

Wind-Driven Gyre Circulations and Boundary Currents

Surface Balance of Forces

- Wind **stress** accelerates surface water
- **Friction** couples surface to underlying water
- Friction always acts exactly **opposite current motion**
- Coriolis acceleration is always **perpendicular to current motion**

RESULT: Surface current directed about 45° (right/left) of wind in (NH/SH)

Ekman Flow

- Combined effects of Coriolis and friction on "stack" of thin layers
- Each layer moves more slowly and further right(left) than layer above ("**spiral hodograph**")
- Mass-weighted **mean motion is 90° cum sole of wind**

Ekman Pumping

NORTHERN HEMISPHERE

- Ekman flow in NH is 90° to the right of the wind stress
- Cyclonic wind forces divergence in water, and upwelling
- Anticyclonic wind forces convergence and downwelling

Idealized Geostrophic Gyre

- Convergence of Ekman flow raises sea surface
- Rotating "dome" results

Figure 3.24 The generation of geostrophic current flow in a gyre driven by anticyclonic winds in the Northern Hemisphere. This current is driven by the wind only indirectly and persists below the wind-driven (Ekman) layer.

Asymmetric Gyre

- Real world doesn't look like Sverdrup flow
- Boundary currents are not symmetric
- Why?

Figure 4.11 Schematic illustration of the asymmetrical North Atlantic gyre (blue) and the more or less symmetrical wind field which overlies it.

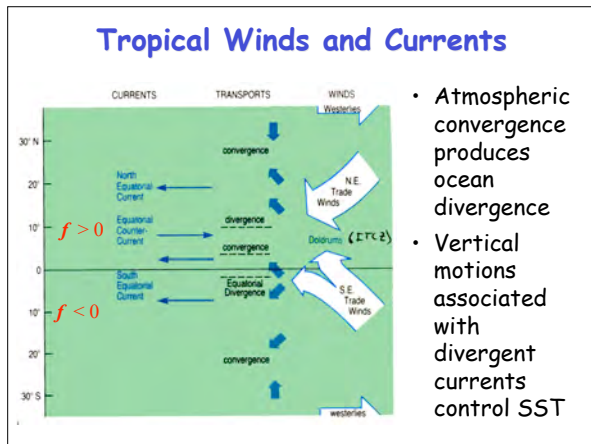
Eastern Boundary Currents (NH)

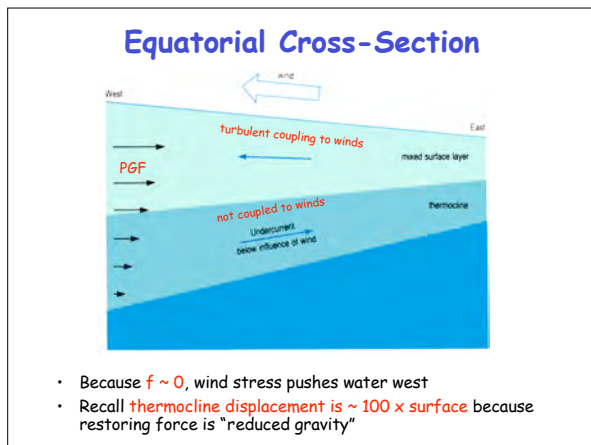
- Sloping sea-surface in upwelling regions produces **equatorward geostrophic current**
- Strongly **baroclinic conditions** eliminate and even reverse pressure gradient at depth (**countercurrent**)

