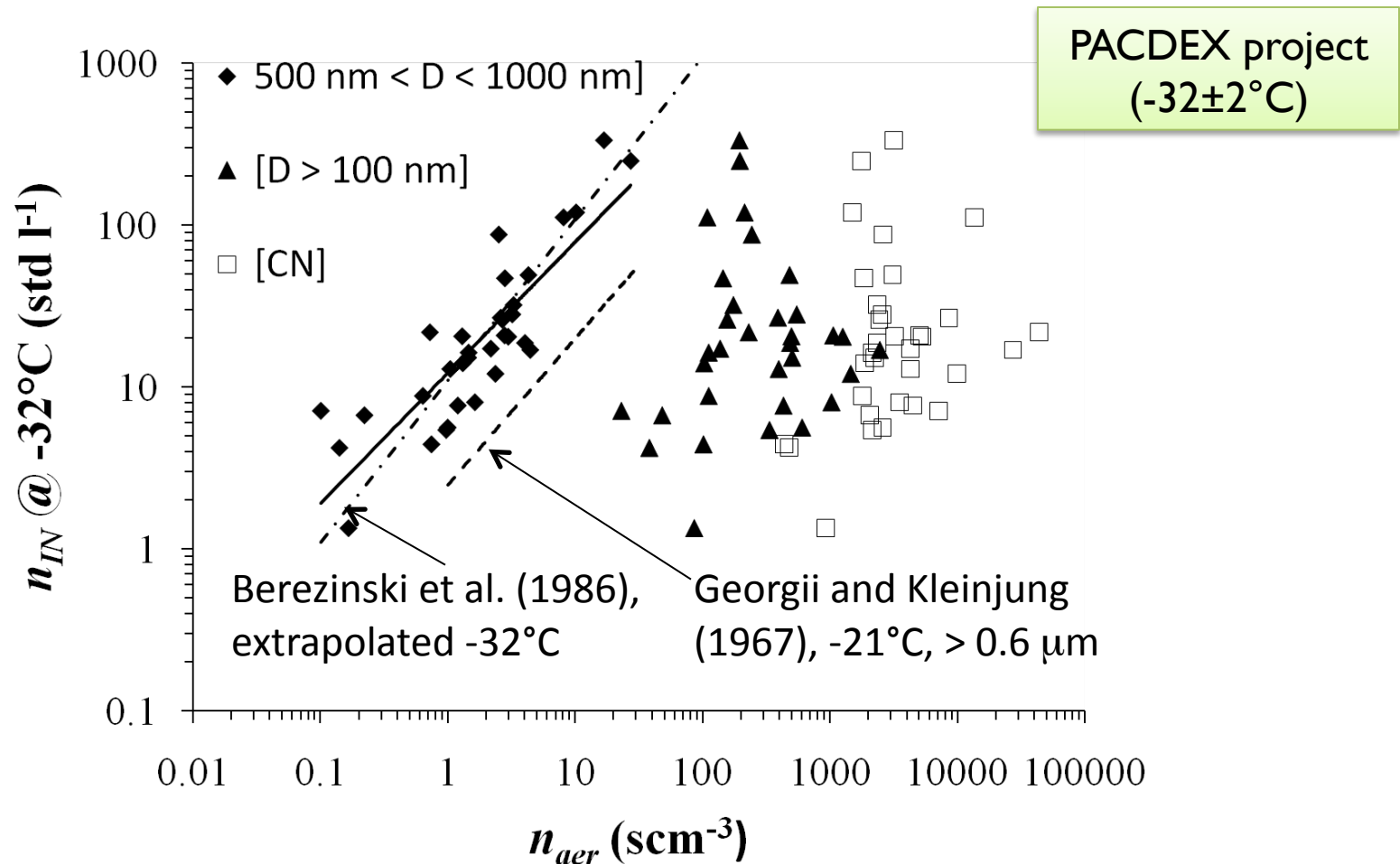
[DeMott et al., 2009]

# Ice nuclei concentrations over several projects (10-30 min. averages)



[DeMott et al., 2009]

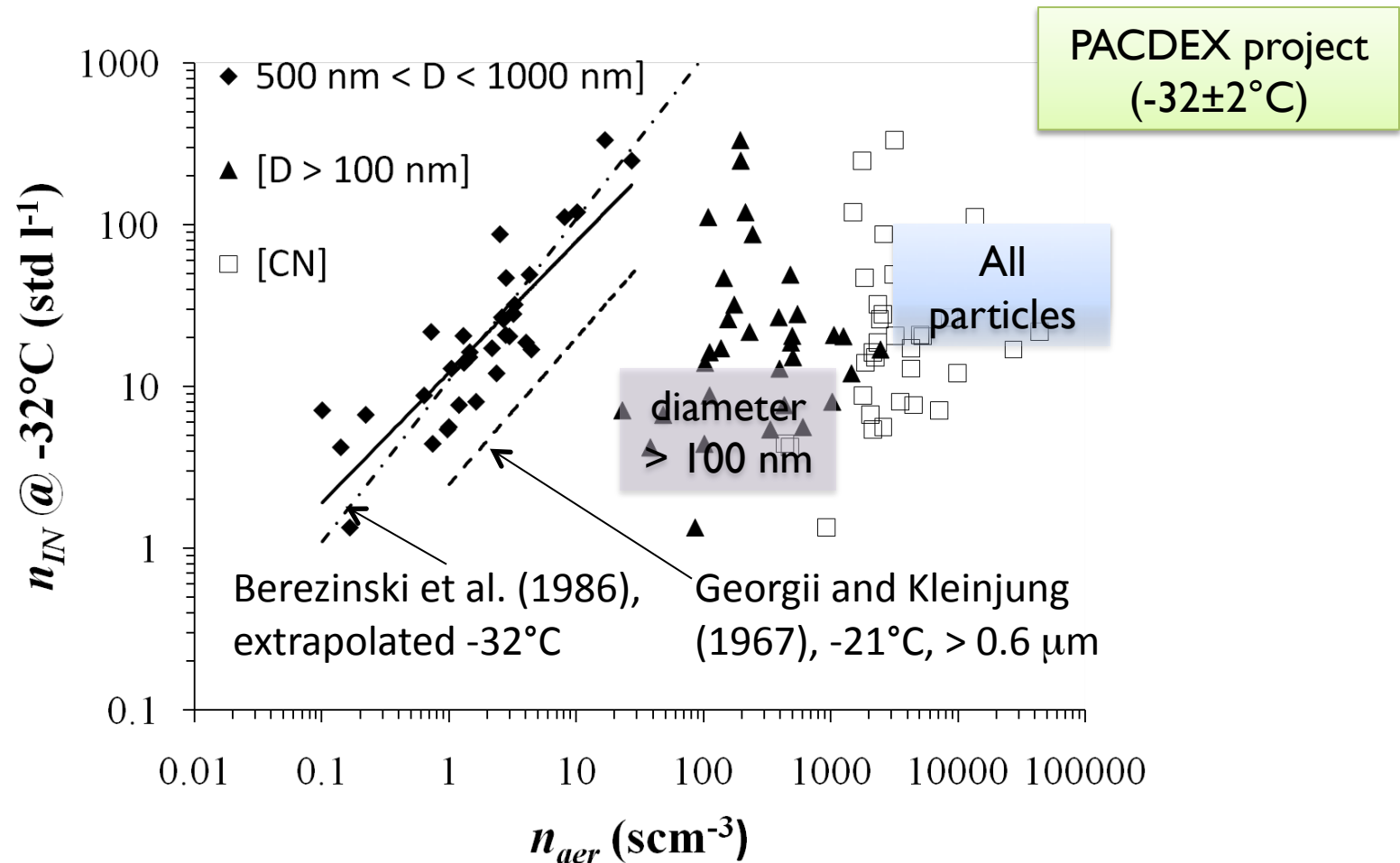
# IN trend with aerosol concentrations when stratified by **size** and **temperature**



DeMott et al. (2009)