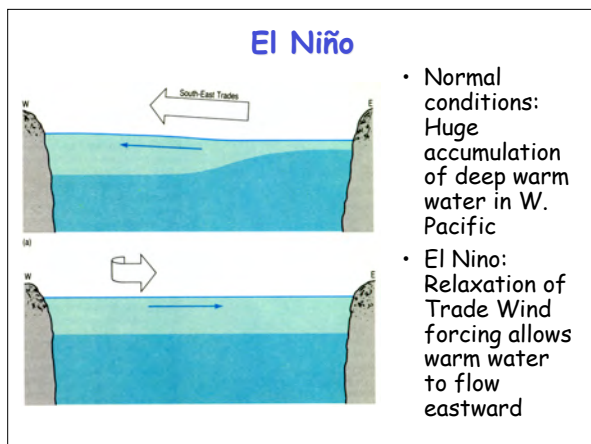
Coastal Upwelling Effect on SST

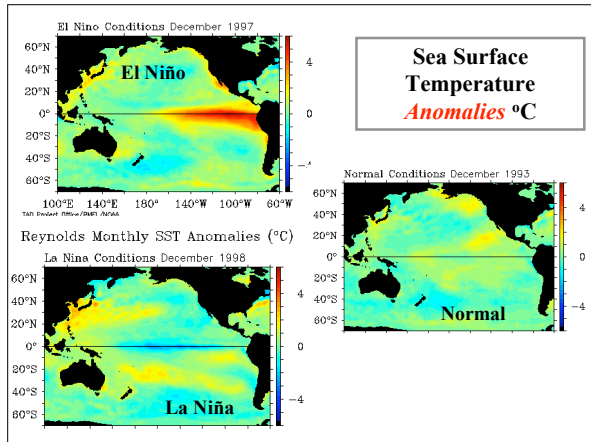
- Deviation from zonal mean SST
- Cold coastal water due to upwelling and cold geostrophic currents
- Interaction with desert climate on land