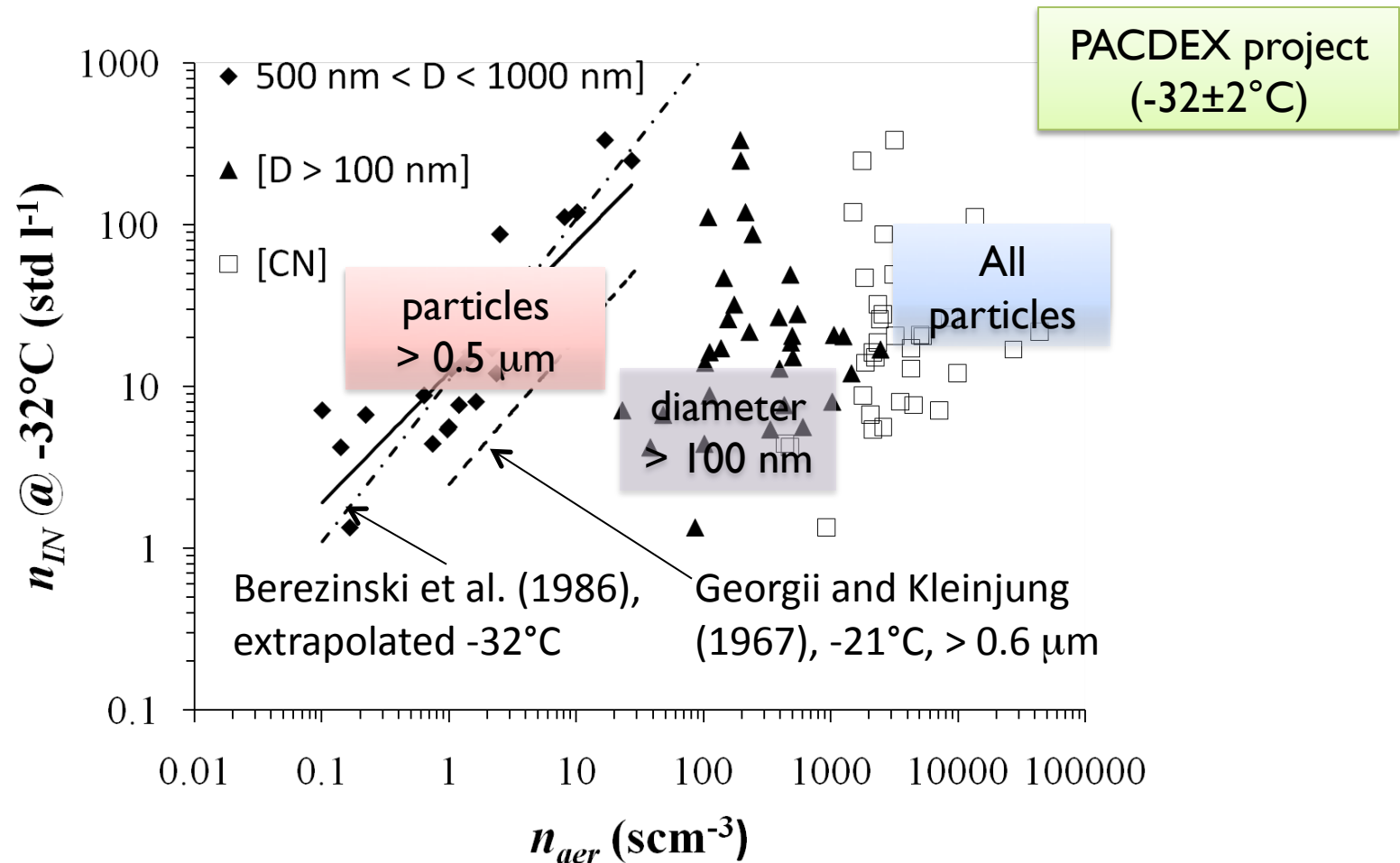


# IN trend with aerosol concentrations when stratified by **size** and **temperature**



DeMott et al. (2009)

# IN trend with aerosol concentrations when stratified by **size** and **temperature**



DeMott et al. (2009)

# Ice nucleation parameterizations

- Meyers et al. (1992):  $n_{in} = \exp(12.96(S_i - 1) - 0.639)$   
 (no links to aerosol properties)
 

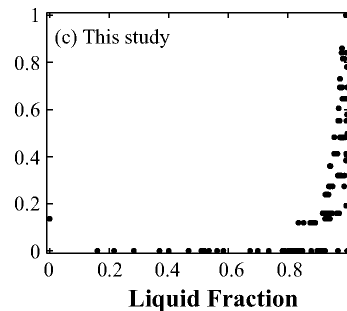
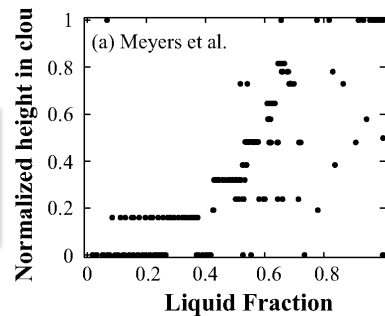
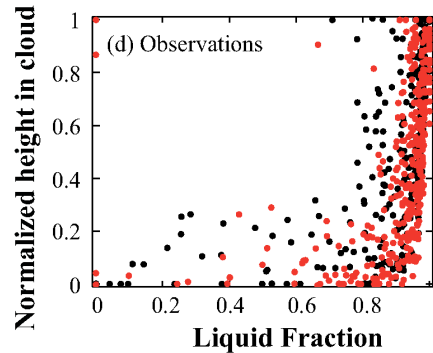
Ice supersaturation dependence only

- Phillips et al. (2008):  
 (surface area, **composition**,  $S_i$ ,  $T$ )
 
$$n_{IN,X} = \alpha_X H_X(S_{i,v}, T) \xi(T) \left( \frac{n_{IN,1.5,*}(T, S_{i,v})}{\Omega_{X,1.5,*}} \right) \Omega_X$$

$\alpha_X = f_{dust} f_{BC} f_{bio}$       Lab based corrections      Scaling to "baseline" IN conc. and sfc. area

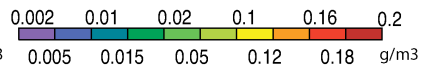
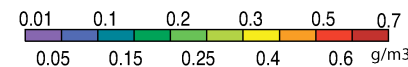
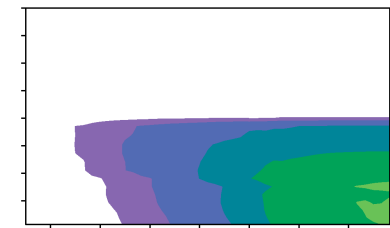
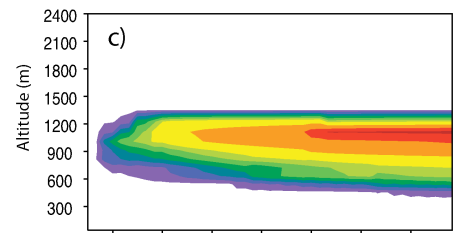
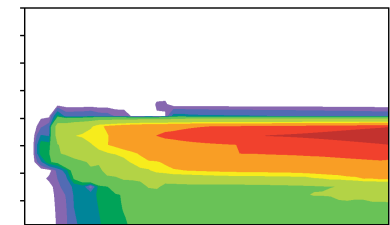
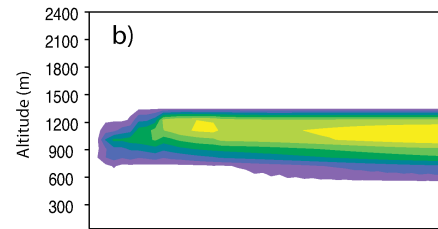
- DeMott et al. (2009):  $n_{IN,T_k} = a(273.16 - T_k)^{3.6434} (n_{aer,0.5})^{b(T_k)}$   
 ( $T, n_{aer} > 0.5 \mu\text{m}$  diameter)

# Regional impacts – Arctic stratus single column global model (SCAM3)



Cloud water content

Ice water content



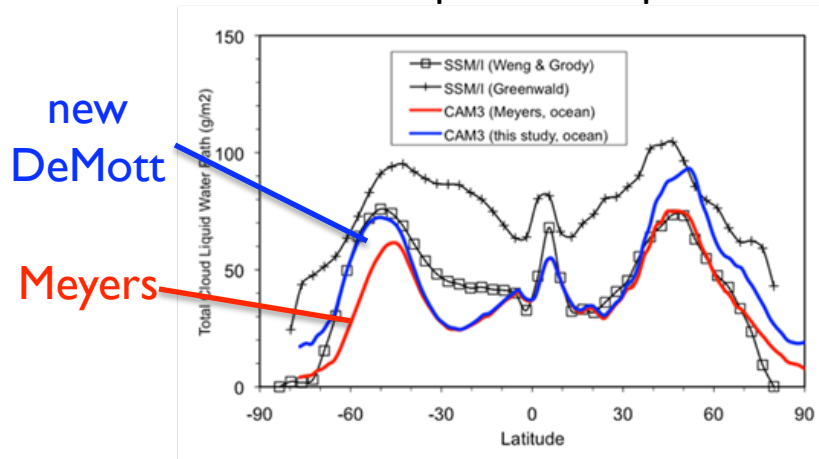
Liu et al. 2-moment  
microphys. + Meyers →