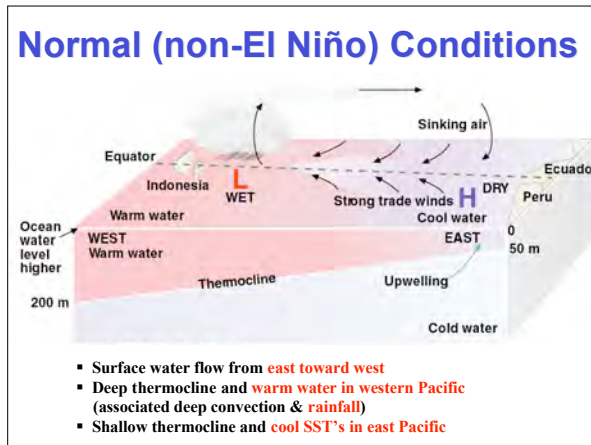
Tropical Oceans and El Niño

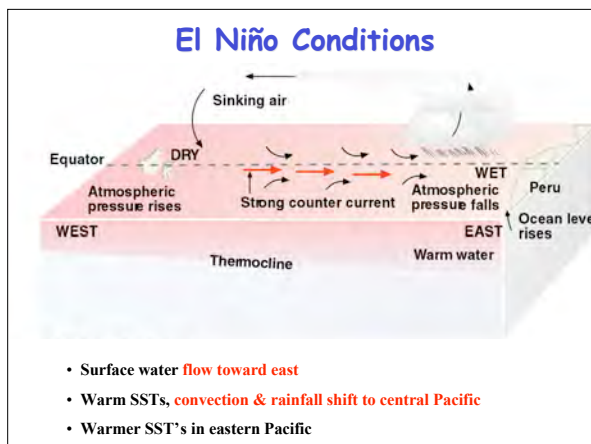


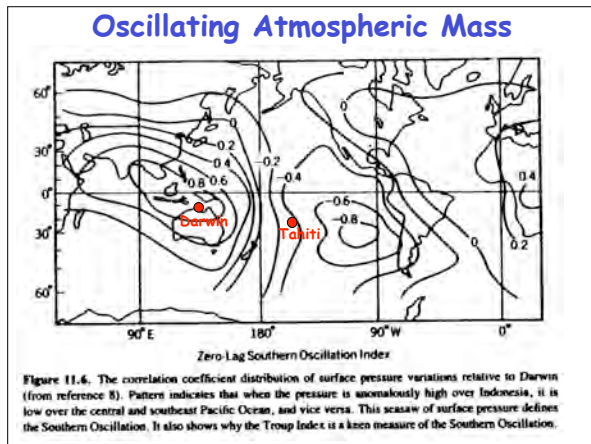


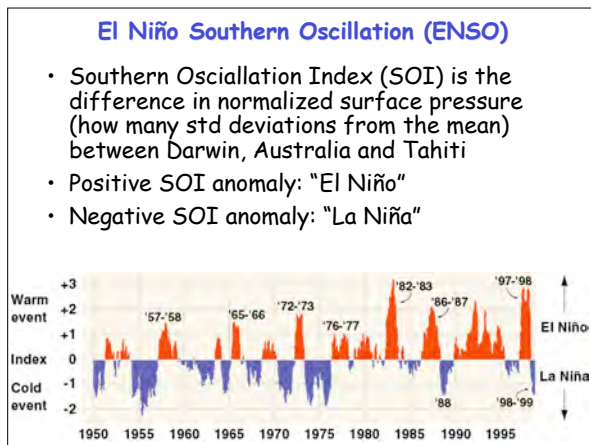






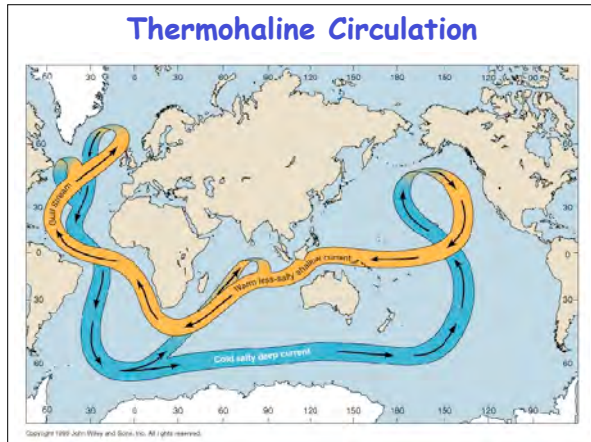


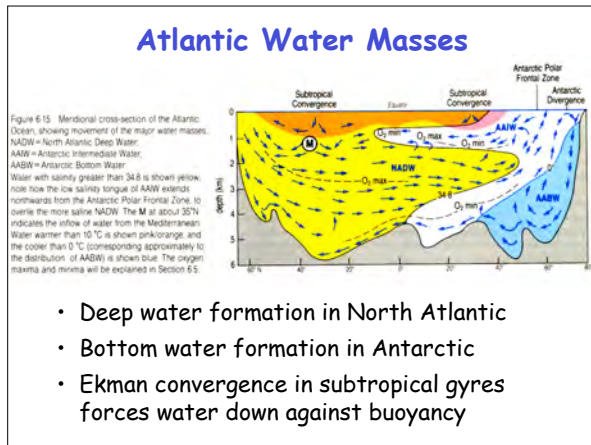


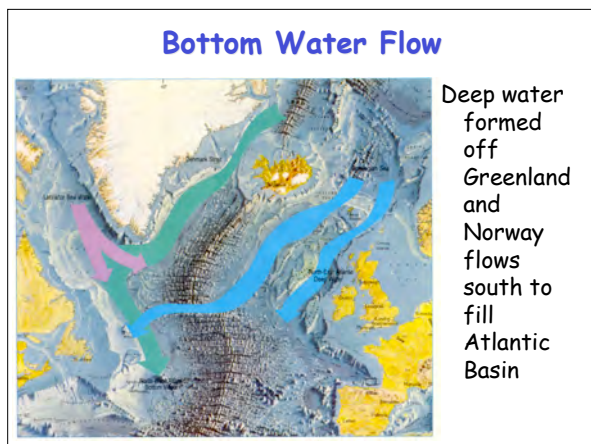


Thermohaline Circulation

Read article by W. Broecker

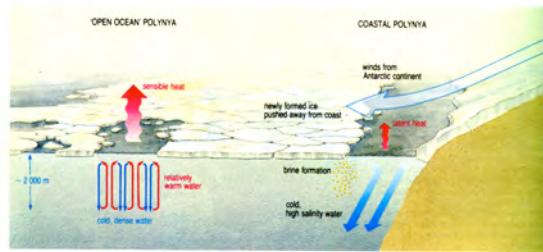






Formation of Antarctic Bottom Water

Figure 6.19 The different roles played by coastal and 'open ocean' polynyas in the production of Antarctic Bottom Water.



Continuous formation of ice along coast and in "leads" or "polynyas" forms extremely dense water

Thermohaline Heat Pump

- Upper limb **inflow to North Atlantic** $\sim 10^\circ \text{C}$
- Lower limb **outflow** $\sim 3^\circ \text{C}$
- $dQ = c \, dT \sim 3 \times 10^7 \text{ J}$ of heat released by each m^3 of water during conversion from upper limb to lower limb water mass
- $20 \text{ Sv} = 20 \times 10^6 \text{ m}^3 \text{ s}^{-1}$ of water makes this transition, releasing $6 \times 10^{14} \text{ J s}^{-1}$ ($= 0.6 \text{ Pw}$) of heat to the atmosphere
- This is **35% of solar heating** of North Atlantic north of 40°N latitude!

Simulated Thermohaline Collapse