As above, BUT  
new IN param →

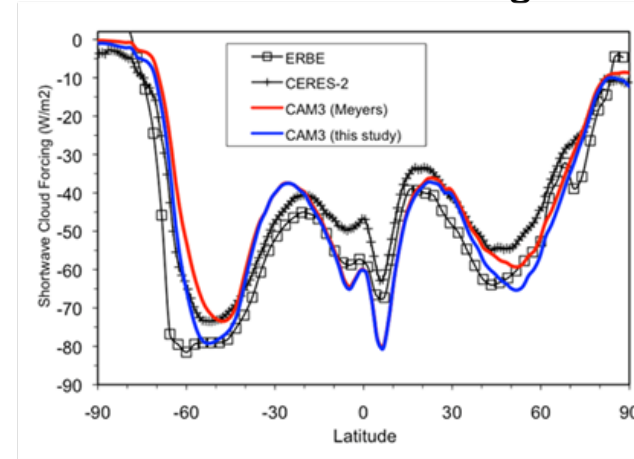
[DeMott et al., 2009]

# Global model (CAM3) 5-year simulations, annual averages

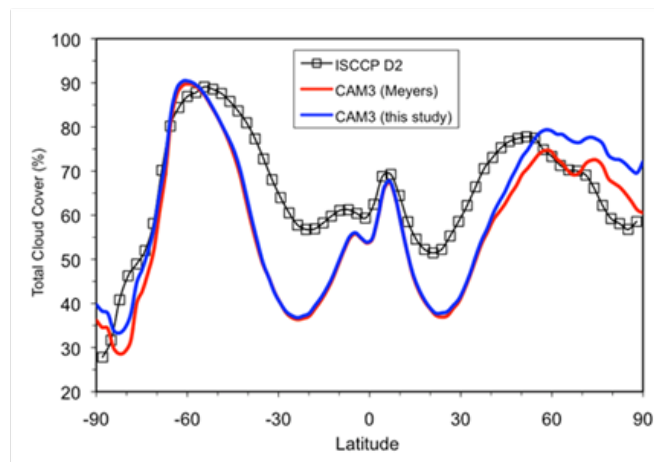
Total liquid water path



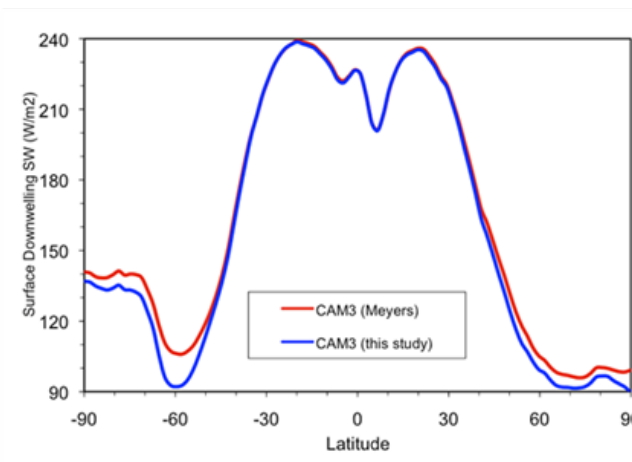
SW cloud forcing



Total cloud cover



Surface downwelling SW



# Summary

- IN measurements relate directly to first ice formation (clear from wave cloud studies, other studies where secondary ice processes can be separated)  
→ *important for predicting phase in many clouds!*
- IN concentrations in mixed-phase cloud T regime can be related to the number concentrations of particles larger than  $\sim 0.5 \mu\text{m}$   
→ *useful in models that carry some information on particle size, eventually particle type*
- Global model simulation sensitivity to IN formulation is quite strong  
→ *our new parameterization yields more water clouds and less ice, especially in Arctic & midlatitude storm tracks*

# Future work

- For CMMAP, implement the parameterization into the SAM model
  - Case studies for different locations
  - Use of CloudSat simulator to compare with obs
- Implementation in the MMF
  - Once aerosols are included!